

## Space-time distribution of the Friuli (1976) earthquakes

V. CAGNETTI (\*) - R. CONSOLE (\*\*)

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RIASSUNTO. - Sono stati esaminati i risultati delle determinazioni ipocentrali delle scosse del periodo sismico iniziato il 6 maggio nel Friuli.

Una particolare considerazione è stata rivolta all'evento principale i cui parametri focali determinati per via analitica, sono stati confrontati con quelli determinati con metodi macrosismici.

L'evoluzione del fenomeno sismico è stata seguita con una rete di stazioni locali che ha permesso l'individuazione degli ipocentri di circa 700 repliche; il metodo utilizzato è stato quello basato sulle differenze  $T_{Sg} - T_{Pg}$  assumendo per il fattore  $K$  il valore di 7.5 km/sec.

Una valutazione degli errori sperimentali effettuata con un nuovo programma di calcolo ha fornito risultati compresi in  $\pm 2$  km.

Lo studio della migrazione degli ipocentri ha messo in evidenza la comparsa di attività sismica sia come premonitrice ad un evento di maggiore magnitudo sia come ad esso conseguente.

L'evoluzione cronologica del fenomeno è stata rappresentata con piante e sezioni limitate per opportuni intervalli di tempo.

I risultati ottenuti hanno evidenziato interessanti aspetti del periodo sismico che si ritiene possano incoraggiare il proseguimento di studi in questa direzione.

SUMMARY. - Results obtained by hypocentral determination of 700 shocks of the Friuli earthquake have been examined.

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(\*) C. N. E. N. - Laboratorio di Ingegneria del Sito-RAD-RS. C. S. N. Casaccia (Roma).

(\*\*) I. N. G. - Osservatorio Geofisico Centrale - Monteporzio Catone (Roma).

Special care has been devoted to the main shock. A comparison between focal parameters determined by analytical and macroseismic methods has been carried out.

The time and space evolution of the phenomenon has been followed by means of a network of local stations which allowed the determination of the hypocenters of about 700 aftershocks.

The analytical method based on  $S_g - P_g$  time interval has been employed, assuming a  $K$ -value of 7.5 km/sec.

An evaluation of experimental errors, performed by means of an original code, yielded an accuracy of  $\pm 2$  km.

The study of hypocenters migration shows seismic activity both to precede and to follow an event of greater magnitude.

Epicenter and hypocenter distributions are plotted in order to depict the time-evolution of the phenomenon.

Results show interesting features of the seismic period which, in our opinion, warrant the pursuit of this kind of investigation.

#### DESCRIPTION OF THE SEISMIC PERIOD.

The seismic period of 1976 in Friuli started, at least in its most remarkable phase, on the evening of May 6 at 19<sup>h</sup> 59<sup>m</sup> 03<sup>s</sup> (G.M.T.) with a shock having a magnitude  $M_L=4.5$  (RMP)\* (probably the only foreshock), followed at 20<sup>h</sup> 00<sup>m</sup> 12<sup>s</sup> by the catastrophic shock of magnitude  $M_L=6.3$  (RMP). The very numerous aftershocks which followed showed a distribution in time which is peculiar of the seismic events of high intensity and their distribution law as a function of the magnitude was characterized by a value of parametr  $b$  usual for the sites of the Mediterranean area.

Such trend, which appeared to be completely normal at the time, continued until the beginning of the month of September, when a renewal of the seismic activity was observed; this manifested itself on September 11 with the shock of magnitude 5.3 (RMP) at 16<sup>h</sup> 31<sup>m</sup> 12<sup>s</sup>, which was followed after only four minutes by the shock of magnitude 5.9(RMP) at 16<sup>h</sup> 35<sup>m</sup> 00<sup>s</sup>. Also in this case the aftershocks showed a relatively rapid decay.

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\* determined at the Central Observatory of the National Institute of Geophysics, Rome.

On September 15, at 03<sup>h</sup> 15<sup>m</sup> 49<sup>s</sup> a new violent increase of the seismic activity was observed with the shock of magnitude  $M_L=6.1$  (RMP). This new activity reached a new maximum at 09<sup>h</sup> 21<sup>m</sup> 18<sup>s</sup> with the earthquake of magnitude  $M_L=6.0$  (RMP). Since that time and until March 1977, the seismic activity showed, generally, a descending feature; however, renewals both in the frequency of shocks and in their intensity were observed sometimes.

#### SOME CONSIDERATIONS ON THE MAIN SHOCK OF MAY 6.

The hypocentral determination of the main shock, obviously carried out from the data coming from the pre-existing seismographic stations, was difficult because of the overlapping of the first pulse to the pulse of the foreshock.

In any case to carry out this determination a calculation programme based on the times of arrival of the first pulse from the different stations was used; the travel times experimentally derived by prof. Caloi (<sup>1</sup>) in the study of the earthquake of Cansiglio (1936) were used.

The result was the following (<sup>2</sup>):

$$\begin{array}{ll}
 H=20^h 00^m 11.8^s & \pm 0.4^s \\
 \varphi=46.266^{\circ} N & \pm 0.027^{\circ} \\
 \lambda=13.250^{\circ} E & \pm 0.029^{\circ} \\
 h=20.1 \text{ km} & \pm 3.8 \text{ km.}
 \end{array}$$

It is necessary to make some considerations on the errors involved in the determination technique.

A first cause of error is related to the error with which the arrival time of the *P* wave, dependent on the velocity of the recording system and on the frequency response of the instrument, is read. Moreover, it should be taken into consideration that the phase considered as the first one in the seismogram does not always correspond to the actual first pulse coming from the source. This may be more especially true for those stations receiving the wave coming from the nodal plane and for those which are situated at particular distances from the epicenter. Furthermore, one should carefully consider the possibility that the nearer stations, that is to say more sensitive, could record as

a first pulse a phase which is otherwise mixed up with the noise in the recording of the farther, that is to say less sensitive, stations.

An evaluation of these errors, even if rather approximate, suggests to consider them within the confidence margins obtained with the least square method, previously discussed.

As far as the hypocentral depth is concerned, the greatest error which can be made in its determination is that of assuming a model of the propagation velocities rather different from reality. This systematic error is very difficult to estimate; several tests carried out using different models resulted in a wide variability for the depth, while, on the contrary, the position of the epicenter remained constant <sup>(3)</sup>.

The comparison of the result thus obtained with the results obtained with the method of macroseismic investigation will be shortly discussed.

The epicentral determination is generally made considering the epicenter as the center of the most damaged area. In this case this was found in an area between Gemona, Trasaghis and Osoppo. The discrepancy between the analytical epicenter and the macroseismic epicenter is only apparent: in fact their definition is essentially different. The analytical epicenter is on the vertical line from the spot where the fracture started. The macroseismic epicenter, located according to the damages caused to the buildings, is normally on the vertical line from the spot of the focal volume nearest to the surface and can, furthermore, be affected by particular conditions of the surface geology.

The focal depth, as it is well known, can be calculated from the distance between isoseismals of different degrees. Generally, the well known formula is used <sup>(7)</sup>:

$$I_0 - I = \nu \log \left( 1 + \frac{R^2}{h^2} \right)^{0.5} \quad [1]$$

In this case the formula gave the value of  $h = 11.8$  km.

Another method to determine the focal depth is based on the Karnik's formula <sup>(5)</sup>:

$$M = 0.66 I_0 + 1.7 \log h - 1.6 \quad [2]$$

which correlates the magnitude, the epicentral intensity and the focal depth. Assuming  $I_0 = IX - X$  MSK and  $M = 6.3$ ,  $h$  resulted = 9.1 km.

This depth is that of the part of the focal volume where the energy release contributed for the greatest amount in causing the damages at the surface.

#### HYPOCENTRAL DETERMINATION OF THE AFTERSHOCKS.

The scientific interest and the remarkable social aspects of the event of May 6 suggested the immediate intervention of three mobile stations of I. N. G. and three stations of C. N. E. N. These stations operated in the area, in cooperation with the stations already existing at dams of E. N. EL. (2).

The development in time of the seismic phenomenon was followed continuously by the local seismic stations, which sometimes were displaced to improve their operation (6).

As known, the determination of the hypocenters can be made, using the data of local stations, through two calculation techniques substantially different. These techniques can be reduced to the use of the only first arrivals, or to the intervals between the arrivals of the longitudinal and transversal waves. Without going details concerning the two methods, it should be observed that the first method requires an accurate determination of the absolute first-arrival times or the availability of a telemetric seismic network. The second one, on the contrary, can leave out of consideration the absolute measurement of the time, but requires an accurate interpretation of the seismogram.

The difficulty involved in the synchronization of the clocks used in the network (due to the bad reception of the time-signals, and to the displacements of the stations) suggested the use of the classic method based on the differences  $T_{Sg} - T_{Pg}$ . Such time is proportional to the hypocentral distance for the recording station, through the factor

$$K = \frac{V_{Pg} \cdot V_{Sg}}{V_{Pg} - V_{Sg}} \quad [3]$$

An evaluation of the parameter  $K$  can be made experimentally when a sufficient number of stations recording the same events (at least 4) is available. The tests carried out in this sense resulted in a mean valued of  $K = 7.5$  km/sec. This value was successively used for all the hypocentral determinations.

Taking into consideration the possibility that the parameter  $K$  could be not perfectly known in the area under examination, various tests were carried out to evaluate how these variations could affect the results obtained. It was thus confirmed that, for a variability range of the  $K$  factor between the values of 7 and 8 km/sec, the scattering of the hypocentral coordinates was included in the variations related to the errors in the readings of times.

#### ERRORS IN THE $S-P$ HYPOCENTRAL DETERMINATION METHOD.

The error in the results of the calculations giving hypocentral coordinates derives from the uncertainty with which the experimental data are read and depend on the number of stations and on their localization with regard to the hypocenter.

The most important difficulty in the reading of the time differences  $S-P$  can be attributed to the detection by the operator of the exact arrival of the  $S$  phase in the seismogram.

Consequently to minimize this cause of error, all the readings of seismograms for which interpretation doubts existed, were rejected. Moreover, in the successive phases of the elaboration, all the data with large anomalous residues were eliminated; also the analytically impossible solutions were not taken into consideration.

On readings without interpretation doubts, in the case of recording speeds of 60 mm/min, the possible error was evaluated at a value of about  $\pm 0.2$  sec. Such value which we considered as the maximum value, is undoubtedly higher than that obtained with magnetic tape reproductions, as it is the case for some stations.

The consequences deriving in the detection of the hypocenter from such reading error, have been carefully analysed through a calculation programme on purpose prepared. The results, rather interesting, evidenced the dependence of the error on the relative position of the stations with regard to the hypocenter of the shock.

When the stations are located all around the hypocenter, constituting a network the size of which is of the same order of magnitude of the hypocentral depth, the error on the hypocentral coordinates is less than 2 km, and the error on the depth is lower. On the contrary, if the distance among the stations is remarkably lower than the hy-

poentral depth, the error on the depth is still of the same order of magnitude, while the error on the epicentral localization increases.

In the case of a network composed of stations remarkably more distant than in the previous cases, the epicentral position remains relatively sure in its assessment, while the depths results rather uncertain.

The most unfavourable case is that of a network not surrounding the hypocenter; the stations supply data showing a remarkable scattering both of the epicentral coordinates and of the depth.

#### MIGRATION OF THE HYPOCENTERS.

Interpreting the seismograms recorded by the local stations, it was observed that the values of the hypocentral distances were constant for short periods of time. This observation suggested us to study the possible duration of the seismic activity in given zones of the interested area. For this purpose a calculation programme was adopted, which supplied the graphs showing the location of the hypocenters of the shocks for particular periods of time.

Often the seismograms of a group of shocks, recorded in the same station, showed important analogies.

First of all the epicentral positions of the shocks included between two events with high magnitude were plotted on maps. However, these graphs, though giving a clear representation of the migratory phenomenon, did not completely describe its complex trend.

Successively, the area interested by the seismic phenomenon was subdivided into circular zones having an adequate radius. In most cases the circular zones were centered for the most part on the epicentral positions of the shocks having higher magnitude. Thus the events occurred within the examined zones were evidenced in chronological order. The results obtained showed the appearance of the seismic activity in zones previously inactive, either as forerunners signs of more important events, or as phenomena following such events (see Table 1).

The periods of time characterizing such phenomena supplied us with a subdivision of the events to be used for further maps of the epicenters. With reference to the same periods of time,  $N-S$  sections were also prepared, where the positions of the epicenters are represented.

The maps of the epicenters, obtained from a computer printout on an original scale of 1: 100,000, are shown in figures 1a-26a.

In these maps each event is represented by a number indicating the depth in discrete classes of two kilometers. For instance, the number 0 indicates a depth between 0 and 2 km, and so on: number 9 indicates a depth higher than 18 km. When the number is accompanied by an asterisk, this indicates an event of magnitude equal or higher than 4.

The sections, also derived from an original scale of 100,000, are shown in figures 1b-26b. In these figures the epicenters are indicated with little crosses when the magnitude is lower than 4, with asterisk when the magnitudes are higher.

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\* obtained by maximum amplitudes read on a seismograph previously calibrated with the standard « Wood-Anderson » instruments.

## LIST OF THE STUDIED SHOCKS.

## Meaning of the symbols.

<i>M</i>	= month
<i>D</i>	= day
<i>T</i>	= origin time in hours, minutes, seconds ( <i>GMT</i> )
<i>ML</i>	= local magnitude * (the value 0 corresponds to $M_l < 2.5$ )
<i>LAT</i>	= latitude in degrees and thousandths, of the epicenter
<i>LONG</i>	= longitude in degrees and thousandths, of the epicenter
<i>H</i>	= hypocentral depth in km
1	= zone centered on the point with coordinates 46.380 N 13.220 E
2	= zone centered on the point with coordinates 46.380 N 13.120 E
3	= zone centered on the point with coordinates 46.340 N 13.000 E
4	= zone centered on the point with coordinates 46.310 N 12.210 E
5	= zone centered on the point with coordinates 46.290 N 13.100 E
6	= zone centered on the point with coordinates 46.250 N 12.950 E
7	= zone centered on the point with coordinates 46.240 N 13.270 E
8	= zone centered on the point with coordinates 46.220 N 13.250 E
9	= zone centered on the point with coordinates 46.250 N 13.250 E

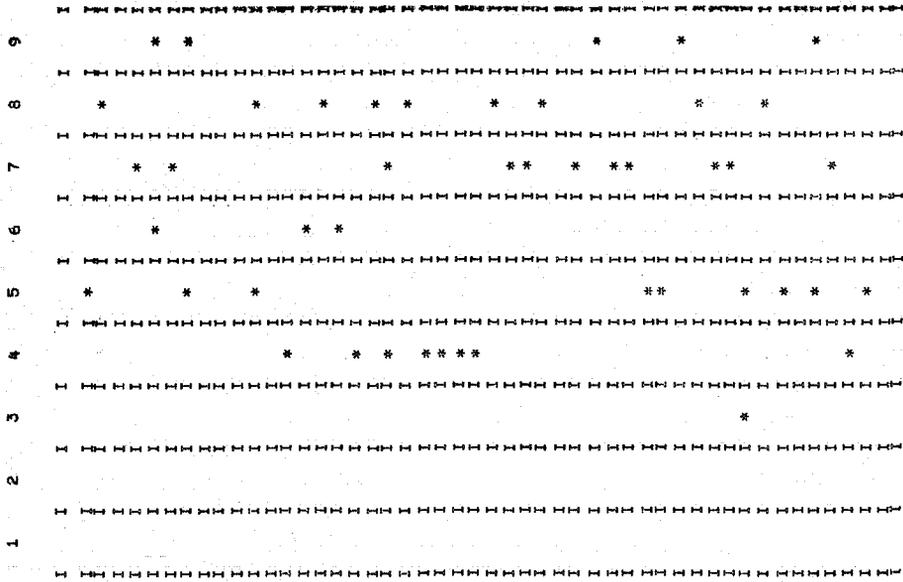
The asterisk means that the event has an epicenter included in the circle with 5 km radius, centered on the point respectively indicated.







M D	T.	ML	LAT.	LONG.	H
5 15	0 39 55	1.9	46.335	13.117	14.7
5 15	1 16 53	2.4	46.365	13.098	7.2
5 15	2 33 12	2.4	46.347	13.146	9.1
5 15	4 26 16	4.2	46.341	13.338	1.0
5 15	8 40 17	3.9	46.333	13.270	15.8
5 15	9 29 39	3.1	46.209	13.986	10.1
5 15	9 29 8	3.4	46.363	13.239	12.7
5 15	9 40 39	3.4	46.347	13.053	15.9
5 15	9 42 54	3.1	46.336	13.190	6.8
5 15	10 26 5	2.6	46.382	13.103	18.3
5 15	11 19 43	2.5	46.347	13.200	6.4
5 15	11 19 43	2.5	46.347	13.116	8.3
5 15	12 13 32	2.5	46.288	13.377	13.4
5 15	14 21 24	3.0	46.259	13.377	13.1
5 15	15 17 24	3.0	46.275	12.942	10.1
5 15	15 25 10	3.3	46.207	13.131	18.8
5 15	16 5 58	3.4	46.213	12.971	18.5
5 15	16 50 48	3.1	46.284	13.210	9.7
5 15	16 58 3	2.4	46.249	13.156	12.4
5 15	18 37 8	3.3	46.275	13.235	8.9
5 15	19 20 16	2.9	46.214	13.142	3.6
5 15	20 13 56	2.7	46.330	13.202	7.0
5 15	20 43 4	2.7	46.278	13.221	7.4
5 15	21 10 26	2.8	46.324	13.238	1.0
5 15	22 34 20	2.8	46.290	13.113	1.0
5 16	0 32 15	2.9	46.232	13.136	6.5
5 16	2 13 55	2.5	46.225	13.327	1.0
5 16	2 37 16	2.5	46.247	13.268	5.1
5 16	4 49 50	2.6	46.237	13.115	7.5
5 16	4 57 5	3.1	46.193	13.216	7.7
5 16	4 58 20	2.8	46.252	13.332	1.0
5 16	5 58 48	2.6	46.235	13.064	17.8
5 16	5 59 23	3.1	46.252	13.255	12.3
5 16	6 28 7	3.0	46.242	13.233	13.8
5 16	7 44 39	3.0	46.312	13.144	18.4
5 16	8 0	3.4	46.333	13.104	18.4
5 16	10 50 17	2.6	46.260	13.039	11.7
5 16	11 56 8	2.5	46.247	13.169	11.0
5 16	12 56 18	3.0	46.272	13.182	7.4
5 16	13 48 27	2.6	46.256	13.325	1.0
5 16	20 3 26	2.6	46.314	13.049	9.8
5 16	20 13 42	2.8	46.259	13.172	9.7
5 16	21 32 0	2.7	46.296	13.093	13.3
5 17	2 55 40	2.9	46.294	13.016	9.2
5 17	9 11 33	2.7	46.273	13.068	16.4
5 17	13 30 5	2.7	46.245	13.286	11.5
5 17	15 23 51	2.4	46.304	13.203	11.8
5 17	16 13 16	4.3	46.296	13.069	1.0
5 17	16 21 27	3.3	46.282	12.987	14.7
5 17	16 23 17	3.3	46.283	12.987	14.7





M D	T	ML	LAT.	LONG.	H	1	2	3	4	5	6	7	8	9
5 22	16 26 0	3.0	46.254	13.114	7.8	I	I	I	I	I	I	I	I	I
5 23	20 30 37	2.4	46.221	13.158	18.7	I	I	I	I	I	I	I	I	I
5 23	0 51 10	4.0	46.228	13.212	12.4	I	I	I	I	I	I	I	I	I
5 23	1 25 49	2.8	46.206	13.145	16.6	I	I	I	I	I	I	I	I	I
5 23	1 53 9	3.0	46.252	13.238	18.2	I	I	I	I	I	I	I	I	I
5 23	3 57 38	2.0	46.254	13.259	10.1	I	I	I	I	I	I	I	I	I
5 23	5 16 8	2.2	46.251	13.117	7.9	I	I	I	I	I	I	I	I	I
5 23	6 1 13	2.7	46.289	13.071	2.8	I	I	I	I	I	I	I	I	I
5 23	6 33 38	2.4	46.224	13.160	1.3	I	I	I	I	I	I	I	I	I
5 23	11 18 46	2.3	46.224	13.160	1.3	I	I	I	I	I	I	I	I	I
5 23	14 1 42	2.0	46.259	13.259	1.5	I	I	I	I	I	I	I	I	I
5 23	23 0 15	2.3	46.259	13.168	10.5	I	I	I	I	I	I	I	I	I
5 24	23 40 14	2.3	46.266	13.107	10.1	I	I	I	I	I	I	I	I	I
5 24	2 38 20	2.3	46.191	13.815	6.8	I	I	I	I	I	I	I	I	I
5 24	4 52 15	2.0	46.317	13.052	11.3	I	I	I	I	I	I	I	I	I
5 24	4 57 26	2.6	46.212	13.474	5.5	I	I	I	I	I	I	I	I	I
5 24	4 7 30 23	2.6	46.241	13.982	9.7	I	I	I	I	I	I	I	I	I
5 24	11 16 43	2.4	46.342	13.063	1.0	I	I	I	I	I	I	I	I	I
5 24	10 37 51	1.3	46.298	13.201	11.6	I	I	I	I	I	I	I	I	I
5 25	0 43 59	3.2	46.258	13.337	17.0	I	I	I	I	I	I	I	I	I
5 25	1 24 35	2.0	46.246	12.911	17.6	I	I	I	I	I	I	I	I	I
5 25	3 32 47	2.3	46.273	13.189	10.0	I	I	I	I	I	I	I	I	I
5 25	4 55 28	1.3	46.256	13.093	5.0	I	I	I	I	I	I	I	I	I
5 25	12 50 9	2.8	46.221	13.110	6.8	I	I	I	I	I	I	I	I	I
5 25	16 25 24	2.6	46.252	13.172	8.7	I	I	I	I	I	I	I	I	I
5 25	19 4 28	2.0	46.205	13.205	1.0	I	I	I	I	I	I	I	I	I
5 25	19 47 59	3.4	46.260	13.096	11.6	I	I	I	I	I	I	I	I	I
5 25	20 21 40	2.7	46.211	12.989	5.2	I	I	I	I	I	I	I	I	I
5 25	22 21 17	3.3	46.280	13.037	7.9	I	I	I	I	I	I	I	I	I
5 25	22 52 37	2.2	46.188	13.132	7.1	I	I	I	I	I	I	I	I	I
5 26	4 39 9	2.2	46.288	13.185	11.9	I	I	I	I	I	I	I	I	I
5 26	7 48 27	2.2	46.371	13.228	14.1	I	I	I	I	I	I	I	I	I
5 26	9 46 25	2.6	46.337	13.213	17.4	I	I	I	I	I	I	I	I	I
5 26	9 59 17	2.5	46.255	13.124	19.2	I	I	I	I	I	I	I	I	I
5 26	11 20 35	3.0	46.366	13.166	13.0	I	I	I	I	I	I	I	I	I
5 26	13 53 11	2.9	46.372	13.252	9.6	I	I	I	I	I	I	I	I	I
5 26	14 11 13	2.2	46.342	13.052	14.0	I	I	I	I	I	I	I	I	I
5 26	14 39 19	2.8	46.258	13.244	9.5	I	I	I	I	I	I	I	I	I
5 26	19 29 0	3.0	46.214	13.138	6.3	I	I	I	I	I	I	I	I	I
5 26	22 6 59	2.3	46.224	13.136	9.1	I	I	I	I	I	I	I	I	I
5 27	3 49 47	3.4	46.308	13.198	7.1	I	I	I	I	I	I	I	I	I
5 27	6 50 20	2.7	46.240	13.294	12.9	I	I	I	I	I	I	I	I	I
5 27	10 3 58	2.5	46.262	13.170	8.4	I	I	I	I	I	I	I	I	I
5 28	10 15 15	2.2	46.379	13.173	8.1	I	I	I	I	I	I	I	I	I
5 28	22 11 29	2.3	46.316	13.110	10.7	I	I	I	I	I	I	I	I	I
5 28	23 13 39	2.5	46.309	13.056	11.1	I	I	I	I	I	I	I	I	I
5 28	23 38 43	2.5	46.211	13.058	11.0	I	I	I	I	I	I	I	I	I

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M D	Y	ML	LAT.	LONG.	H
5 29	1 34 0	2.1	46.233	13.122	7.6
5 29	4 3 8	1.8	46.205	13.104	6.4
5 29	12 57 45	3.1	46.345	13.089	12.9
5 29	14 57 45	3.0	46.334	13.001	10.2
5 29	21 13 13	3.8	46.260	13.249	1.0
5 29	21 13 42	3.8	46.250	13.105	7.6
5 30	7 23 45	3.7	46.298	13.148	18.0
5 31	4 33 49	3.4	46.272	12.910	4.4
6 1	5 42 51	2.5	46.261	13.118	10.2
6 1	5 52 50	2.3	46.286	12.957	8.3
6 1	7 56 29	2.4	46.383	13.165	18.1
6 1	10 4 0	2.6	46.324	13.122	7.5
6 1	15 46 8	3.3	46.292	13.183	12.2
6 1	17 21 11	4.2	46.221	12.860	1.0
6 1	19 16 4	3.5	46.253	13.127	13.4
6 1	21 29 30	2.8	46.312	13.183	1.0
6 2	9 27 2	3.6	46.263	12.565	12.7
6 2	11 56 37	2.9	46.270	13.116	15.5
6 3	13 44 50	3.0	46.230	13.260	14.2
6 3	15 44 45	3.1	46.241	13.118	10.8
6 3	19 23 0	3.7	46.267	13.015	7.8
6 3	19 35 55	3.0	46.170	13.057	13.6
6 3	20 22 10	3.2	46.282	13.169	17.7
6 3	21 39 20	2.4	46.255	13.042	11.1
6 3	23 15 10	2.7	46.255	12.932	1.0
6 4	2 49 37	2.5	46.272	13.058	13.3
6 4	7 49 15	3.5	46.262	12.939	7.4
6 4	11 56 58	3.5	46.266	13.017	15.0
6 6	18 12 36	3.3	46.264	13.069	5.4
6 6	18 30 59	3.2	46.264	12.988	8.0
6 6	18 37 59	3.2	46.270	13.073	8.8
6 7	3 34 19	2.5	46.281	13.137	10.3
6 7	3 15 8	2.6	46.338	13.203	1.0
6 7	12 35 21	2.8	46.212	12.952	5.6
6 7	18 18 22	2.4	46.272	12.922	13.7
6 7	22 31 0	2.8	46.253	13.005	13.2
6 7	23 5 33	2.6	46.248	12.938	10.6
6 8	4 33 32	3.6	46.342	13.102	18.2
6 8	4 29 29	3.6	46.307	13.122	18.1
6 8	6 34 23	3.0	46.298	13.116	18.1
6 8	12 19 27	3.3	46.224	12.605	5.9
6 8	12 19 27	3.3	46.301	13.162	17.2
6 8	12 40 55	2.8	46.291	13.320	1.0
6 9	18 48 14	4.2	46.260	13.000	13.0
6 9	19 12 17	3.2	46.240	12.959	9.0
6 10	0 15 15	3.2	46.263	12.974	11.0













M D	T	ML	LAT.	LONG.	H
9 26	9 40 8	3.1	46.313	13.187	4.4
9 26	13 28 3	2.5	46.158	13.235	11.7
9 26	14 52 23	3.7	46.287	13.228	3.6
9 26	17 50 27	2.6	46.307	13.086	6.5
9 27	14 37 31	3.5	46.261	13.145	14.0
9 27	17 52 23	3.2	46.331	12.993	8.8
9 28	0 10 14	2.7	46.297	13.227	10.5
9 29	5 25 47	0.0	46.333	13.094	3.6
9 29	19 15 30	0.0	46.400	13.175	1.8
9 29	21 53 41	0.0	46.433	13.177	1.0
9 29	22 6 12	0.0	46.427	13.098	1.0
9 29	22 39 19	0.0	46.310	13.135	9.9
9 29	22 40 3	0.0	46.261	13.112	9.8
9 30	22 47 18	0.0	46.351	13.113	1.0
9 30	9 27 17	0.0	46.382	13.080	1.0
9 30	9 27 46	2.5	46.340	13.042	6.9
9 30	19 43 43	0.0	46.330	13.112	7.8
9 30	19 43 23	0.0	46.281	13.112	7.8
10 1	4 27 9	0.0	46.342	13.092	5.3
10 1	4 29 28	0.0	46.351	13.113	1.0
10 1	9 8 57	0.0	46.242	13.070	5.8
10 1	11 20 13	0.0	46.256	13.262	7.6
10 1	11 46 35	0.0	46.282	13.279	8.2
10 1	14 46 12	2.8	46.273	13.257	11.3
10 1	15 28 58	0.0	46.300	13.093	8.6
10 1	15 33 15	0.0	46.340	13.042	6.9
10 1	17 23 53	0.0	46.304	13.168	6.5
10 1	18 8 29	0.0	46.356	13.113	10.2
10 1	18 34 57	3.3	46.365	13.042	10.1
10 1	18 34 57	0.0	46.272	13.102	10.1
10 1	22 5 5	0.0	46.292	13.038	6.6
10 2	0 51 22	0.0	46.272	13.102	9.2
10 2	10 15 25	0.0	46.295	13.189	8.9
10 2	10 53 12	0.0	46.335	13.228	3.4
10 2	11 43 49	0.0	46.346	13.227	2.2
10 2	17 49 49	0.0	46.383	13.127	1.0
10 2	17 49 43	0.0	46.326	13.128	1.0
10 2	18 59 25	0.0	46.358	13.099	1.0
10 3	0 8 14	3.3	46.325	13.255	5.2
10 3	0 8 14	0.0	46.339	13.085	11.6
10 3	17 23 51	0.0	46.400	13.175	1.8
10 3	17 49 52	3.4	46.391	13.010	1.0
10 3	19 33 2	0.0	46.362	13.042	1.0
10 3	19 39 42	0.0	46.352	13.068	4.2
10 3	21 20 33	0.0	46.348	13.306	7.9
10 3	22 59 41	0.0	46.321	13.245	7.6
10 4	11 12 1	2.8	46.360	13.071	1.0
10 4	26 17 51	0.0	45.312	13.164	12.0



M	D	T	ML	LAT.	LONG.	H	1	2	3	4	5	6	7	8	9
10	4	23 30 9	2.5	46.339	13.190	4.1	I	I	I	I	I	I	I	I	I
10	5	4 17 1	.0	46.323	13.199	6.7	I	I	I	I	I	I	I	I	I
10	5	7 33 12	3.3	46.353	13.179	9.1	I	I	I	I	I	I	I	I	I
10	6	3 23 51	3.0	46.302	13.105	7.2	I	I	I	I	I	I	I	I	I
10	6	10 21 41	.0	46.334	13.113	1.0	I	I	I	I	I	I	I	I	I
10	9	18 29 14	.0	46.339	13.023	1.0	I	I	I	I	I	I	I	I	I
10	7	2 45 11	.0	46.315	13.145	9.9	I	I	I	I	I	I	I	I	I
10	7	10 24 39	.0	46.327	13.016	1.0	I	I	I	I	I	I	I	I	I
10	7	20 4 13	.0	46.342	13.104	2.0	I	I	I	I	I	I	I	I	I
10	8	3 23 44	.0	46.346	13.112	1.0	I	I	I	I	I	I	I	I	I
10	8	5 9 25	.0	46.321	13.112	9.6	I	I	I	I	I	I	I	I	I
10	8	5 10 38	.0	46.322	13.032	8.0	I	I	I	I	I	I	I	I	I
10	8	8 15 54	2.5	46.280	13.117	5.6	I	I	I	I	I	I	I	I	I
10	8	9 58 39	.0	46.447	13.039	11.8	I	I	I	I	I	I	I	I	I
10	8	13 00 9	.0	46.344	13.167	1.0	I	I	I	I	I	I	I	I	I
10	8	19 26 14	.0	46.356	13.076	9.6	I	I	I	I	I	I	I	I	I
10	9	0 24 31	.0	46.177	13.206	1.0	I	I	I	I	I	I	I	I	I
10	9	4 20 11	2.9	46.200	13.220	3.1	I	I	I	I	I	I	I	I	I
10	8	13 23 23	2.8	46.333	13.128	16.8	I	I	I	I	I	I	I	I	I
10	9	13 34 34	.0	46.352	13.038	5.1	I	I	I	I	I	I	I	I	I
10	9	21 21 29	.0	46.236	13.236	5.7	I	I	I	I	I	I	I	I	I
10	9	22 5 6	.0	46.299	13.197	7.8	I	I	I	I	I	I	I	I	I

DEI.

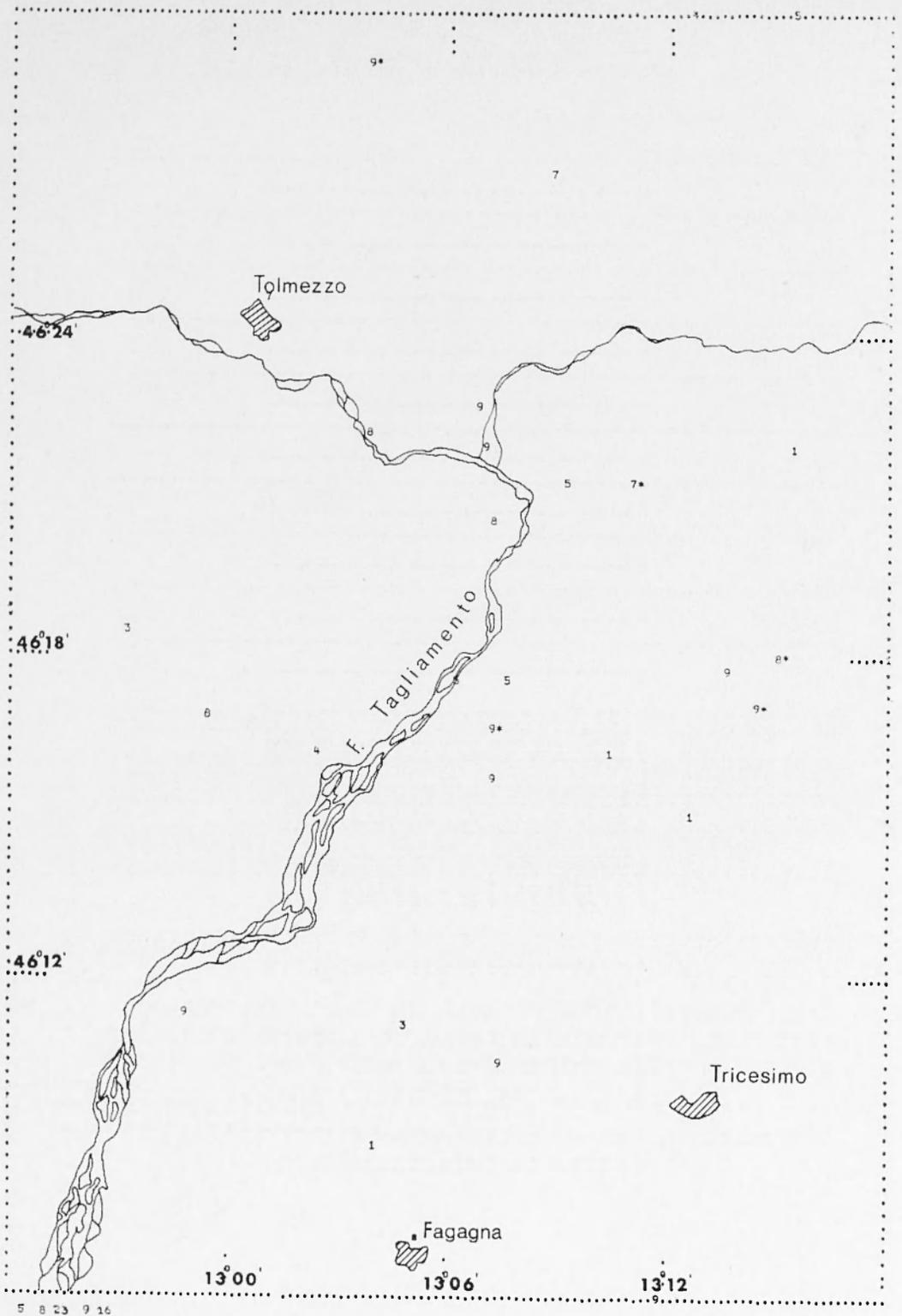


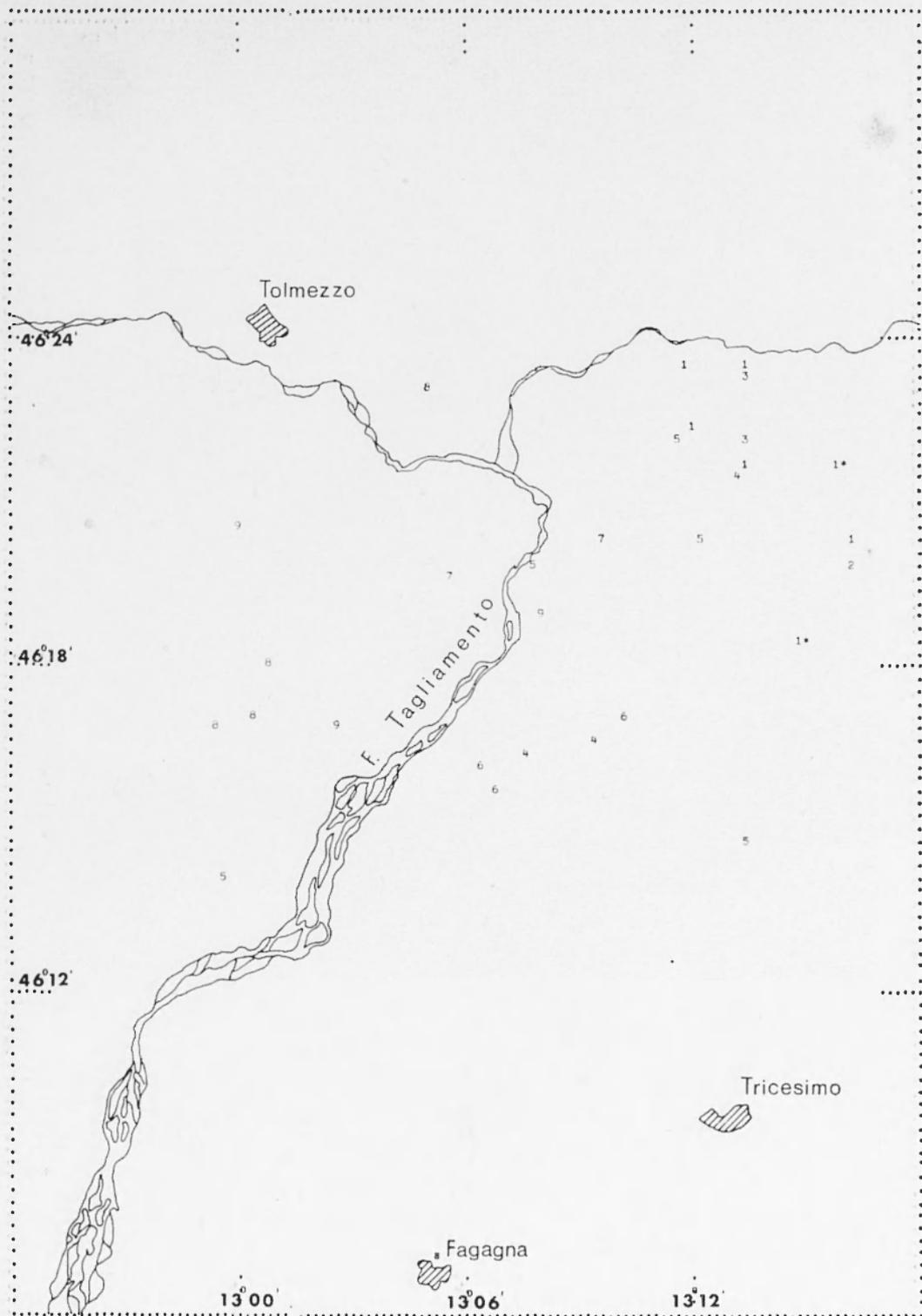
Fig. 1a-1b — From 5.6 19<sup>h</sup> 59<sup>m</sup> to 5.8 23<sup>h</sup> 09<sup>m</sup>.

Seismic activity distributed on a remarkably wide area, showing a high paroxysm of the focal volume, without preferential alignments. It should



5 6 23 9 16

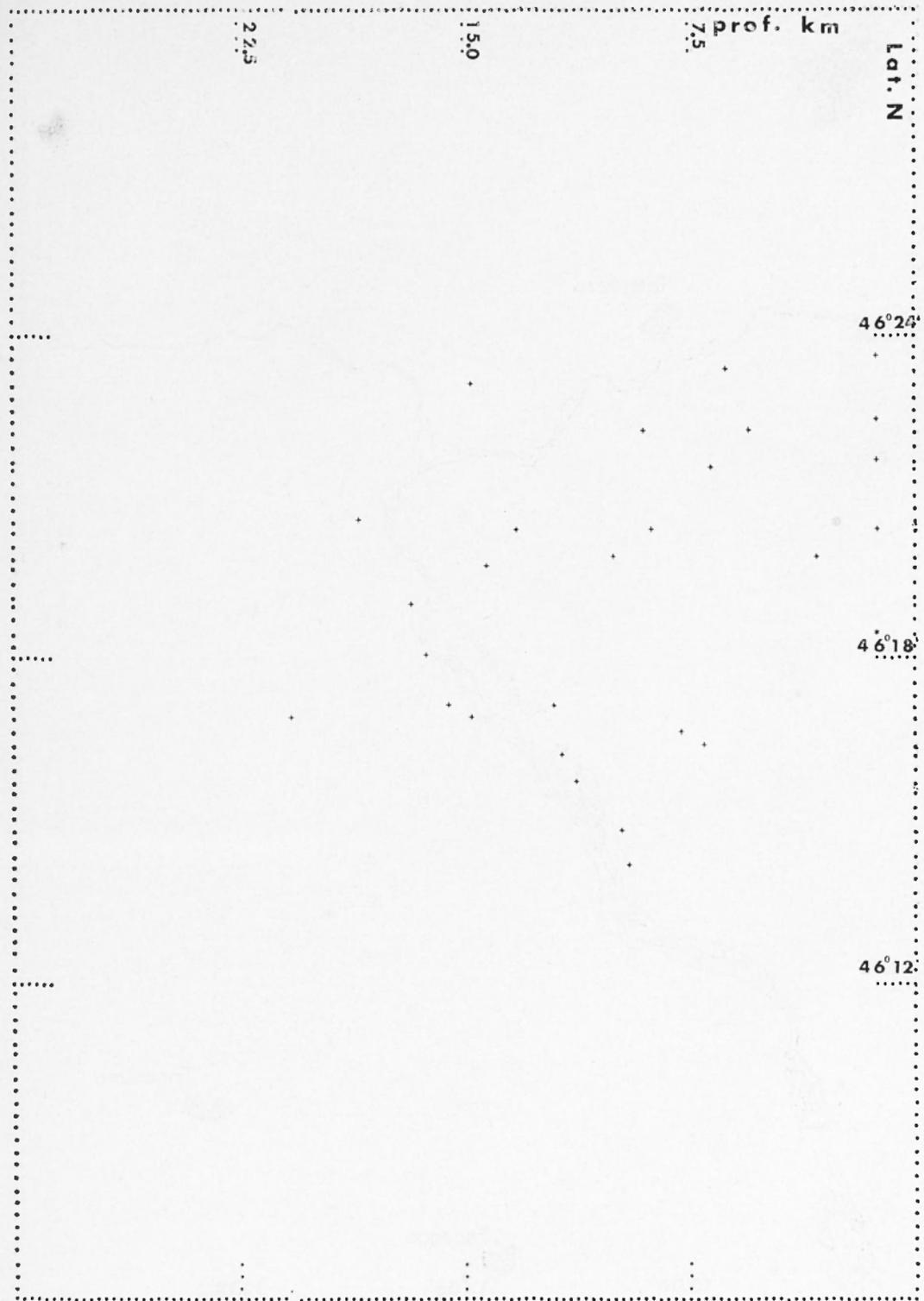
be noted that the number of the events reported is only a small part of the events actually occurred.



5 10 15 49 34

Figg. 2a-2b — From 5.8 25<sup>h</sup> 09<sup>m</sup> to 5.10 15<sup>h</sup> 49<sup>m</sup>.

The group of shocks with coordinates around the values 46.38 and 13.20 is a forewarning sign of the event with magnitude 4.7 on May 10 at 04<sup>h</sup> 35<sup>m</sup>. The zone was previously inactive. The prevailing alignment is WSW-ENE.



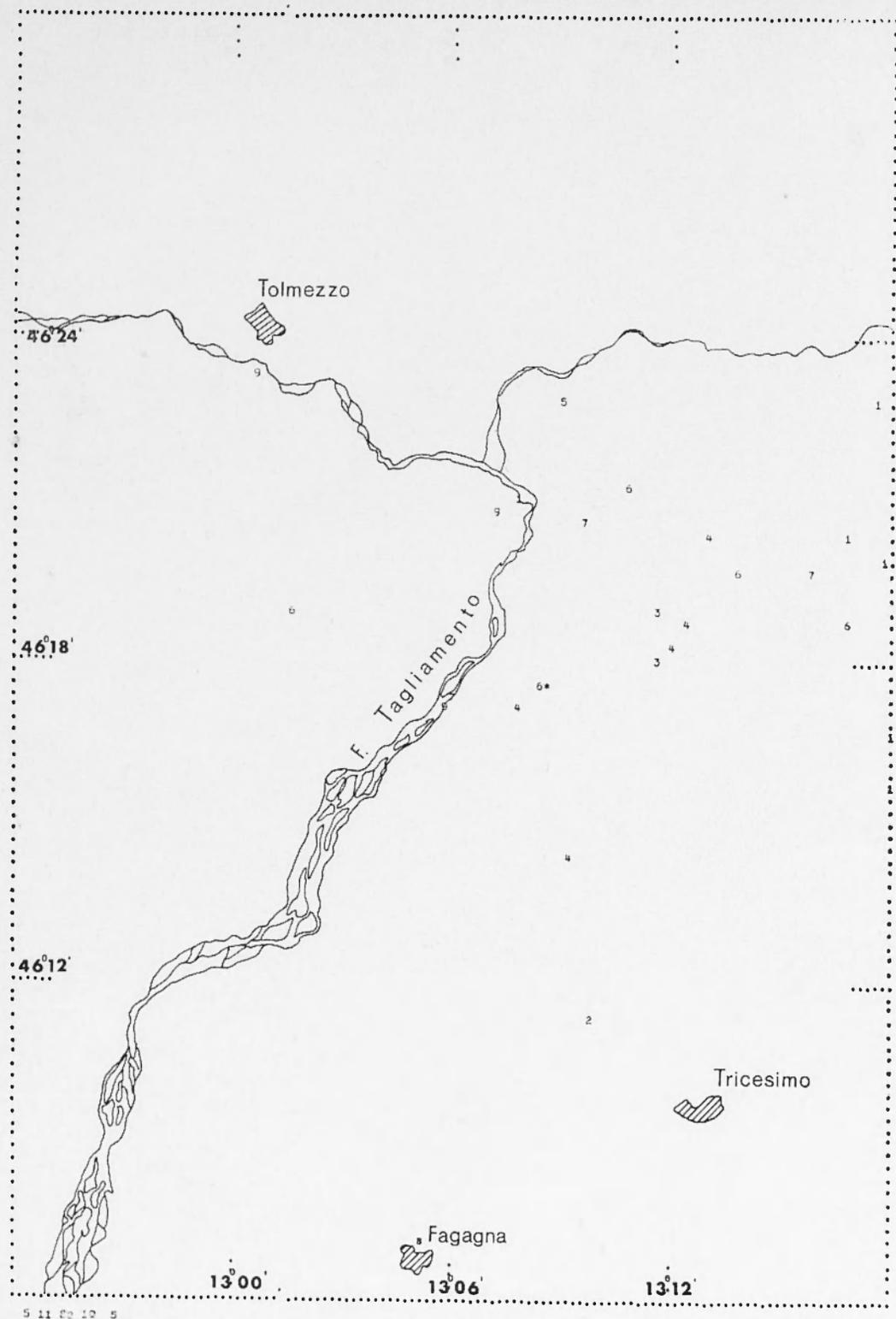
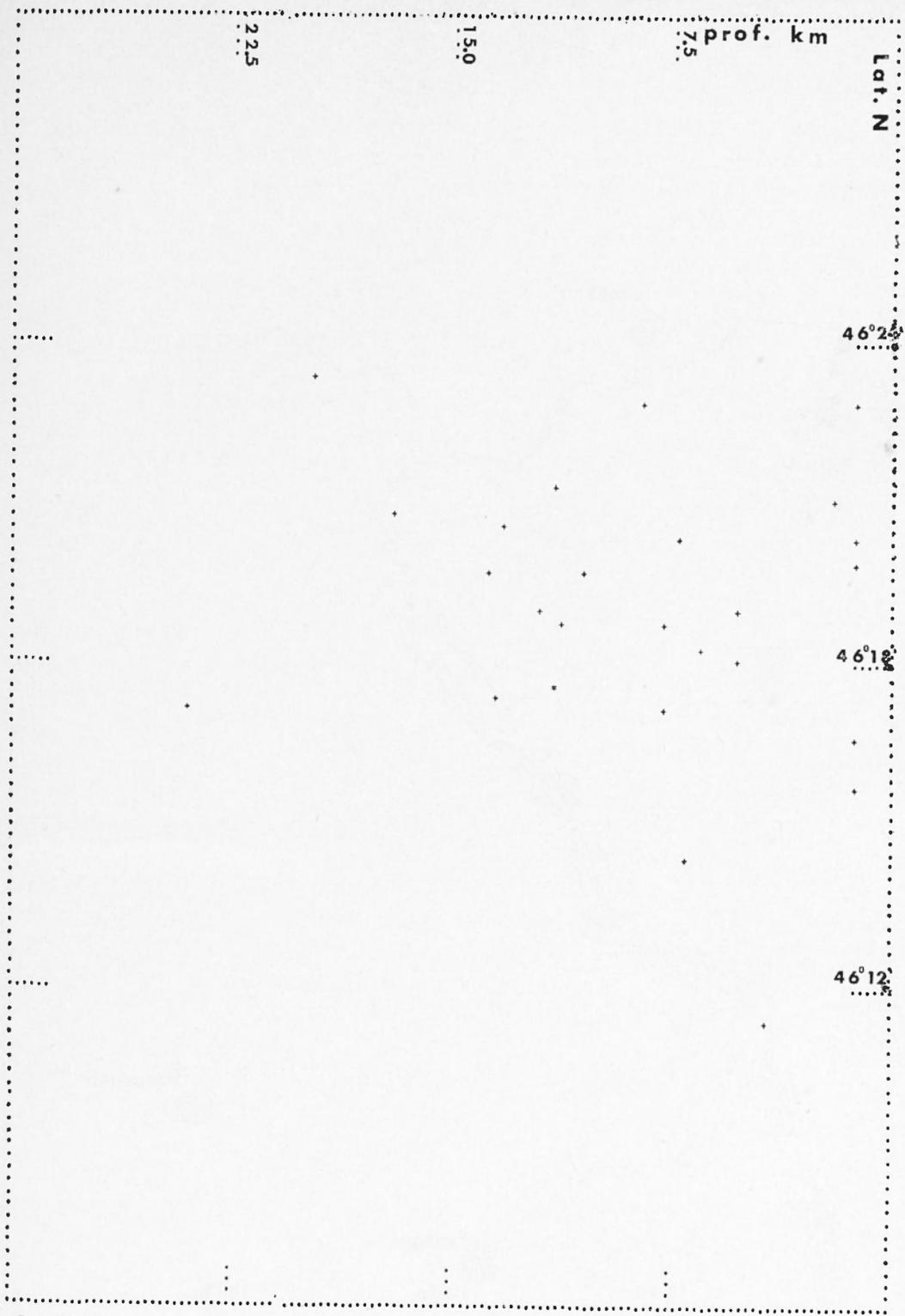


Fig. 3a-3b — From 5.10 15<sup>h</sup> 49<sup>m</sup> to 5.11 22<sup>h</sup> 18<sup>m</sup>.  
 The activity is lasting in the eastern part of the map.



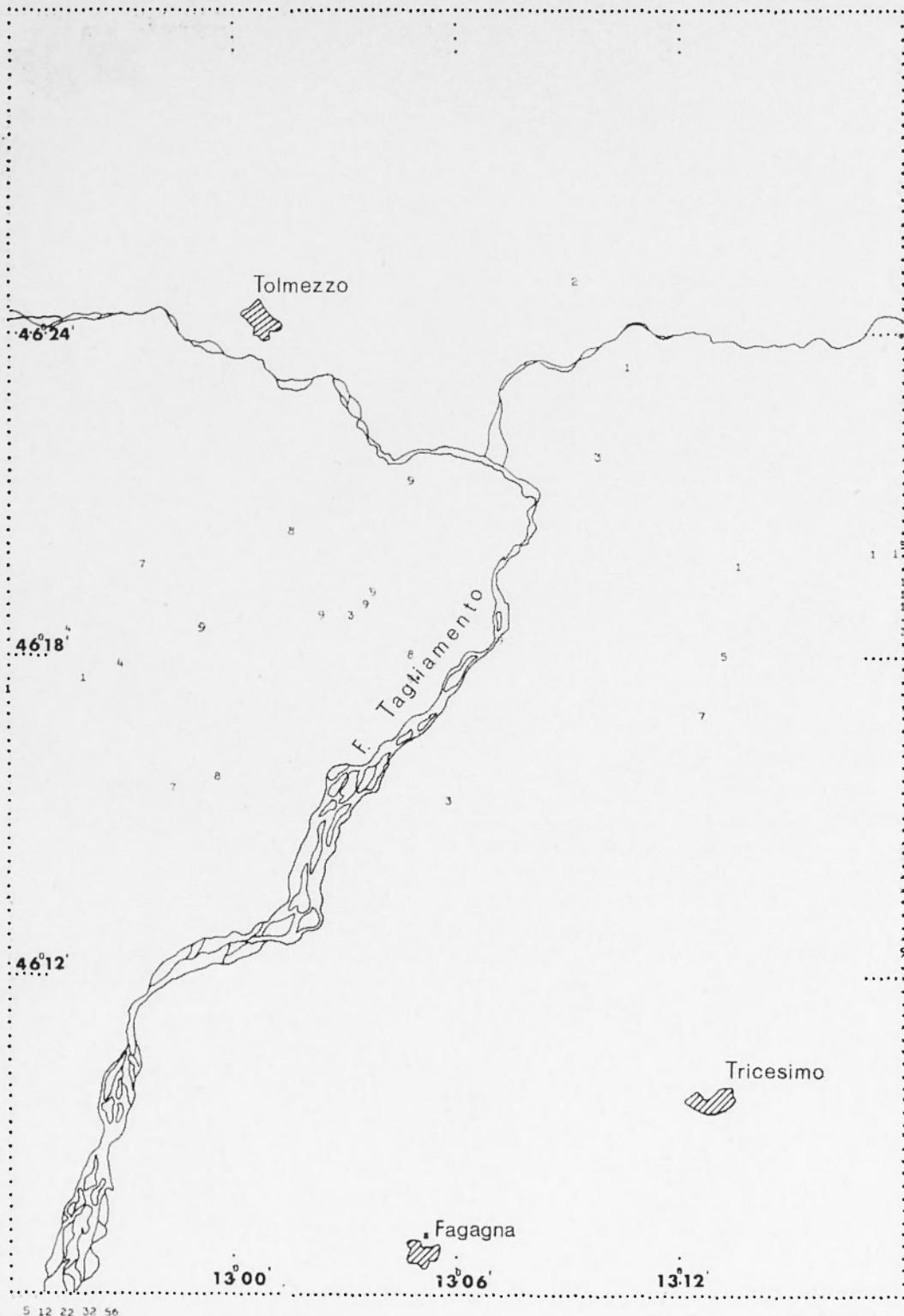
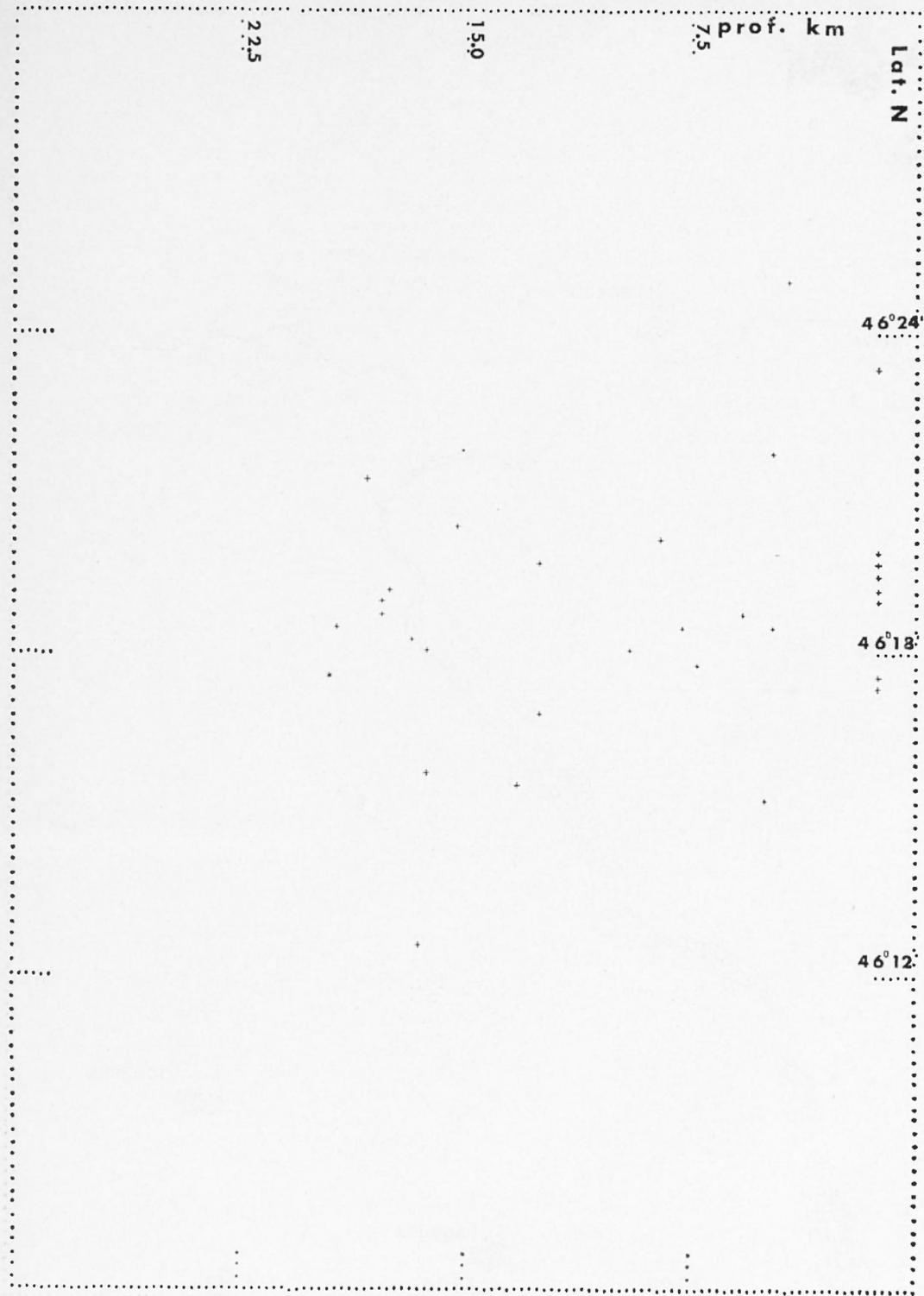
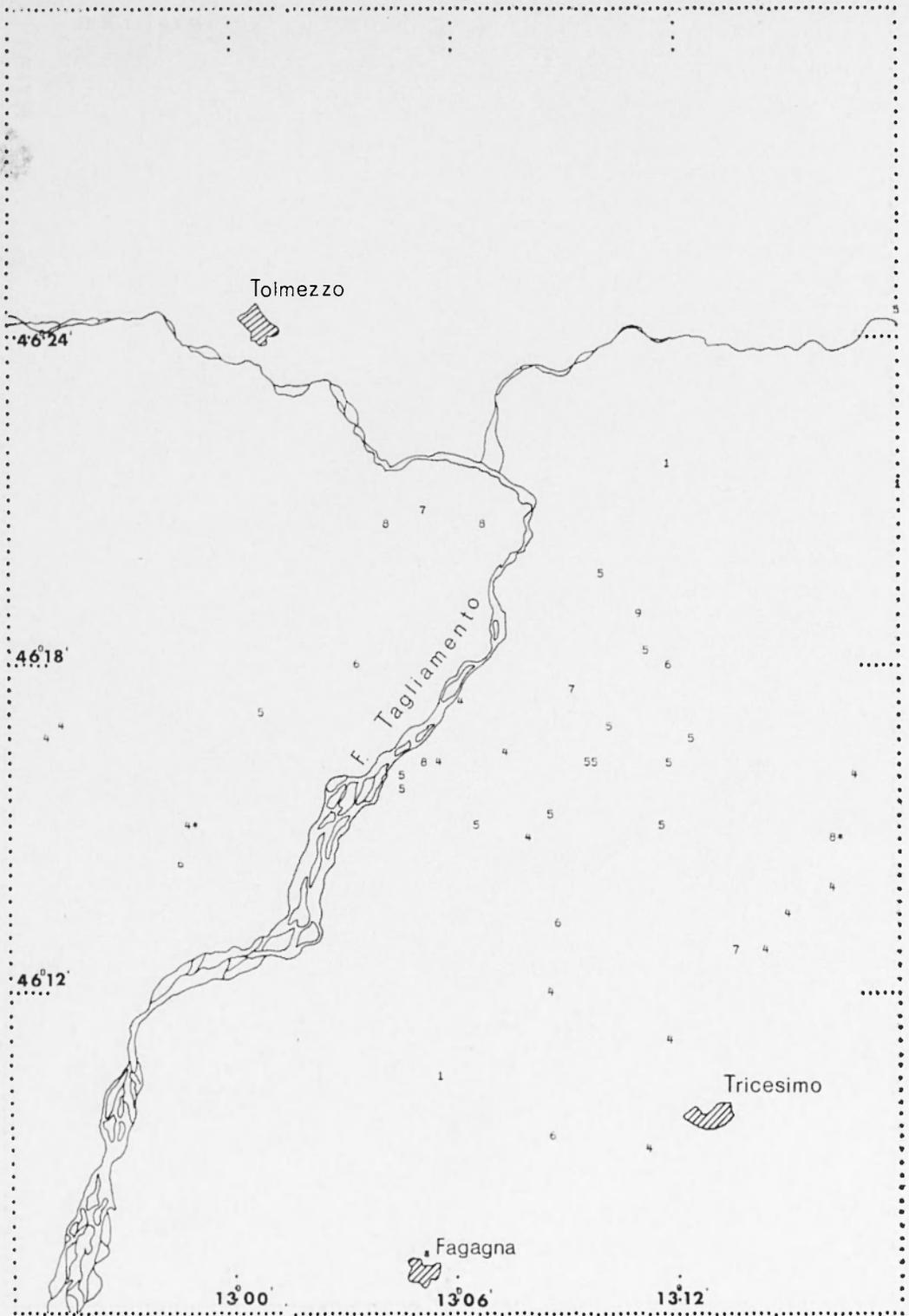


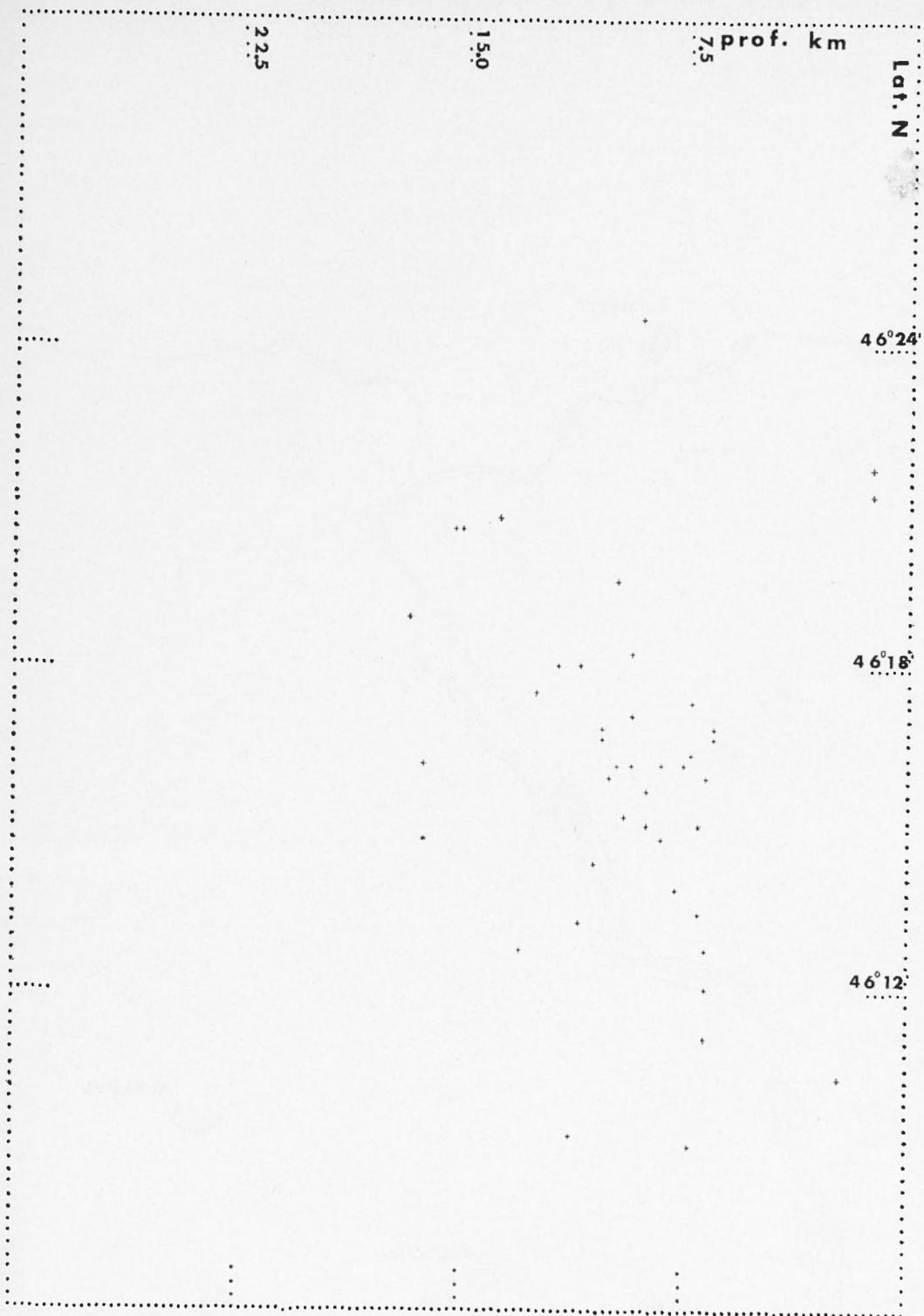
Fig. 4a-4b -- From 5.11 22<sup>h</sup> 18<sup>m</sup> to 5.12 22<sup>h</sup> 52<sup>m</sup>.  
 The shock of magnitude 5.3 on May 11 at 22<sup>h</sup> 44<sup>m</sup> is accompanied by a series of shocks which interested the same zone. In the western section of the map, an activity appears.

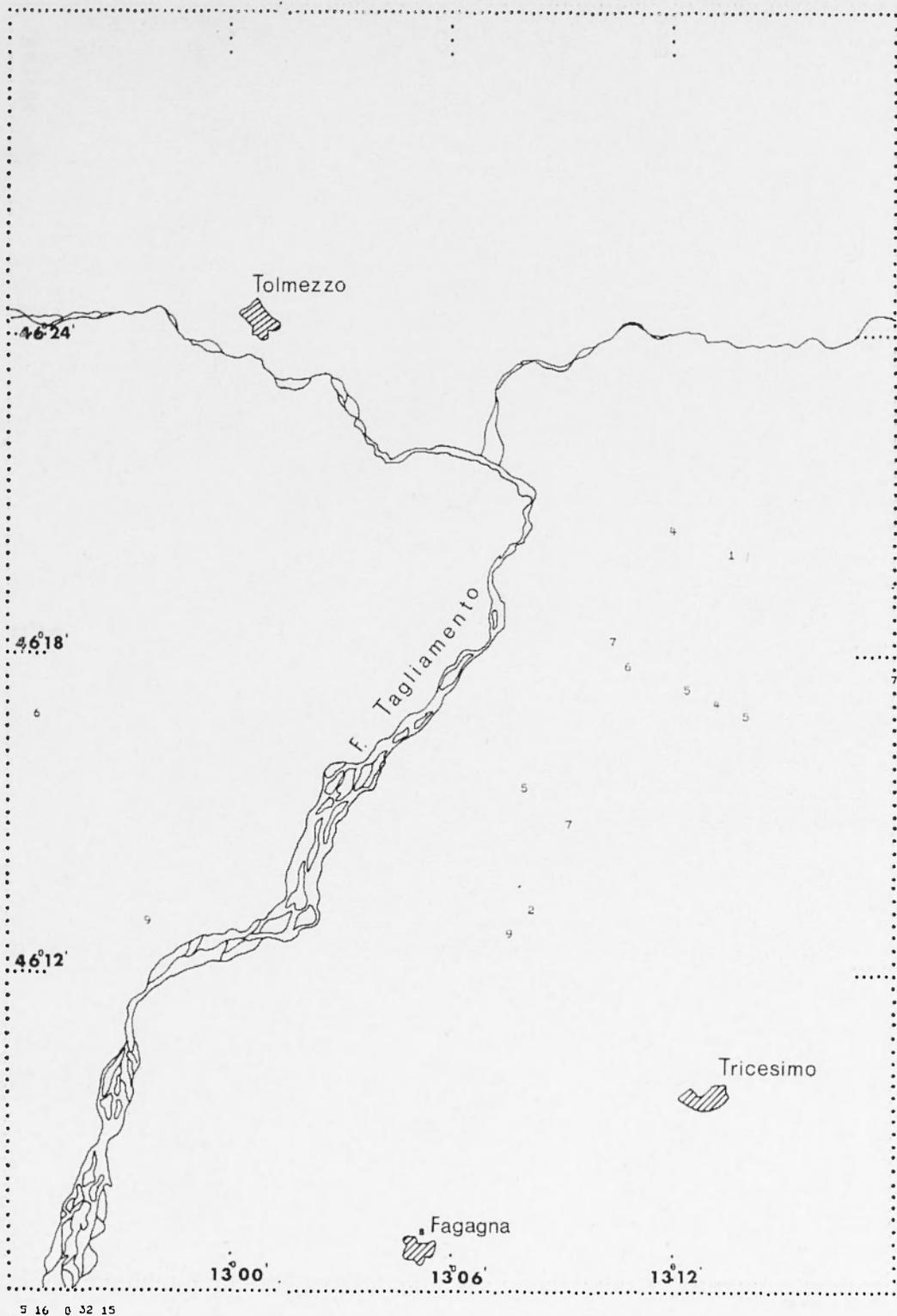




5 15 11 19 63

Fig. 5a-5b — From 5.12 22<sup>h</sup> 32<sup>m</sup> to 5.15 11<sup>h</sup> 19<sup>m</sup>.  
 The seismic activity, remaining in the central zone, shows a widening towards a south-eastern zone and its continuation in the western section with a shock of magnitude 4.7.

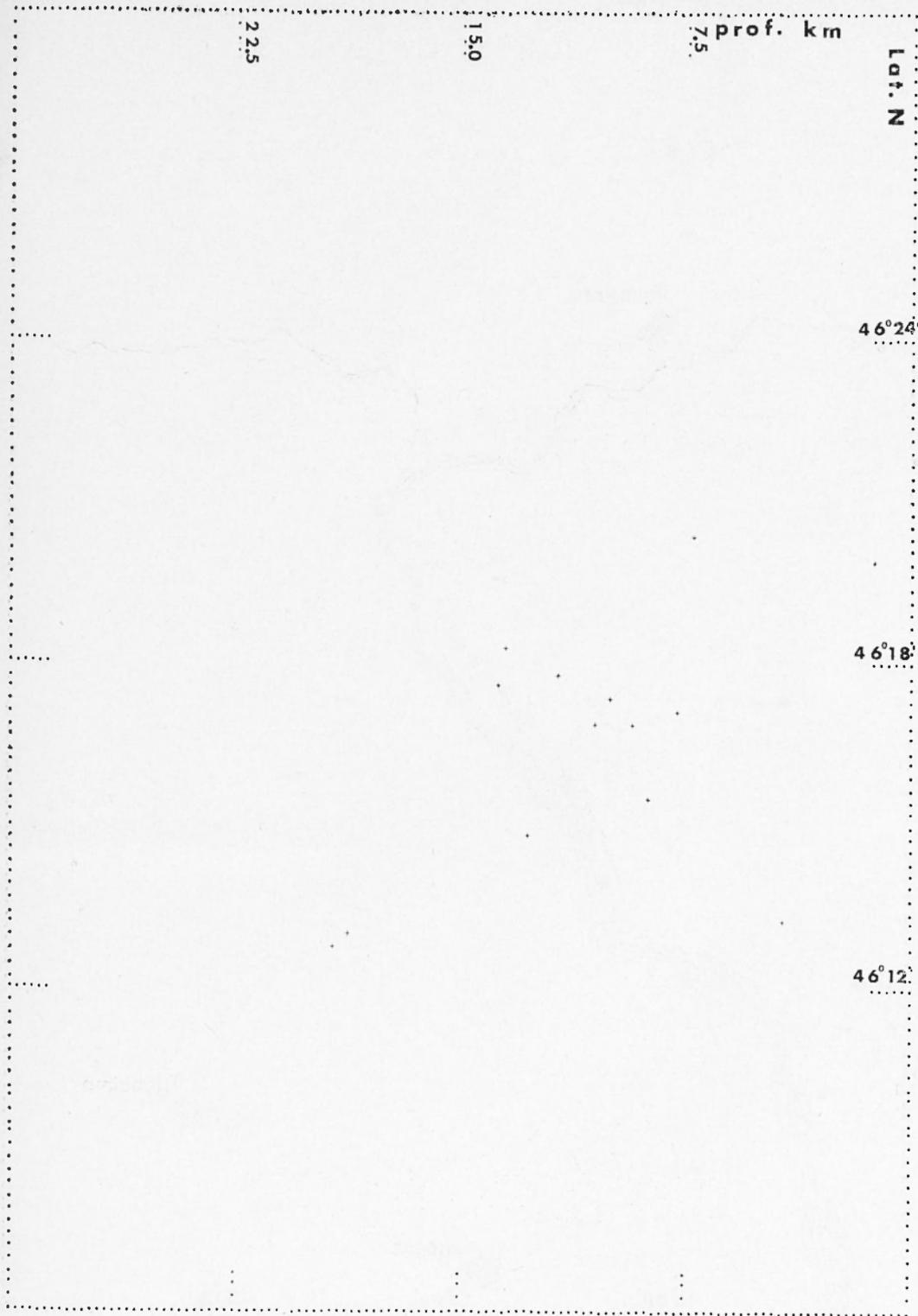




5 16 0 32 15

Fig. 6a-6b — From 5.15 11<sup>h</sup> 19<sup>m</sup> to 5.16 00<sup>h</sup> 52<sup>m</sup>.

The activity is concentrated in the eastern section with a SW-NE alignment of epicenters.



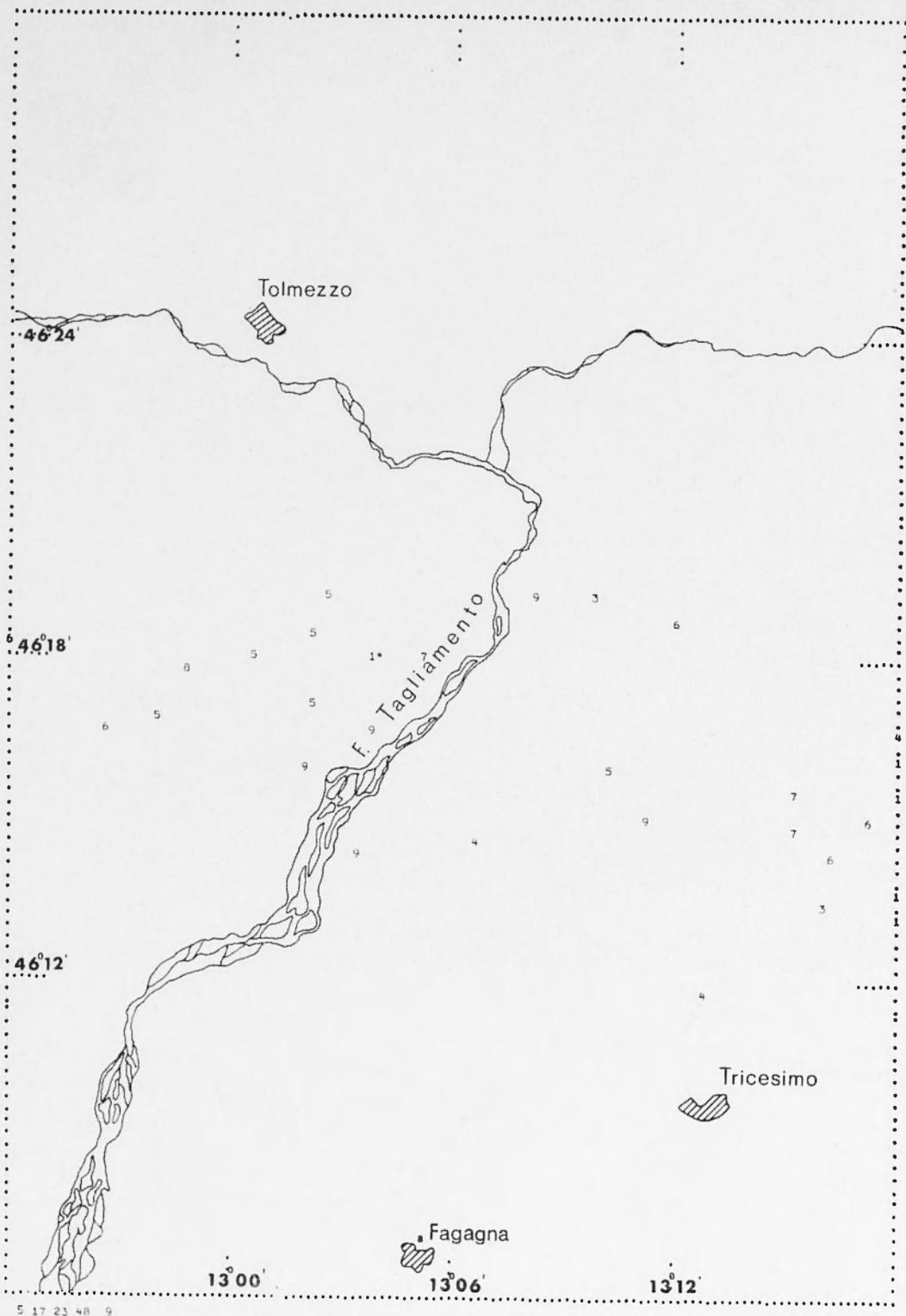


Fig. 7a-7b — From 5.16 00<sup>h</sup> 32<sup>m</sup> to 5.17 25<sup>m</sup> 48<sup>m</sup>.  
 The prevailing alignment is approximately E-W. The south-eastern section shows a certain activity.

Lat. N

prof. km

15.0

22.5

46°24'

46°18'

46°12'

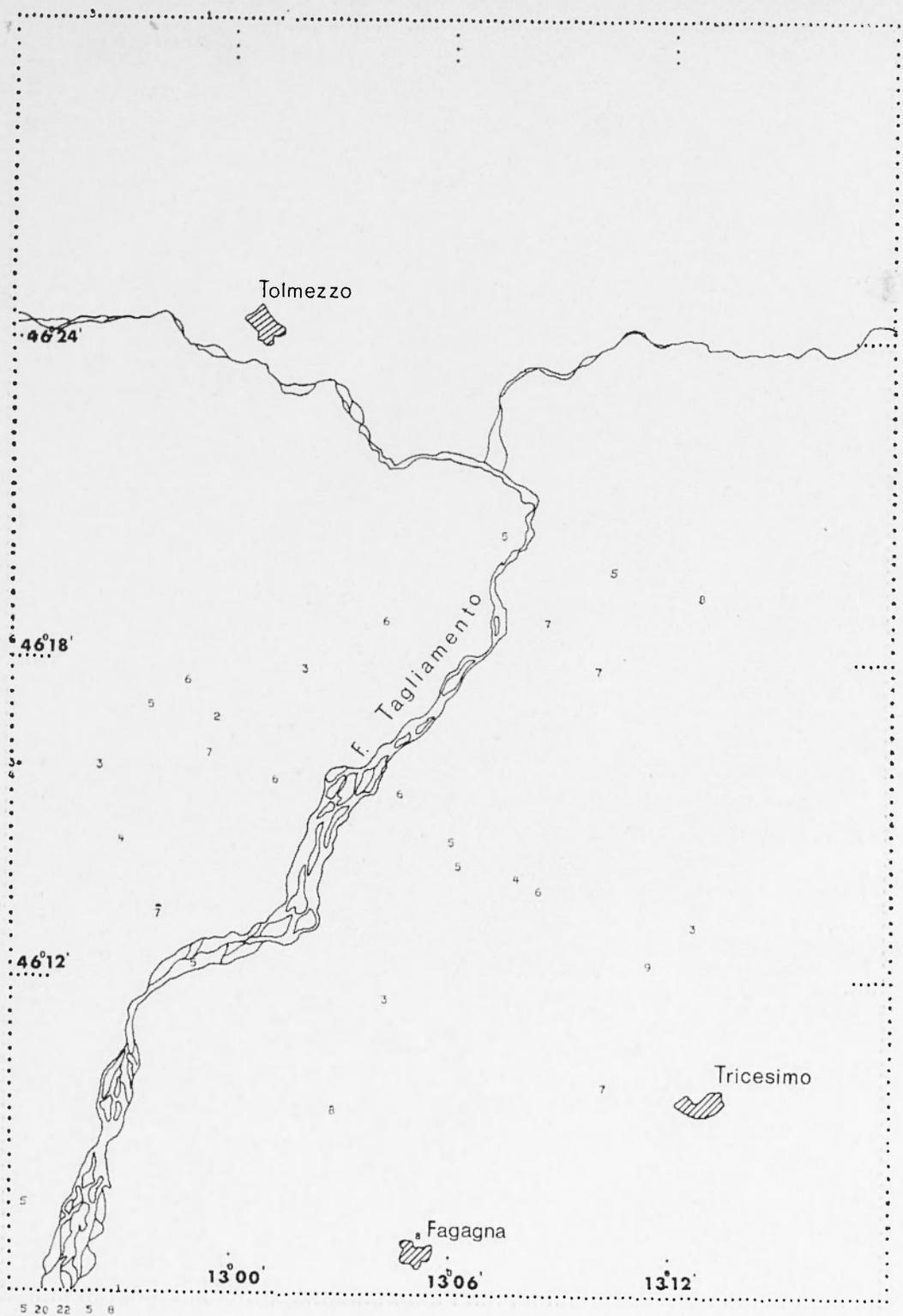
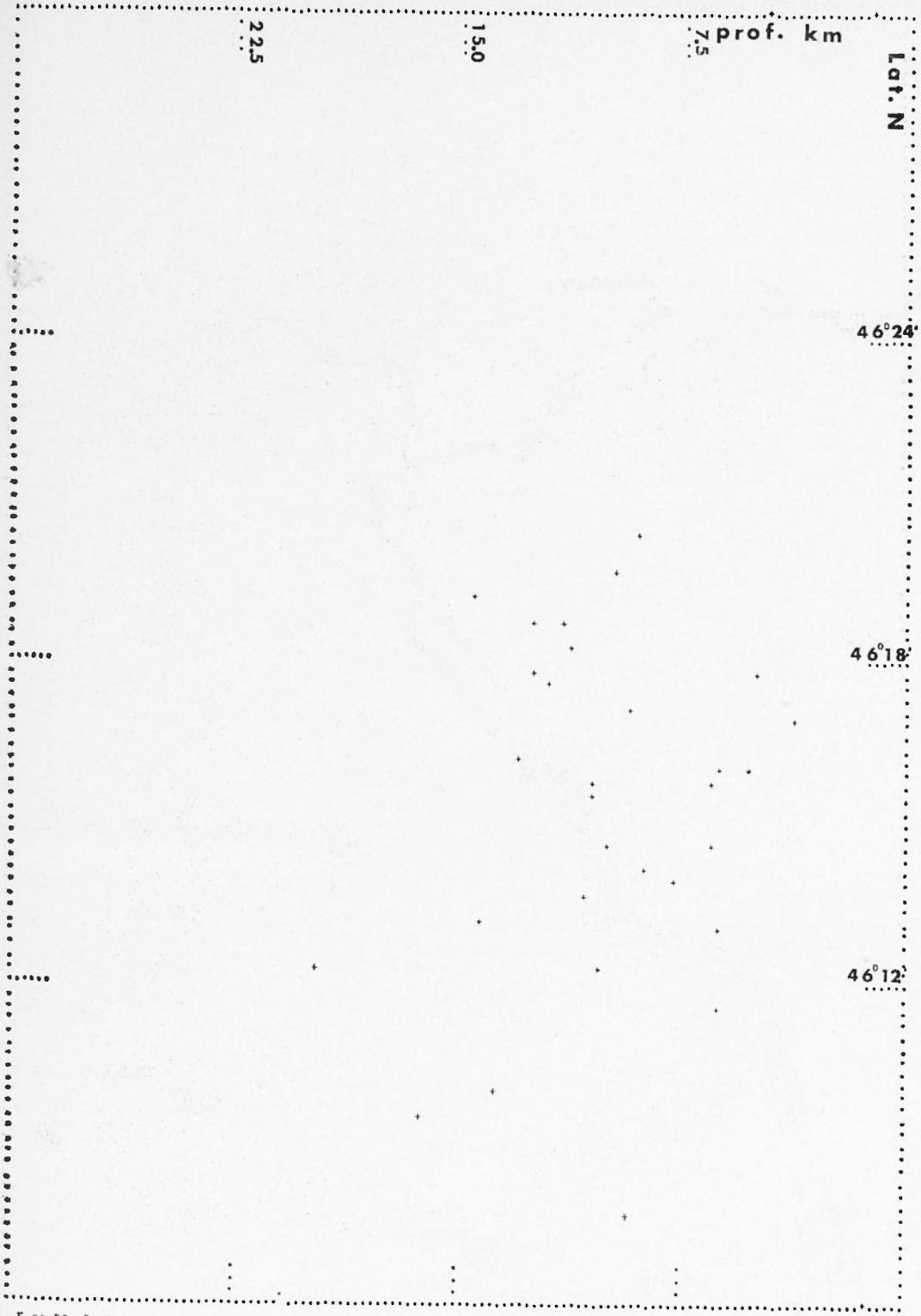


Fig. 8a-8b — From 5.17 23<sup>h</sup> 48<sup>m</sup> to 5.20 22<sup>h</sup> 05<sup>m</sup>.  
 The activity is widely diffused and reaches also the southern part of the map.



5 20 22 5 6

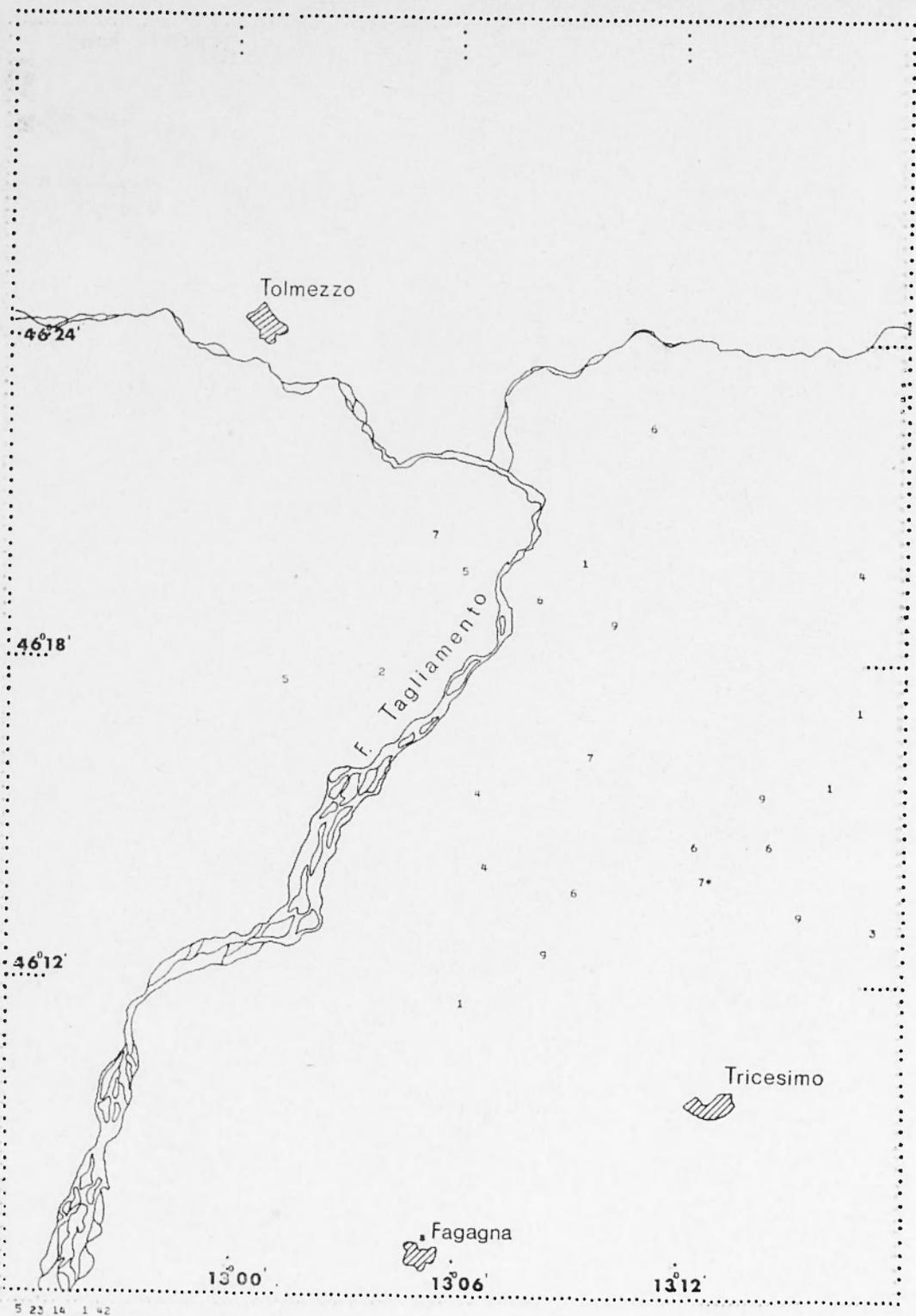
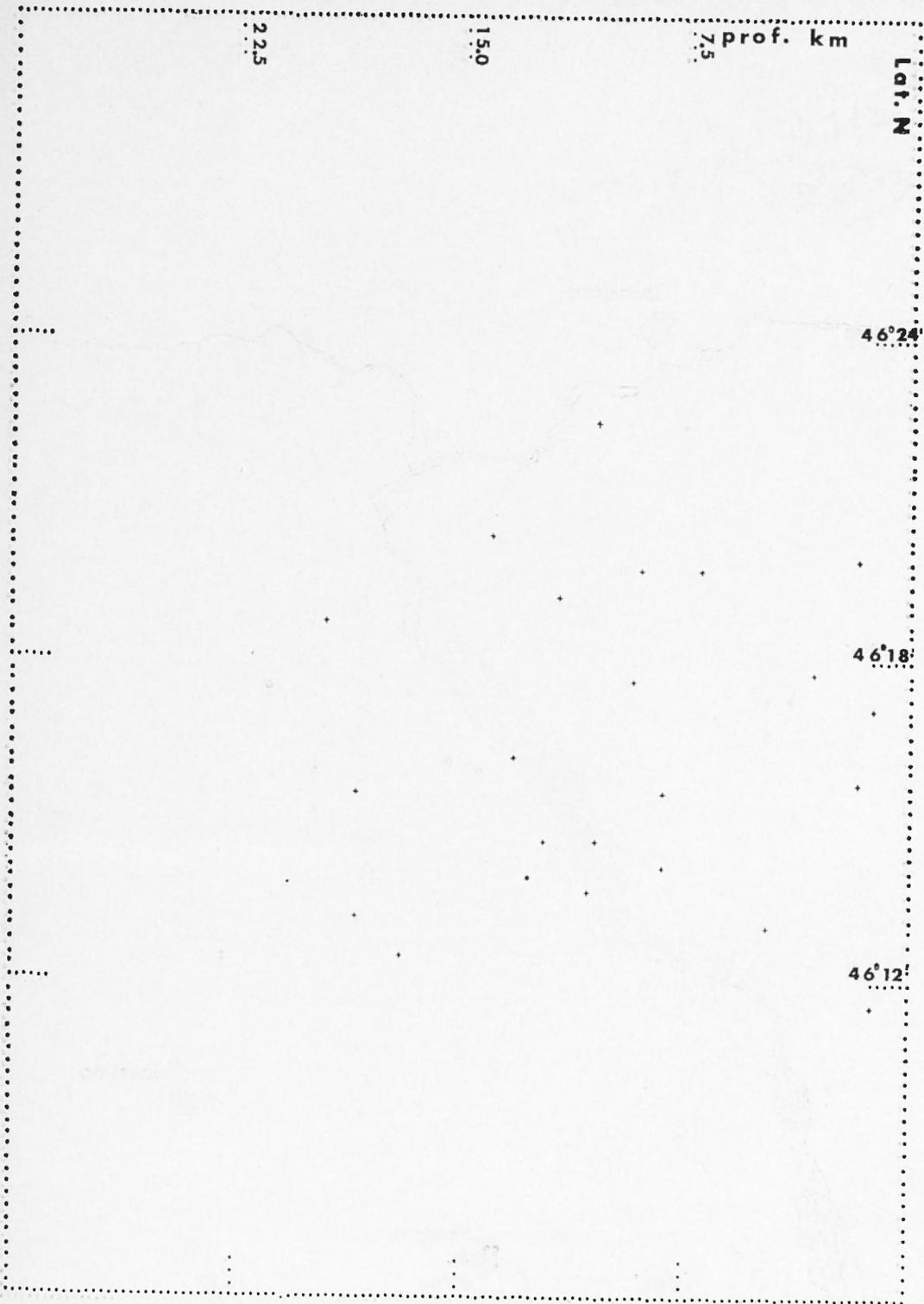


Fig. 9a-9b — From 5.20 22<sup>h</sup> 05<sup>m</sup> to 5.23 14<sup>h</sup> 01<sup>m</sup>.  
 In the southern section a shock of magnitude 4.0 appears; it is accompanied by minor events.



*[Faint, illegible text at the bottom of the page, likely bleed-through from the reverse side.]*

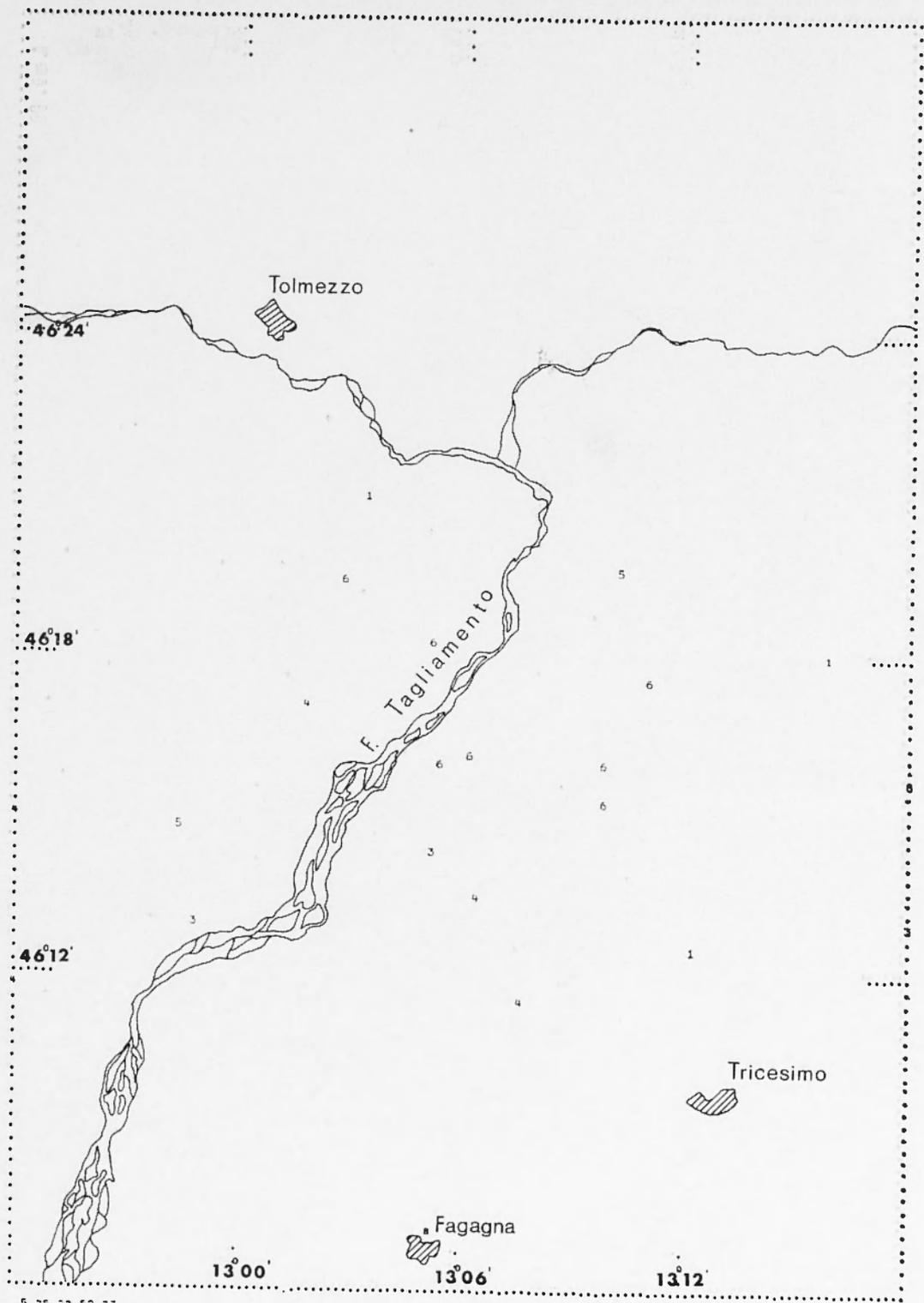
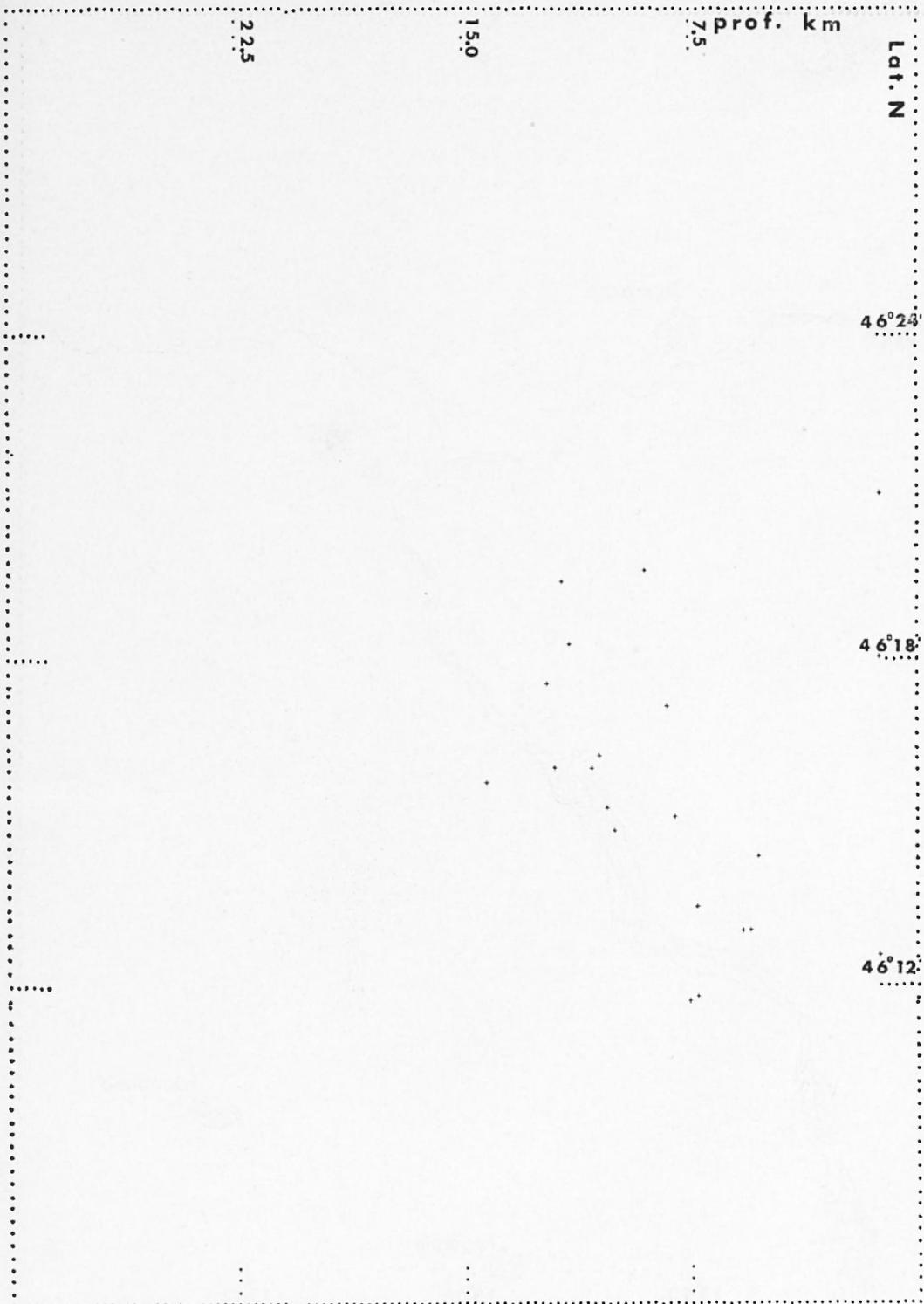


Fig. 10a-10b — From 5.23 14<sup>h</sup> 01<sup>m</sup> to 5.25 22<sup>h</sup> 52<sup>m</sup>.

The seismic activity, remarkably decreased and not accompanied by high magnitude shocks, covers in an almost uniform manner, all the map. In



5 25 22 52 37

the section an alignment of hypocenters still dipping towards N can be observed.

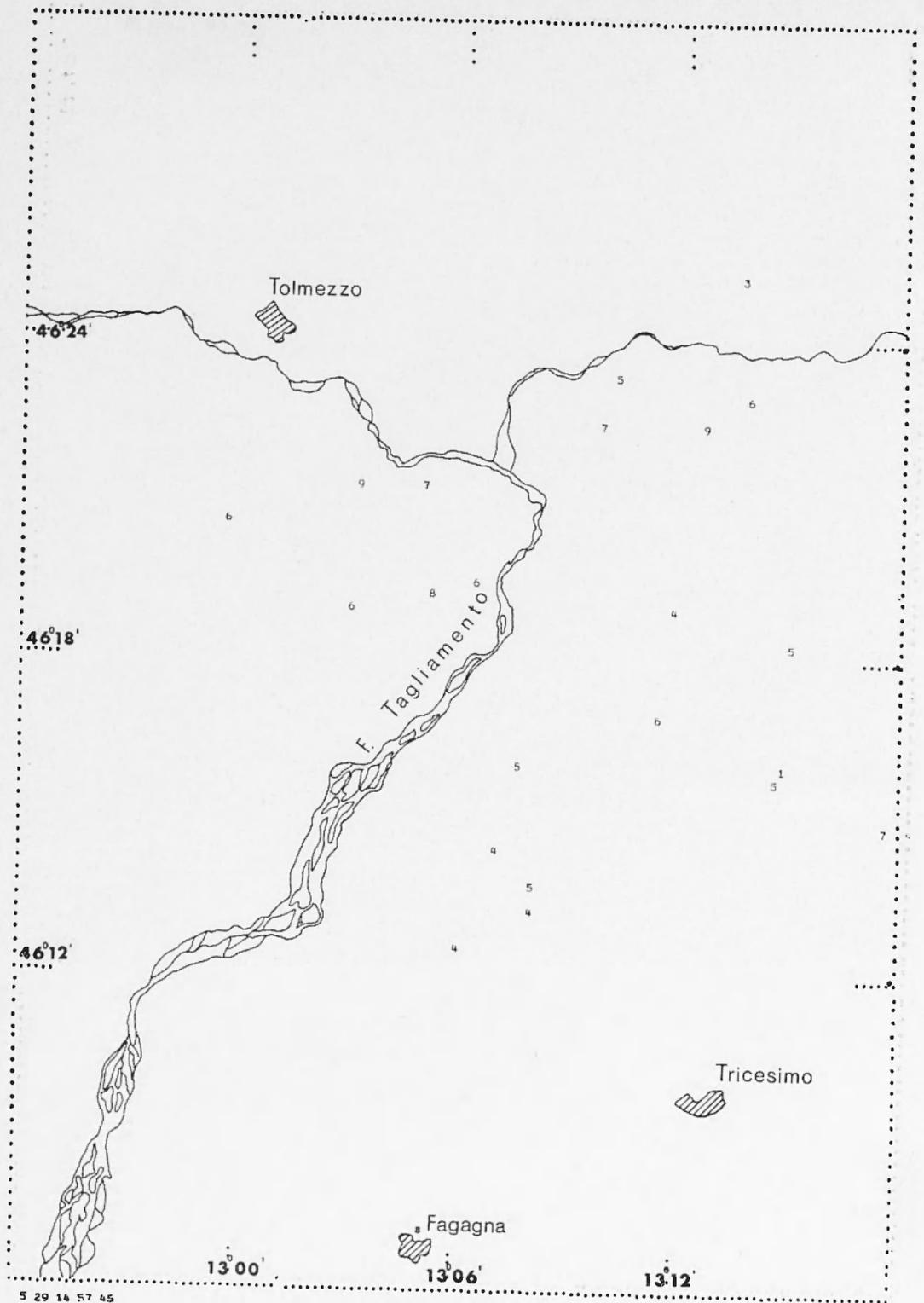
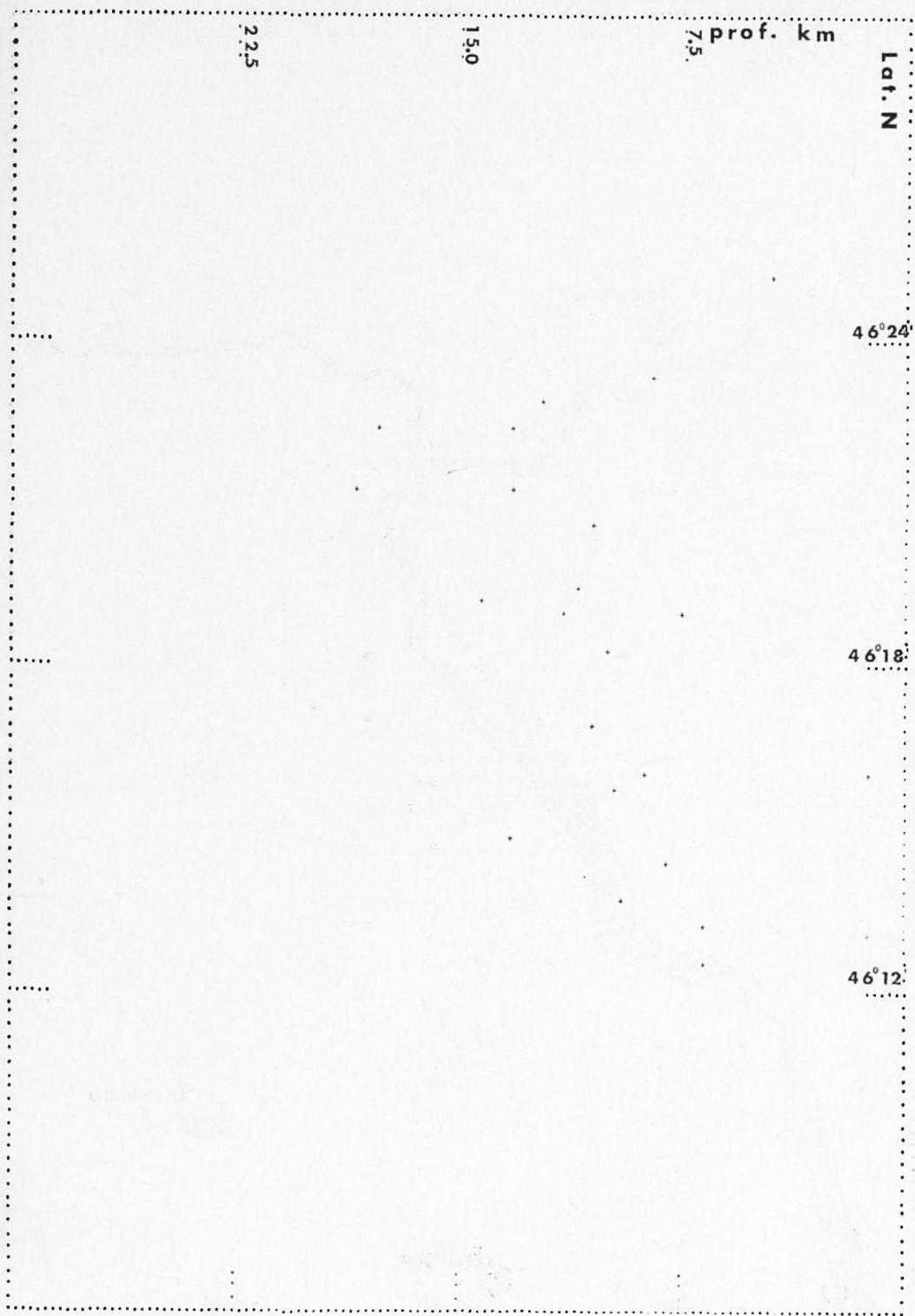


Fig. 11a-11b — From 5.25 22<sup>h</sup> 52<sup>m</sup> to 5.29 14<sup>h</sup> 57<sup>m</sup>.

The low magnitude activity, diffused in an almost uniform manner, is still visible; the alignment is that of the previous figure in the section N-S,



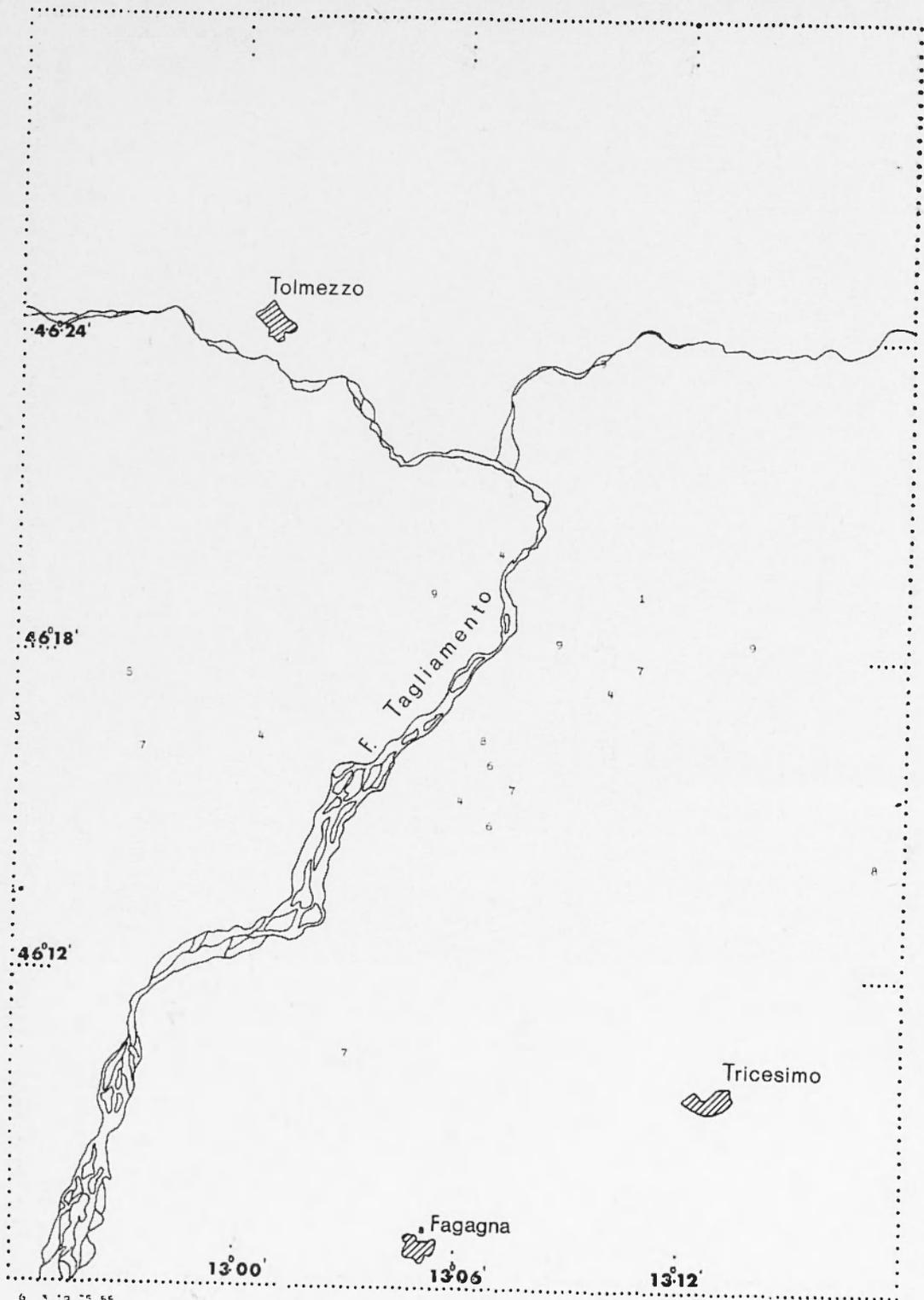
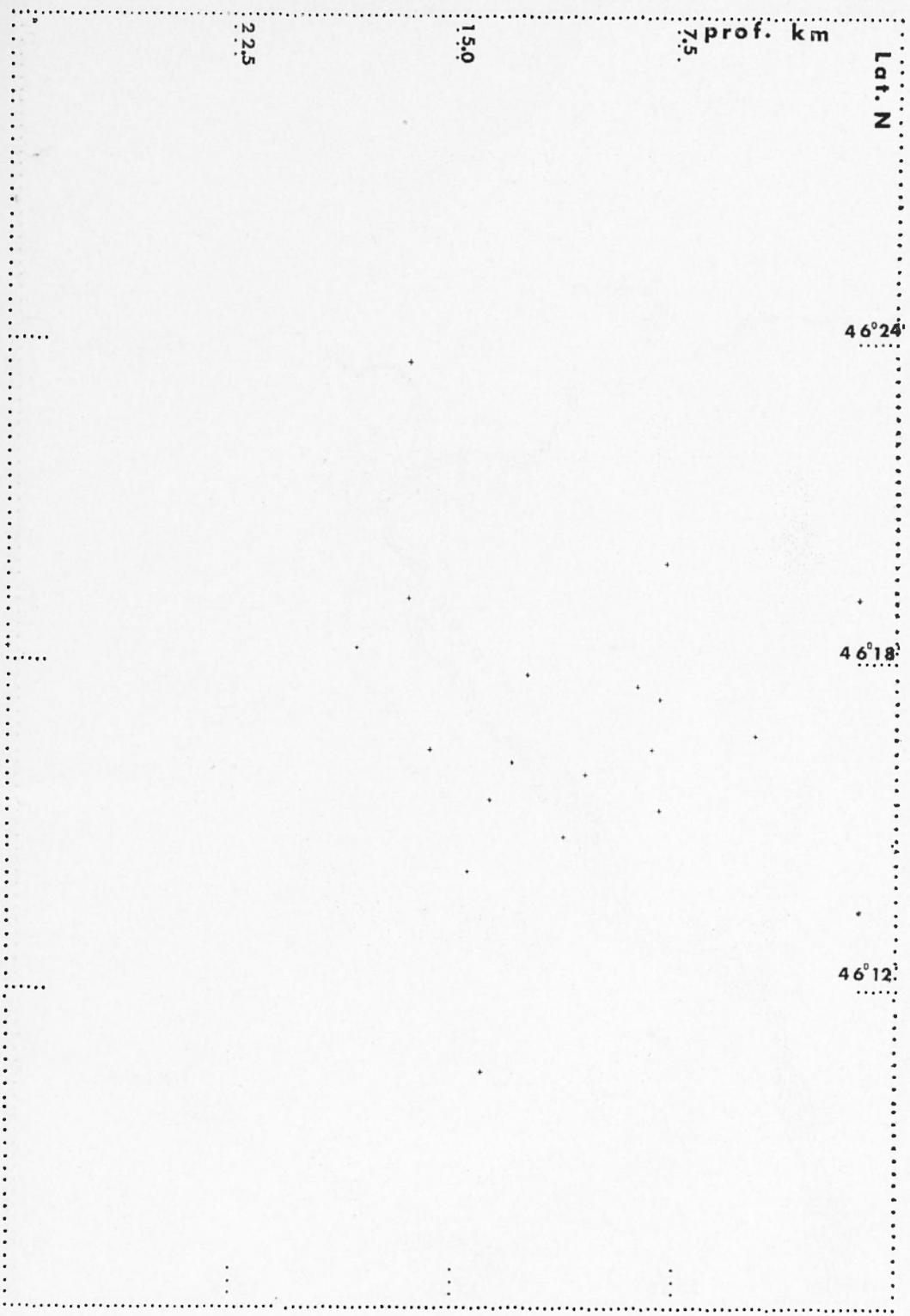


Fig. 12a-12b — From 5.25 14<sup>h</sup> 57<sup>m</sup> to 6.3 19<sup>h</sup> 35<sup>m</sup>.  
The seismic activity is remarkably decreased.



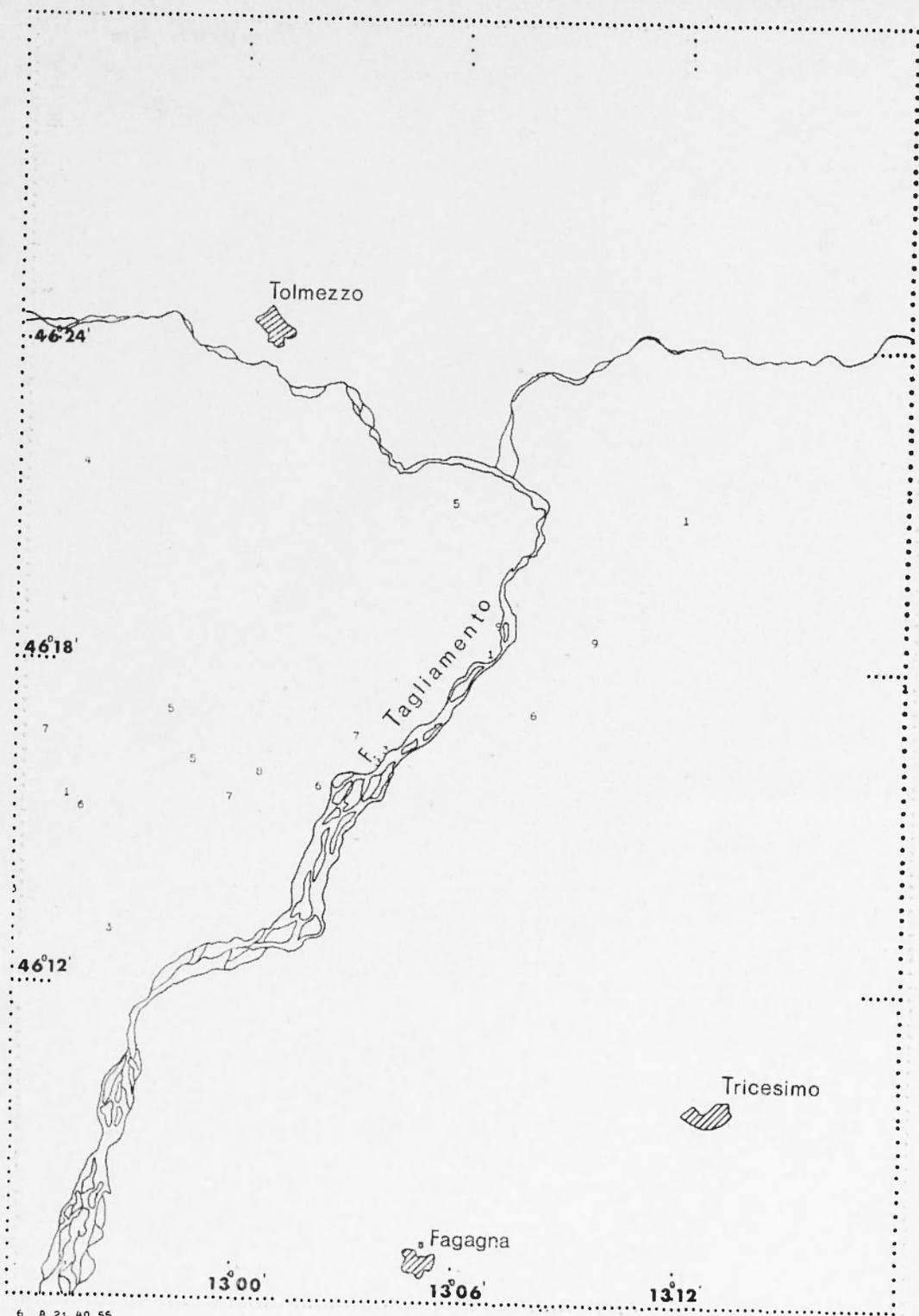
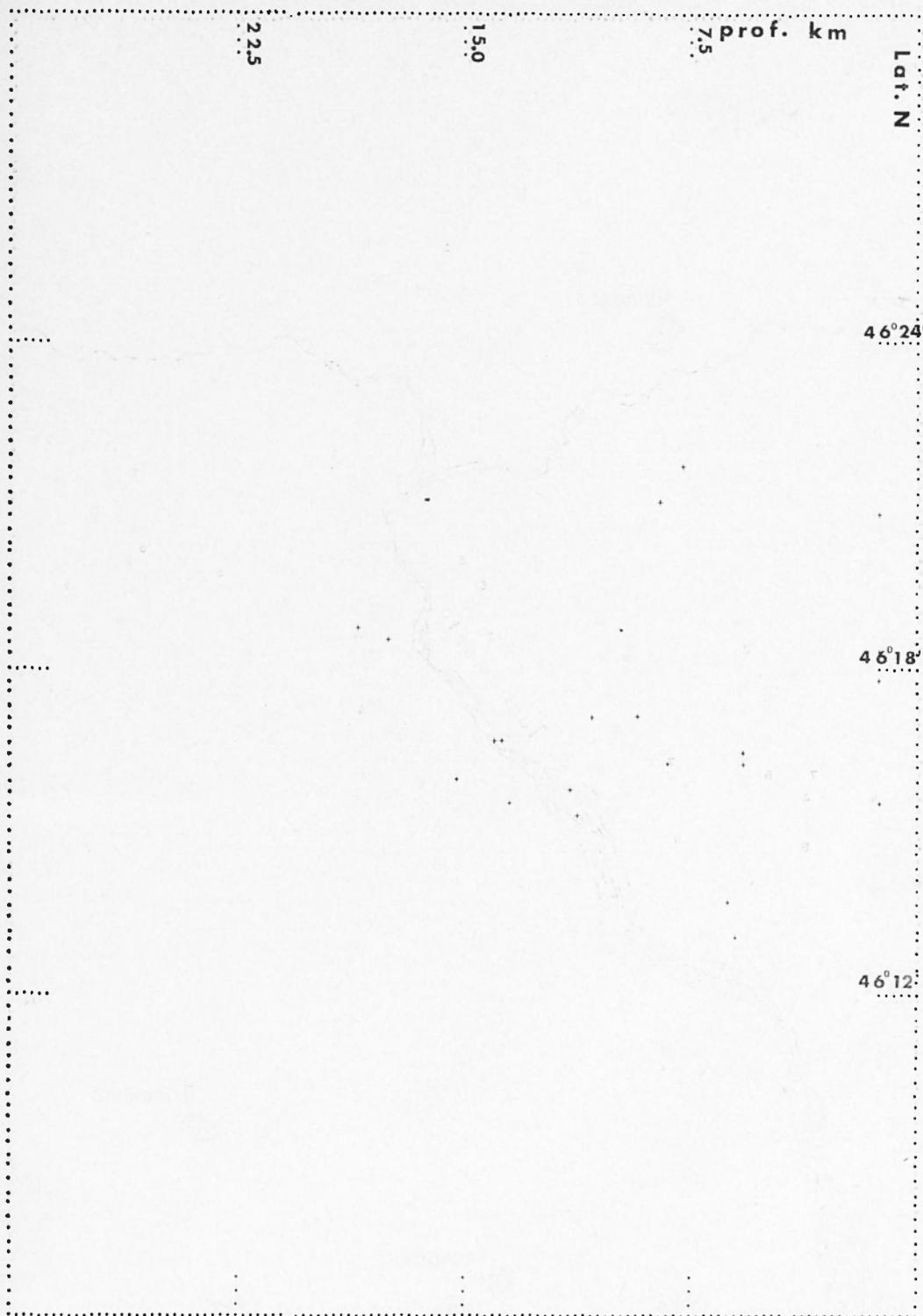


Fig. 13a-13b — From 6.3 19<sup>h</sup> 35<sup>m</sup> to 6.8 21<sup>h</sup> 40<sup>m</sup>.  
 The epicenters seem to be aligned in a WSW-ENE direction, with a higher concentration to W.



This is a plot of depth (prof. km) versus latitude (Lat. N). The x-axis represents depth in kilometers, with major ticks at 22.5, 15.0, and 7.5. The y-axis represents latitude in degrees North, with major ticks at 46°12', 46°18', and 46°24'. The plot shows a complex, multi-lobed curve that generally trends from higher latitudes and greater depths towards lower latitudes and shallower depths. Several data points are marked with small dots along the curve.

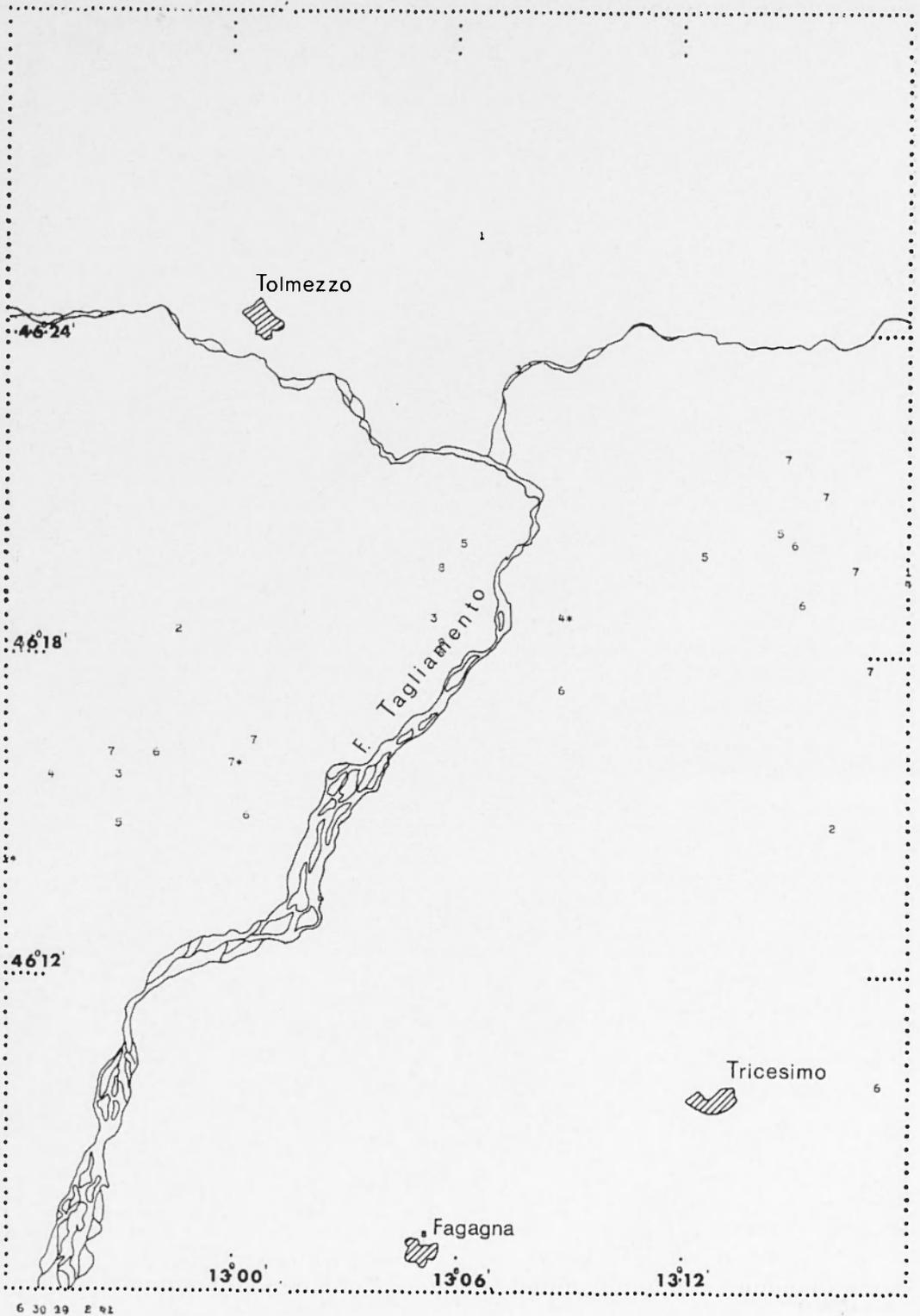
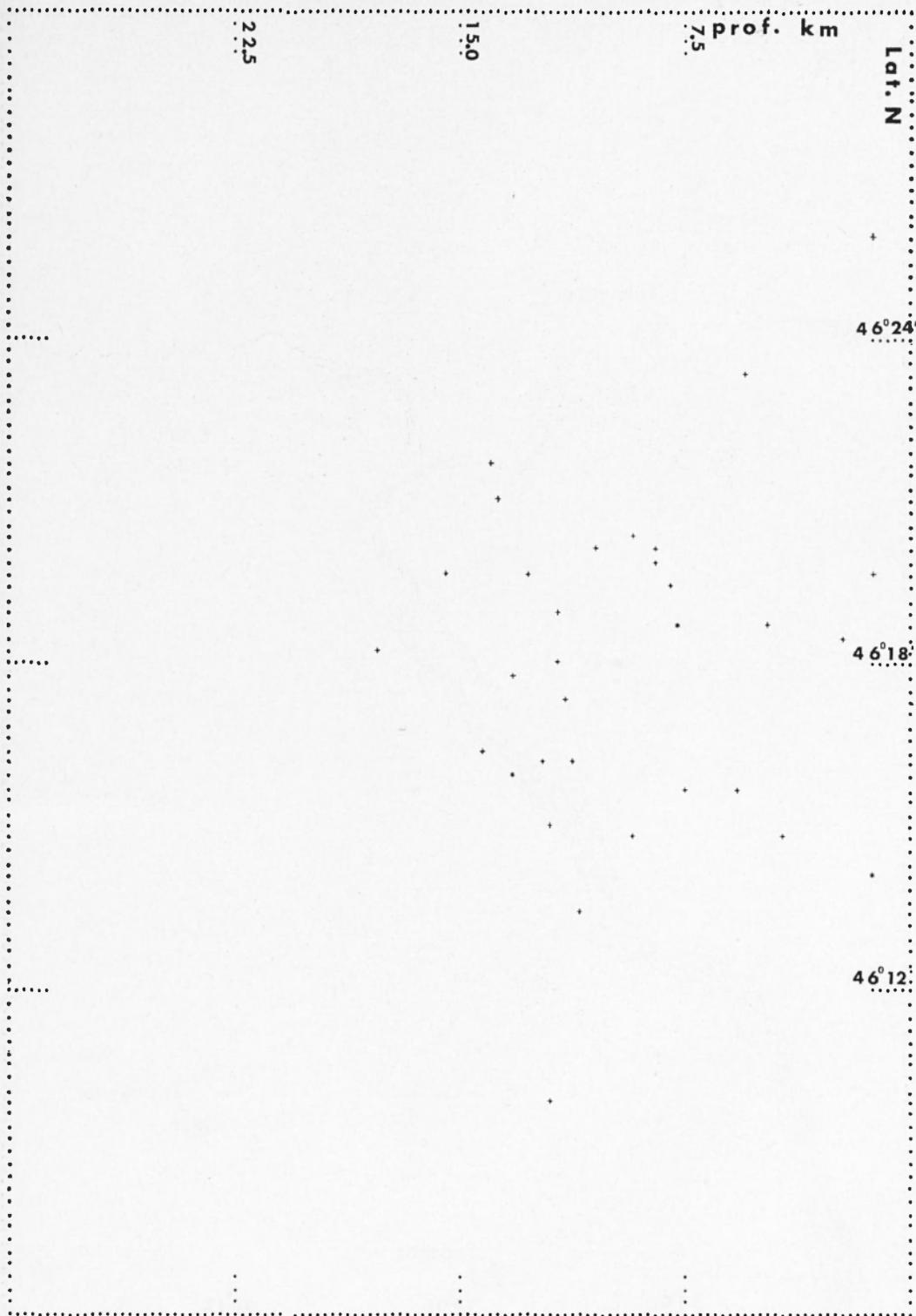


Fig. 14a-14b — From 6.8 21<sup>h</sup> 40<sup>m</sup> to 6.30 19<sup>h</sup> 02<sup>m</sup>.

The distribution of the epicenters is interesting, mainly, three aligned sectors. The seismic activity in the western zone can be correlated with



6 30 19 2 41

the shock of magnitude 4.2 on June 9, at 18<sup>h</sup> 48<sup>m</sup>. The activity of the central sector ends with the shock on June 26, at 11<sup>h</sup> 13<sup>m</sup>, having a magnitude 4.3. The eastern sector shows the appearance of low magnitude events.

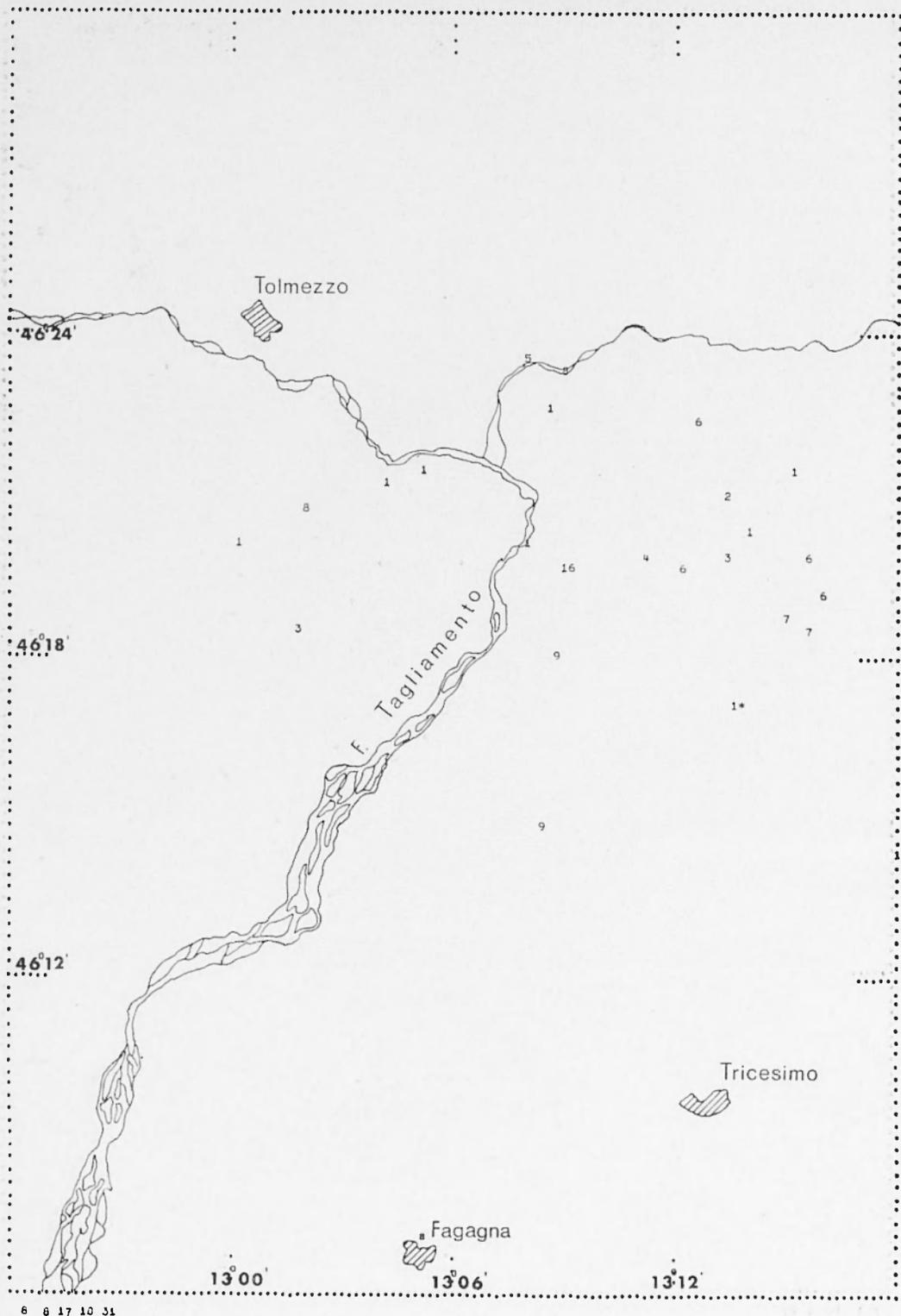
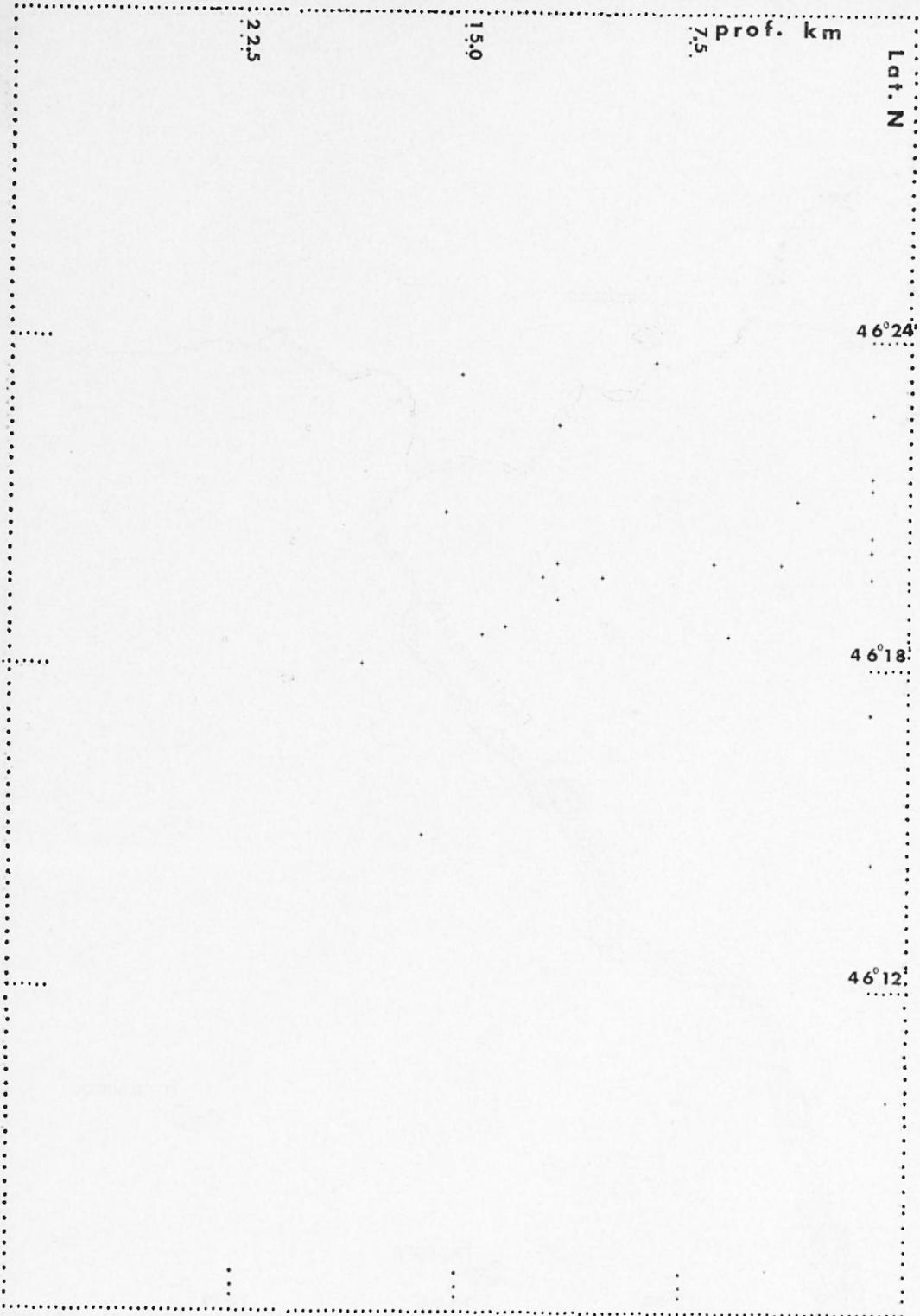


Fig. 15a-15b — From 6.30 19<sup>h</sup> 02<sup>m</sup> to 8.8 17<sup>h</sup> 10<sup>m</sup>.  
 In the eastern sector on July 10 at 04<sup>h</sup> 11<sup>m</sup> an event of magnitude 4.1  
 occurs. Absence of activity in the western sector.



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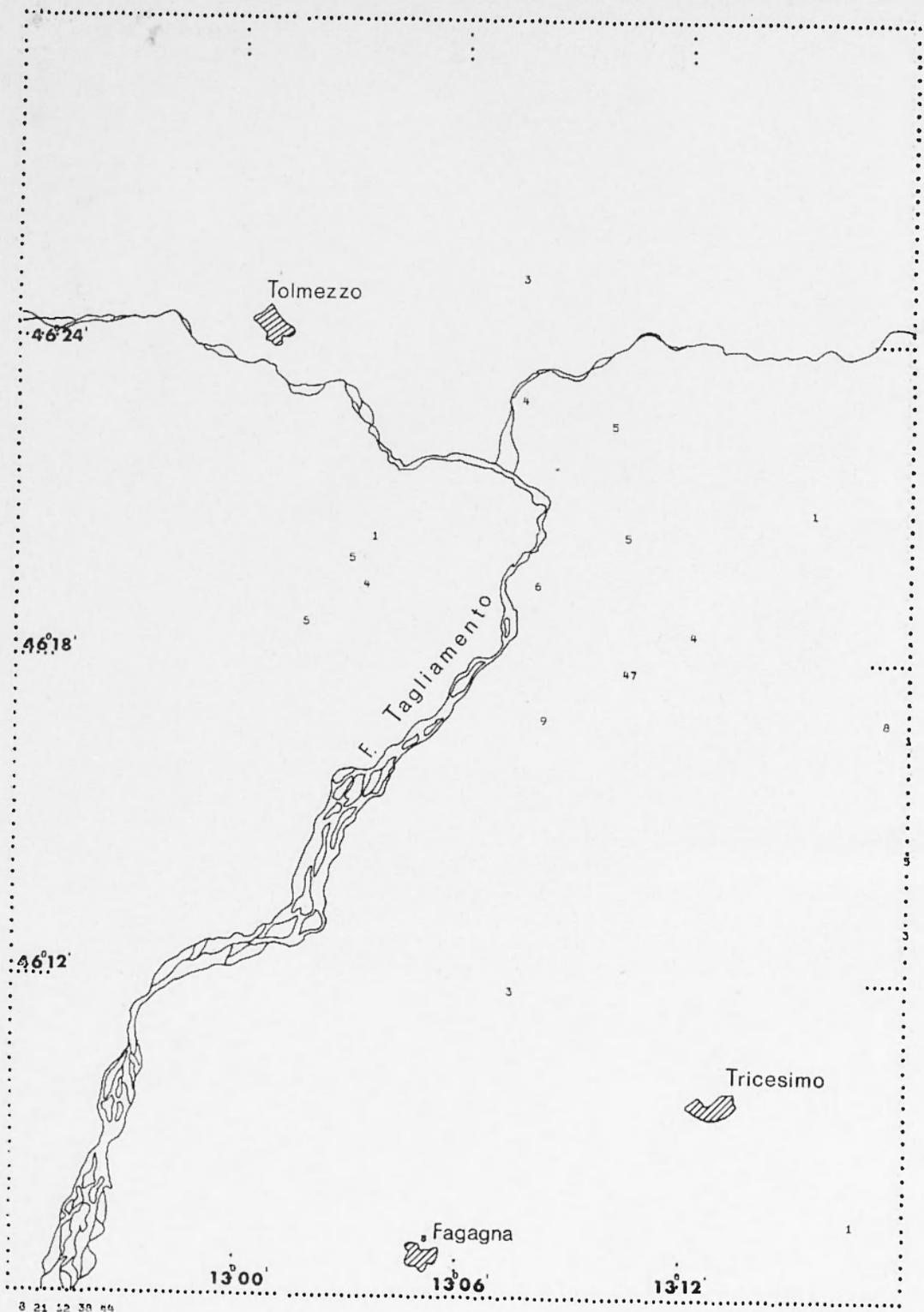
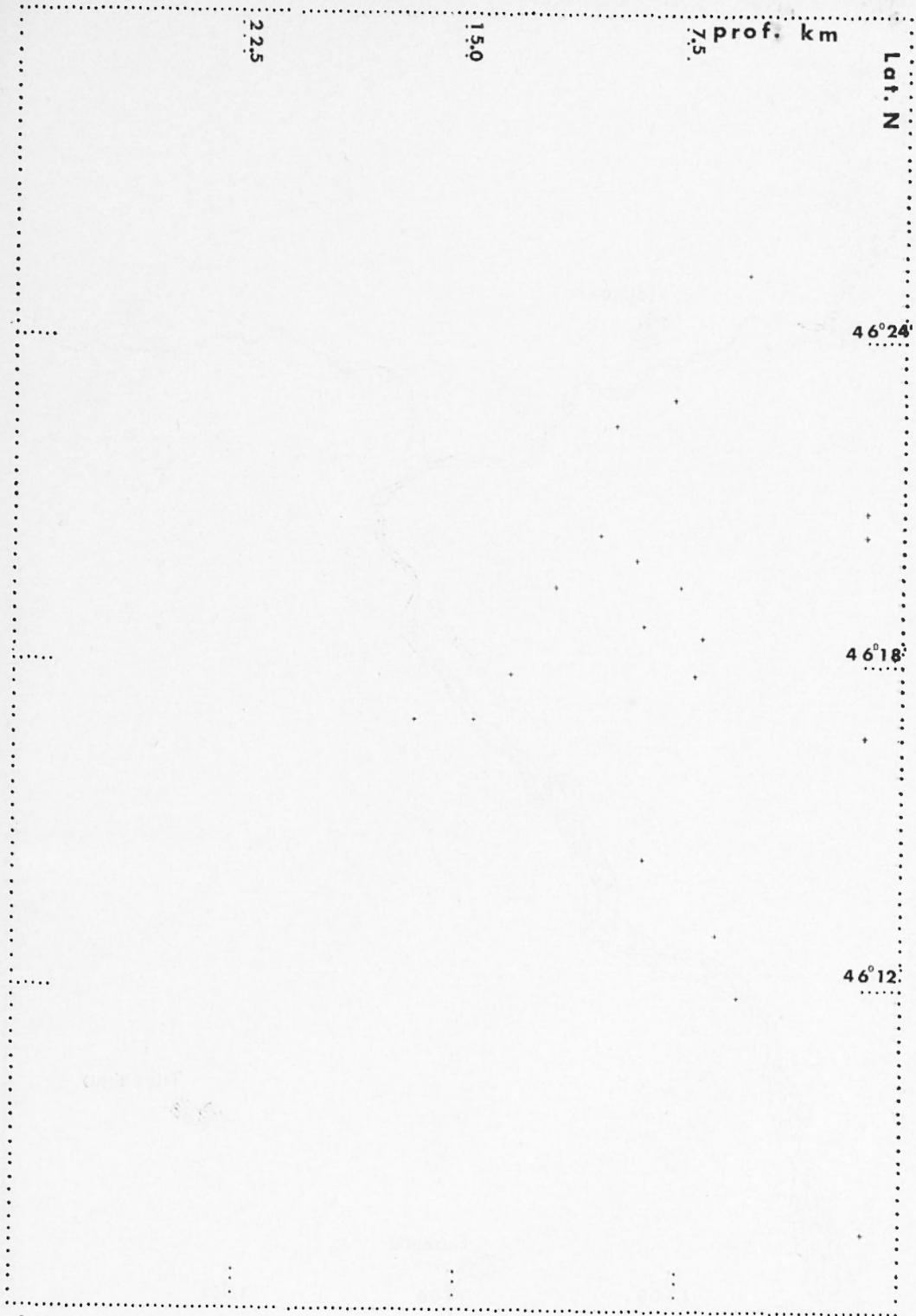


Fig. 16a-16b — From 8.8 17<sup>h</sup> 10<sup>m</sup> to 8.21 12<sup>h</sup> 38<sup>m</sup>.

The activity is very weak. Epicenters appear in the northern part of the zone under examination.



8 21 12 30 44

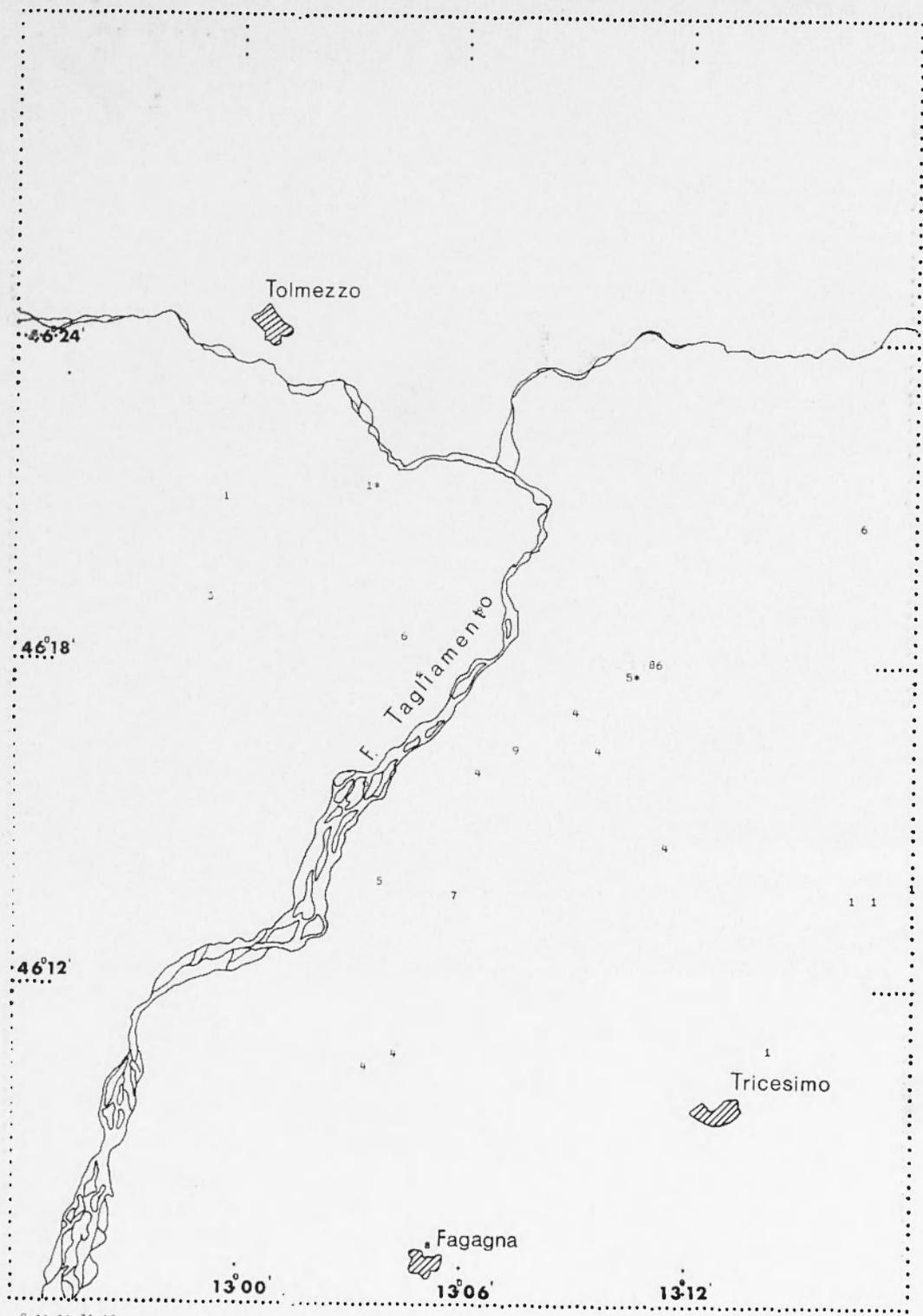
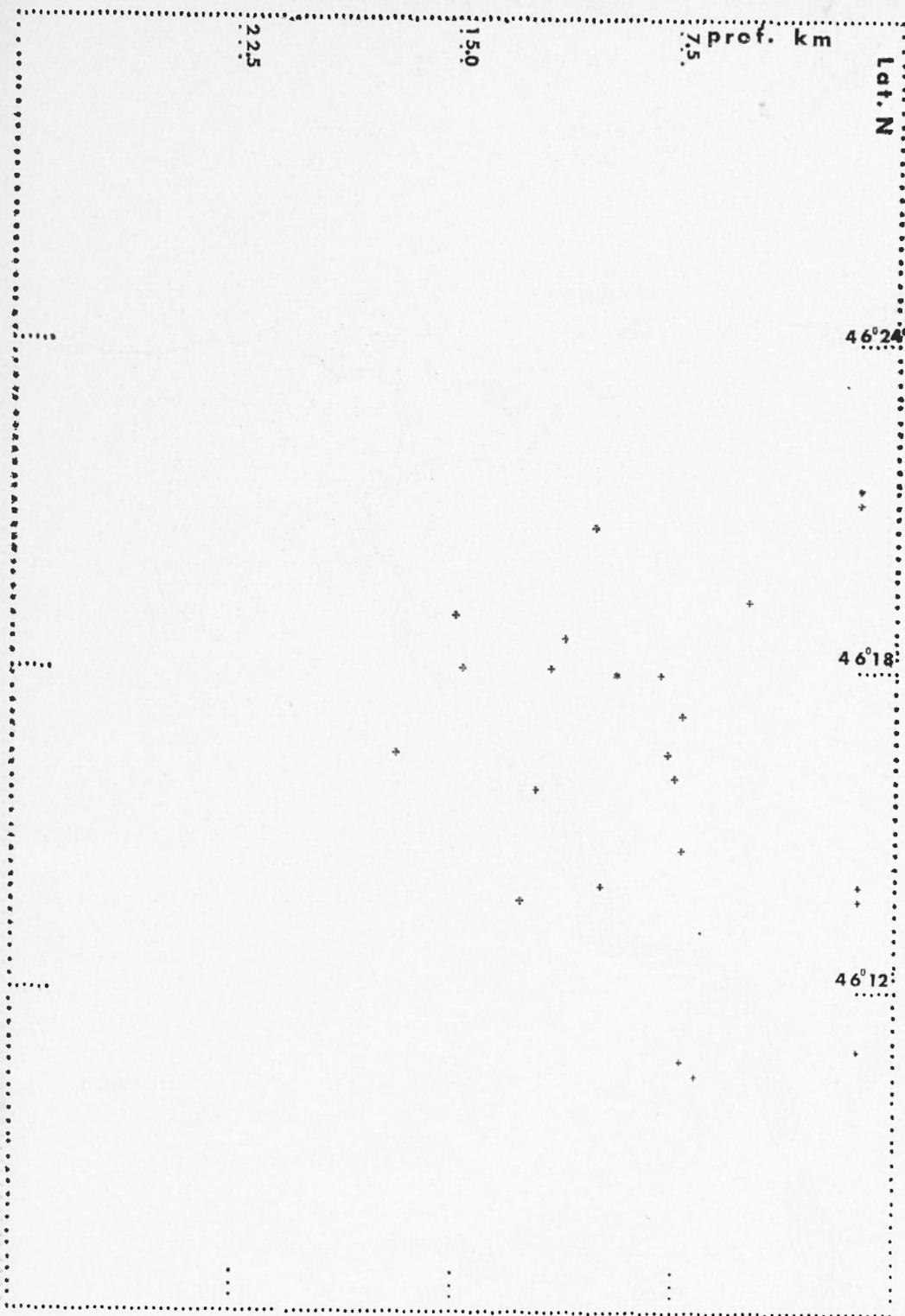


Fig. 17a-17b — From 8.21 12<sup>h</sup> 28<sup>m</sup> to 9.11 16<sup>h</sup> 31<sup>m</sup>.  
 The weak activity in the central sector ends with the shock of magnitude 5.5 on September 11, at 16<sup>h</sup> 31<sup>m</sup>.



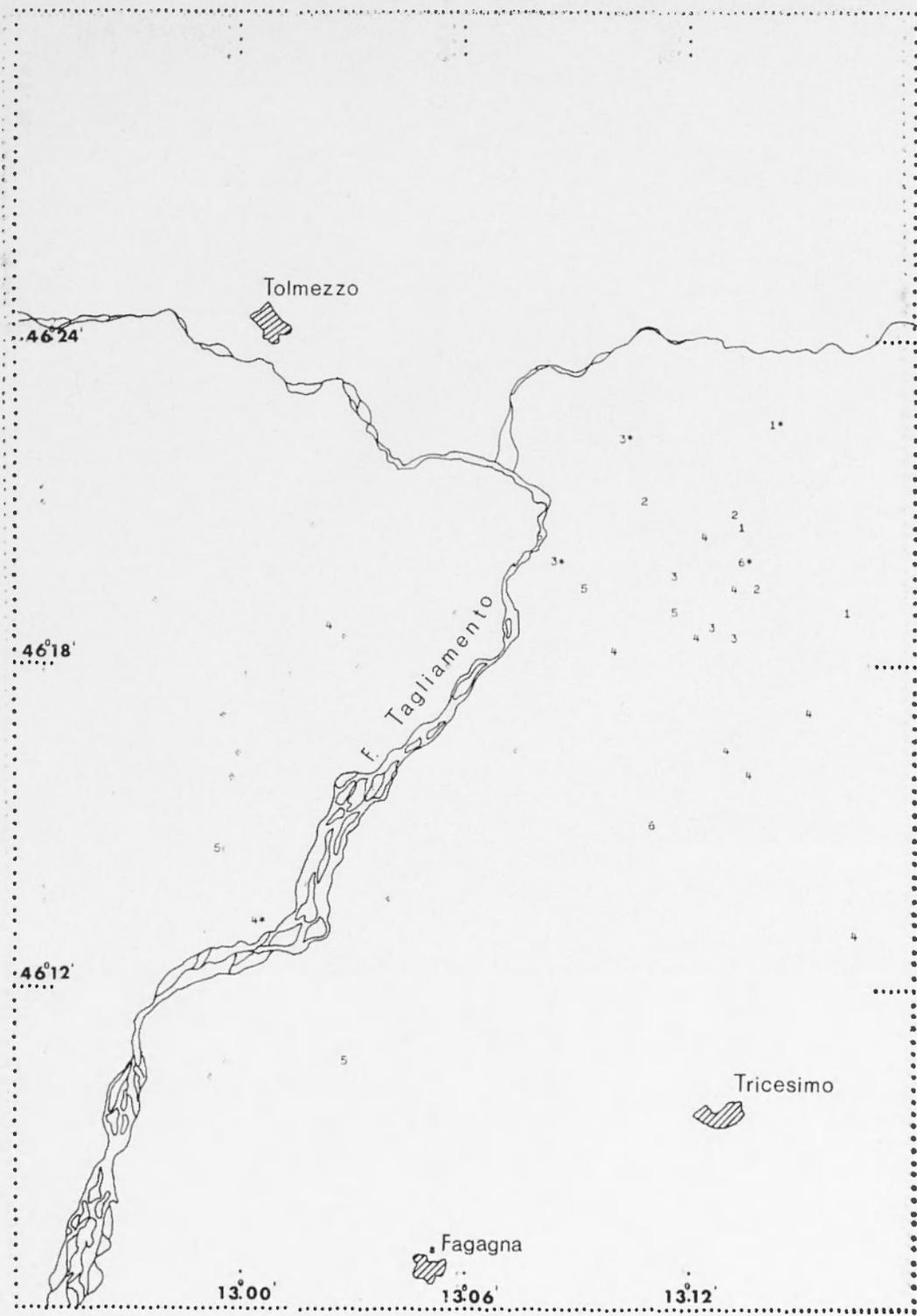
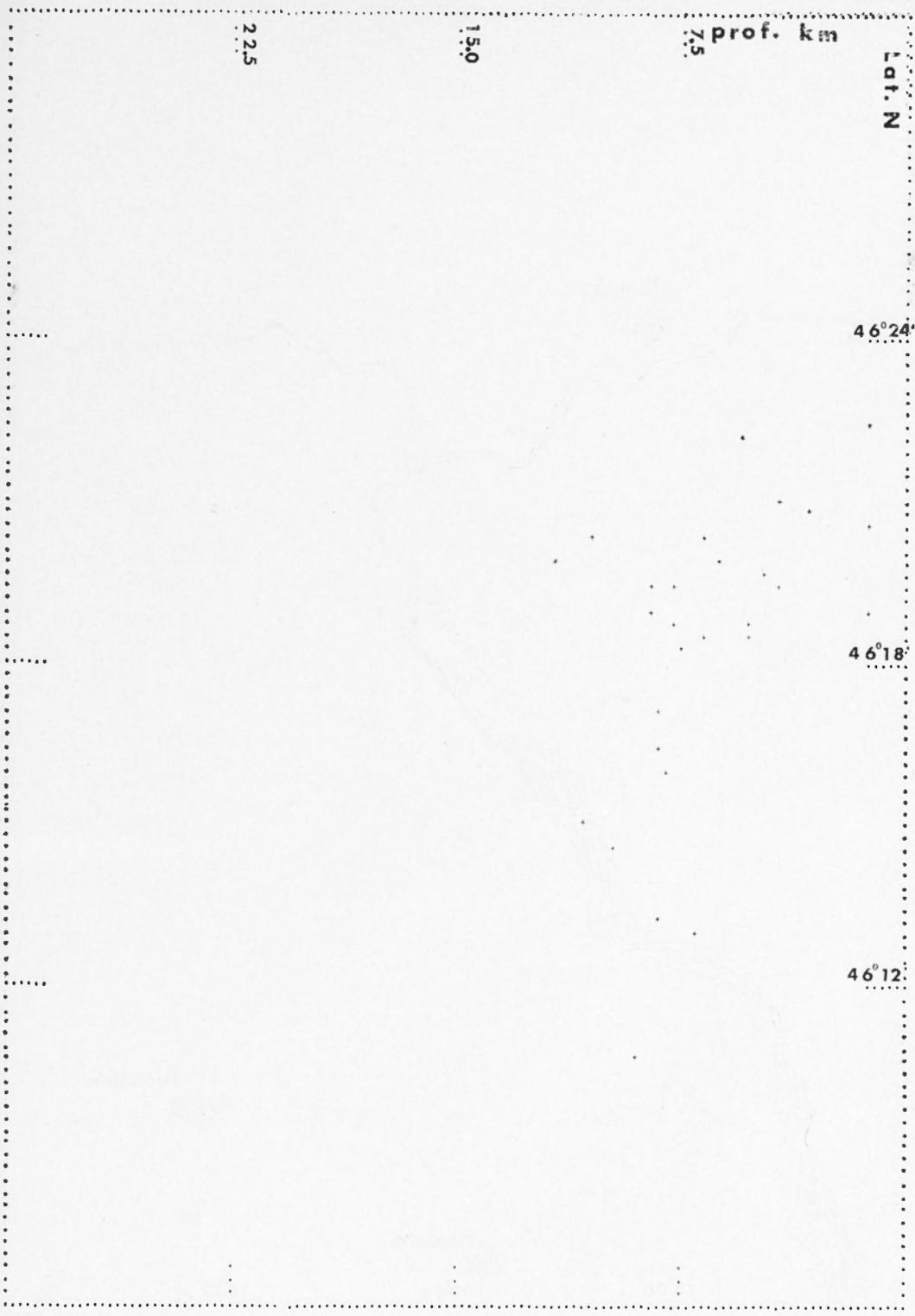


Fig. 18a-18b — From 9.11 16<sup>h</sup> 31<sup>m</sup> to 9.13 21<sup>h</sup> 25<sup>m</sup>.  
 The activity begins with a shock of magnitude 5.9 on September 11 at 16<sup>h</sup> 35<sup>m</sup>, forming a new seismic period, the concentration of which is evidently situated around the point with coordinates 46.32° and 13.20°.



7.5  
15.0  
22.5  
prof. km

Lat. N

46°24'

46°18'

46°12'

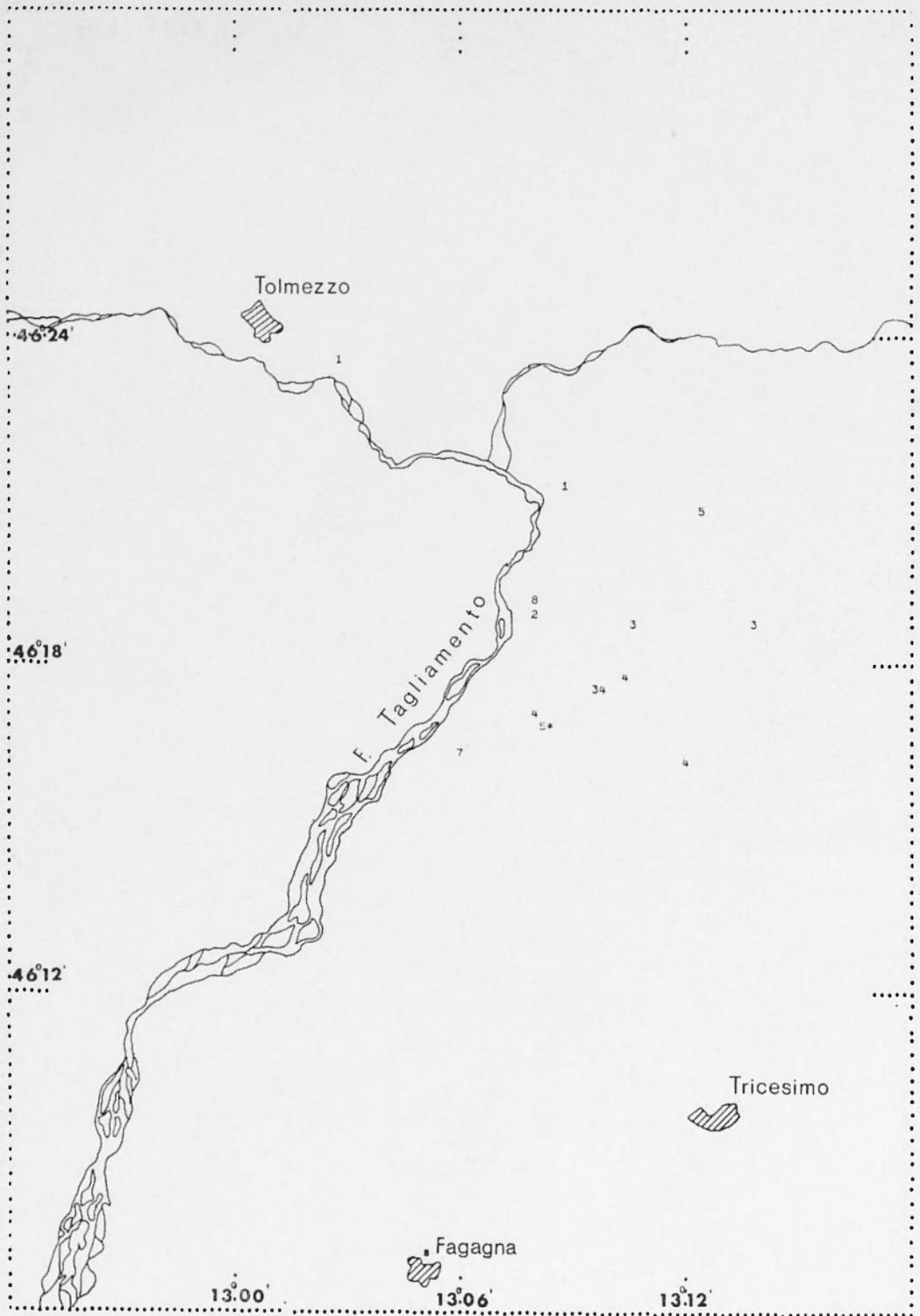
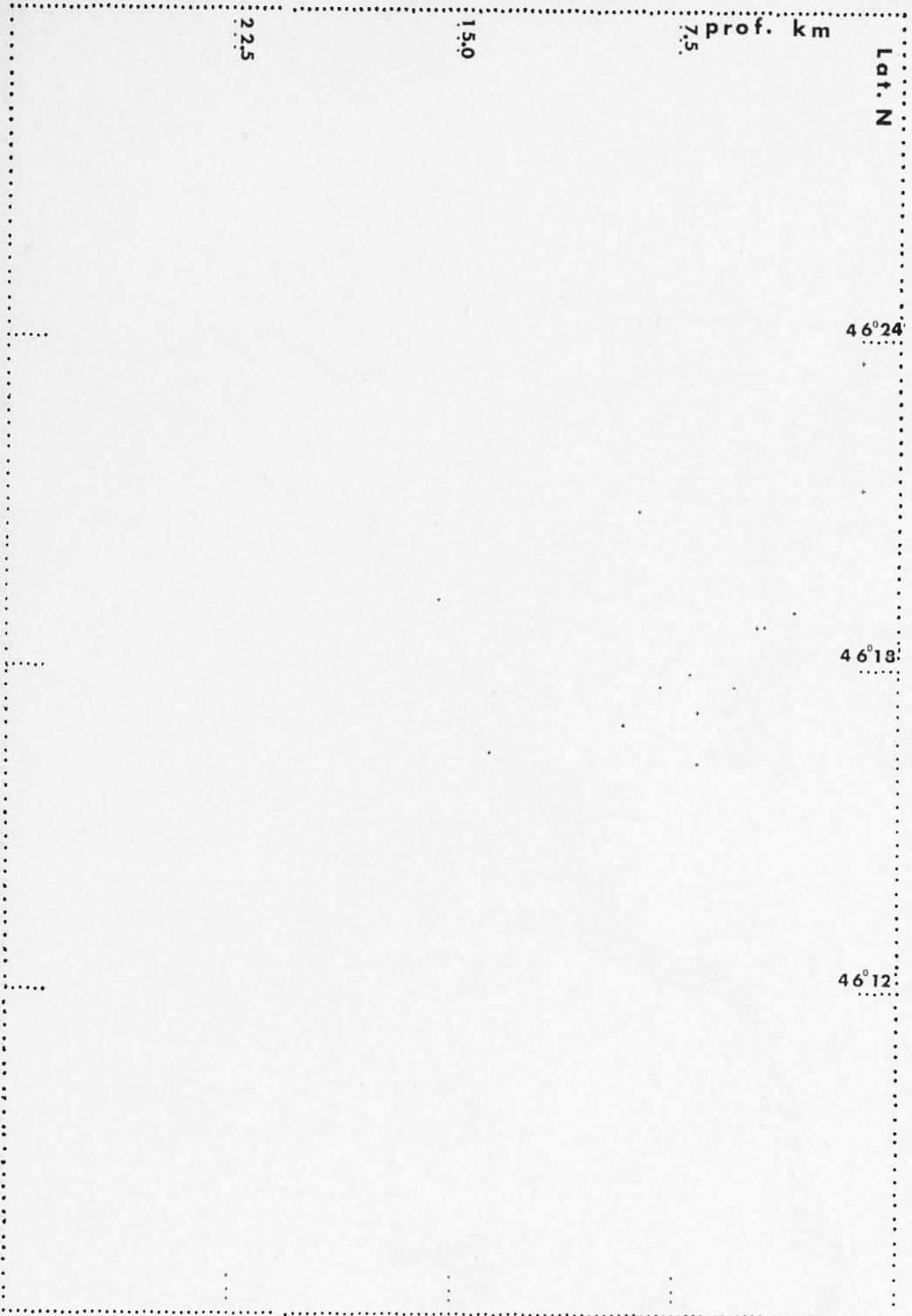


Fig. 19a-19b — From 9.13 21<sup>h</sup> 25<sup>m</sup> to 9.15 03<sup>h</sup> 55<sup>m</sup>.  
 The activity concentrates in the central part with low magnitude shocks, followed in the same zone by the event on August 15 at 03<sup>h</sup> 15<sup>m</sup>, of magnitude 6.1.



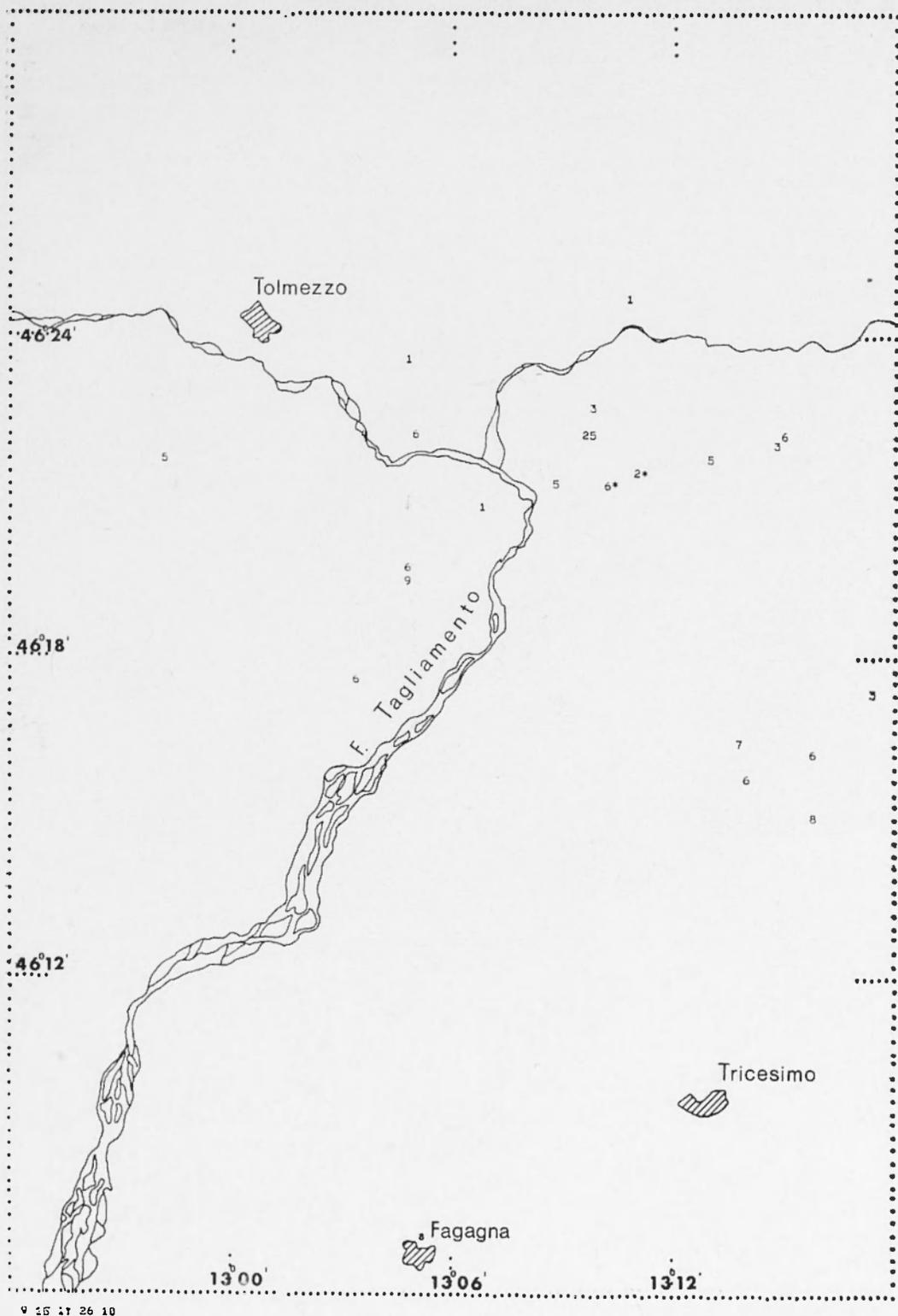


Fig. 20a-20b — From 9.15 03<sup>h</sup> 55<sup>m</sup> to 9.15 17<sup>h</sup> 26<sup>m</sup>.  
 The seismic activity, represented by a series of shocks of different magnitude, shows a sharp shifting towards N. A weak activity, concentrated to S-E, is also appearing.



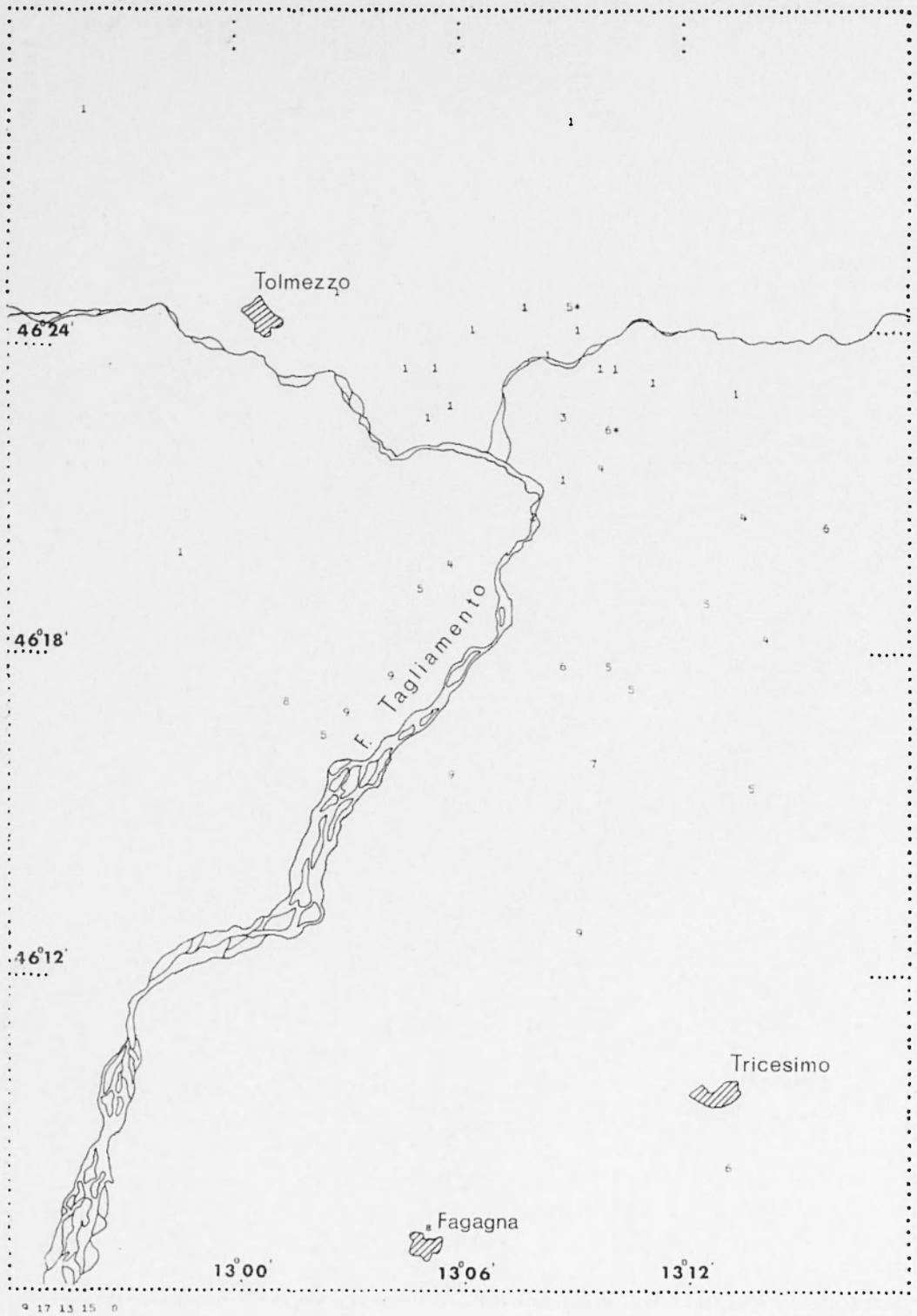
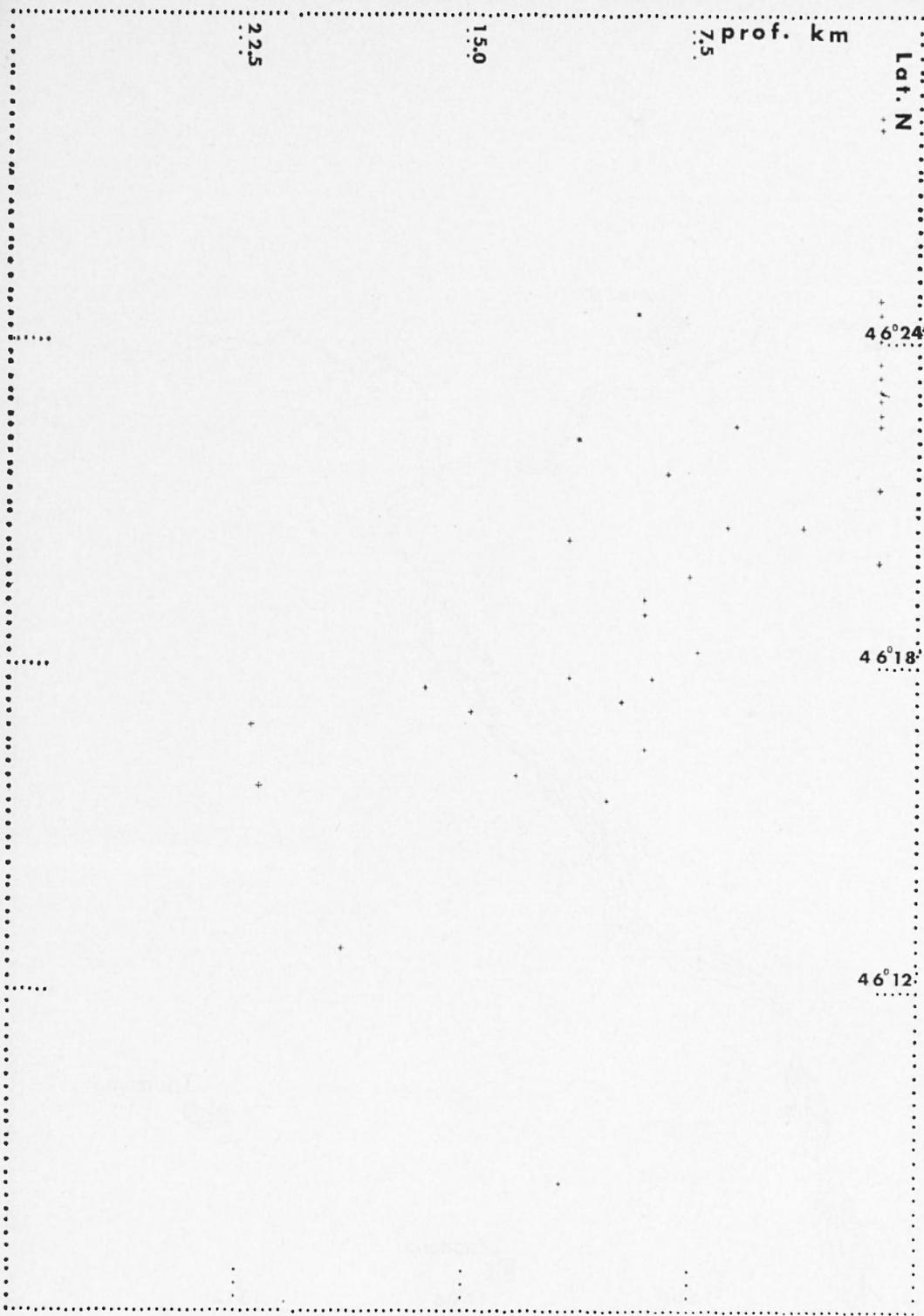
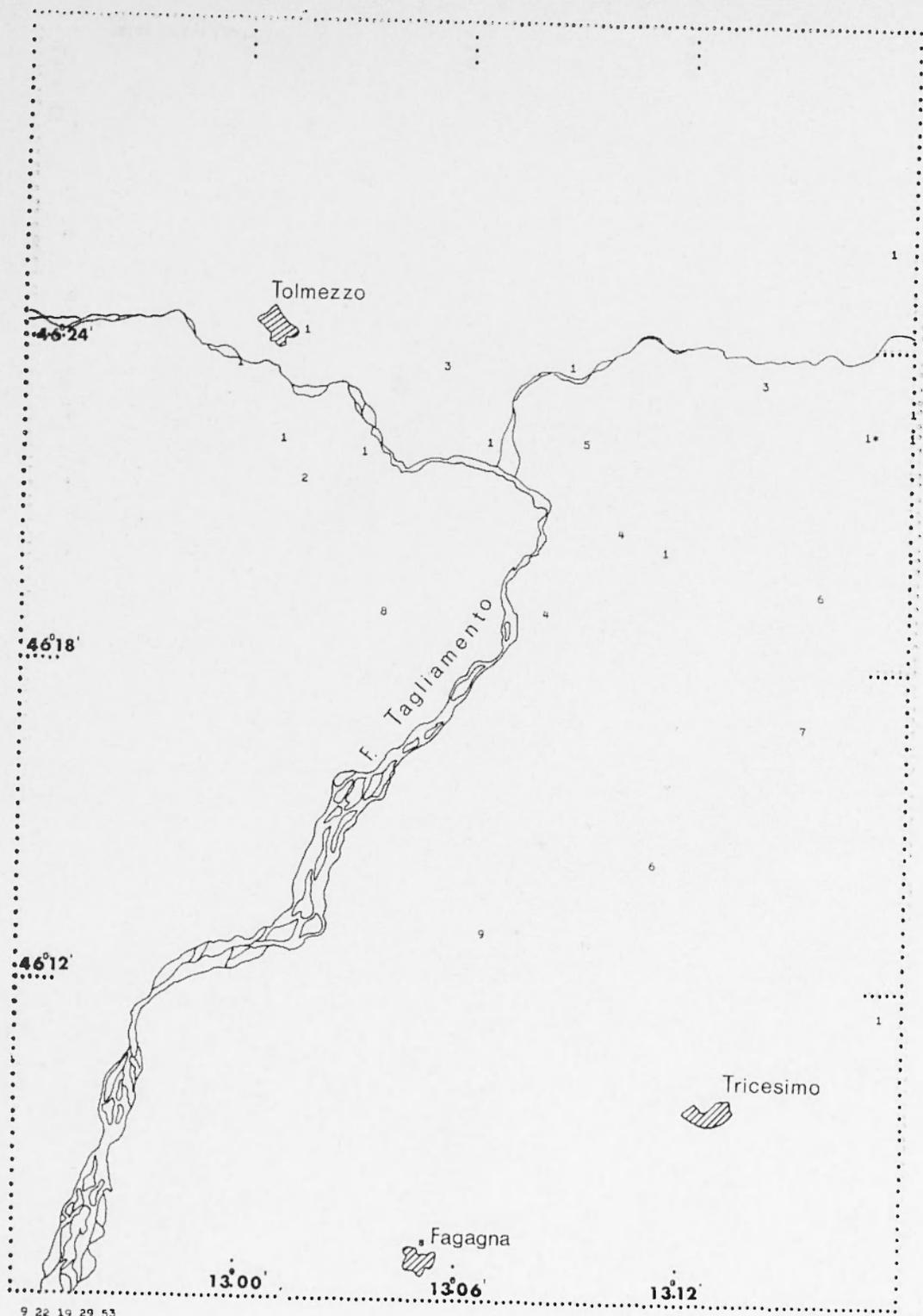


Fig. 21a-21b — From 9.17 17<sup>h</sup> 26<sup>m</sup> to 9.17 13<sup>h</sup> 15<sup>m</sup>.  
 The distribution of the epicenters, though diffused, shows a higher concentration in the N sector. The alignment of the epicenters goes towards S.

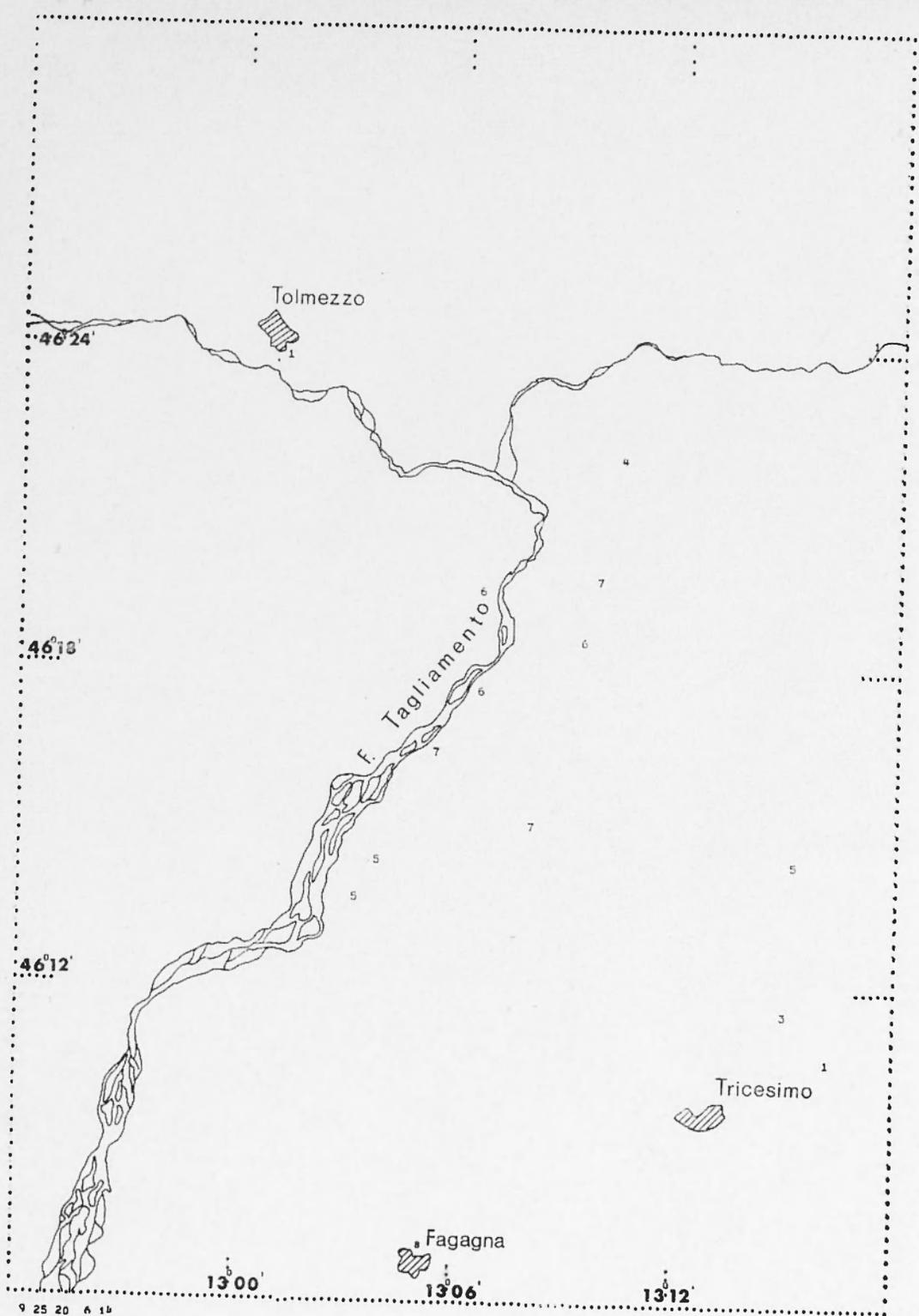




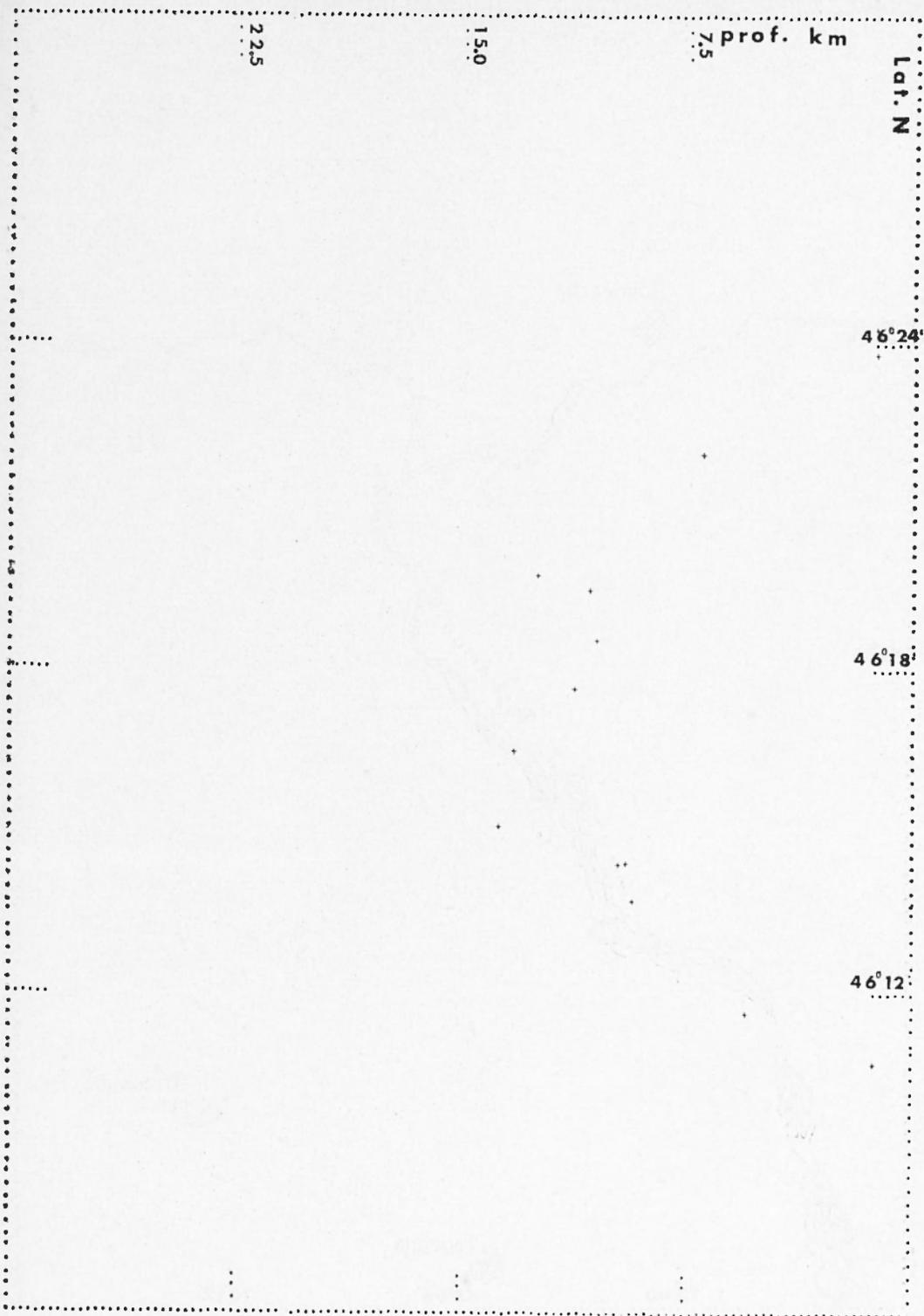
9 22 19 29 53

Figg. 22a-22b — From 9.17 13<sup>h</sup> 15<sup>m</sup> to 9.22 19<sup>h</sup> 29<sup>m</sup>.  
 The activity decreases. A dip towards S of the alignment of the epicenters  
 can still be observed.





Figg. 25a-25b — From 9.22 19<sup>h</sup> 29<sup>m</sup> to 9.25 20<sup>h</sup> 06<sup>m</sup>.  
 No high magnitude shocks can be seen. A modest activity is re-appearing to S-E.



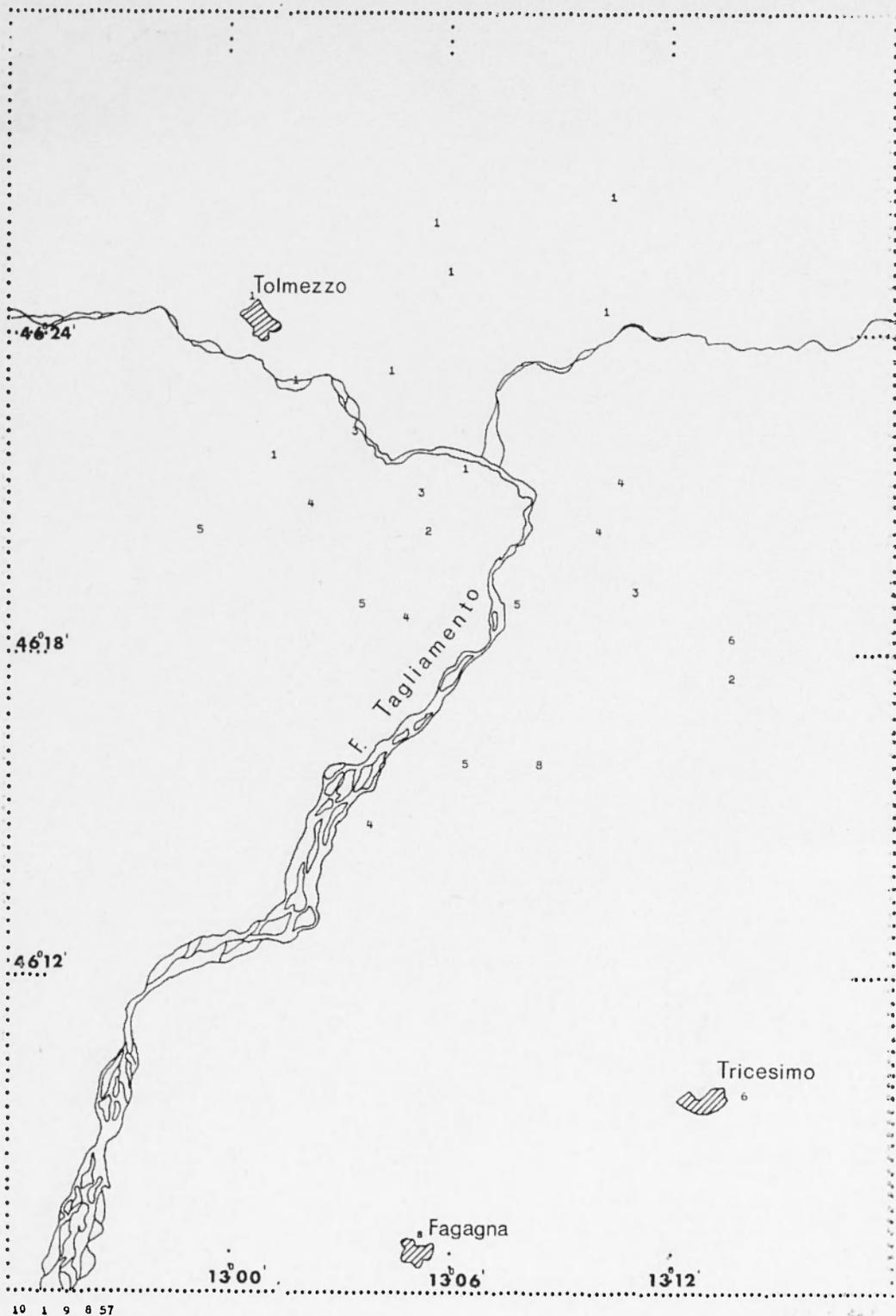
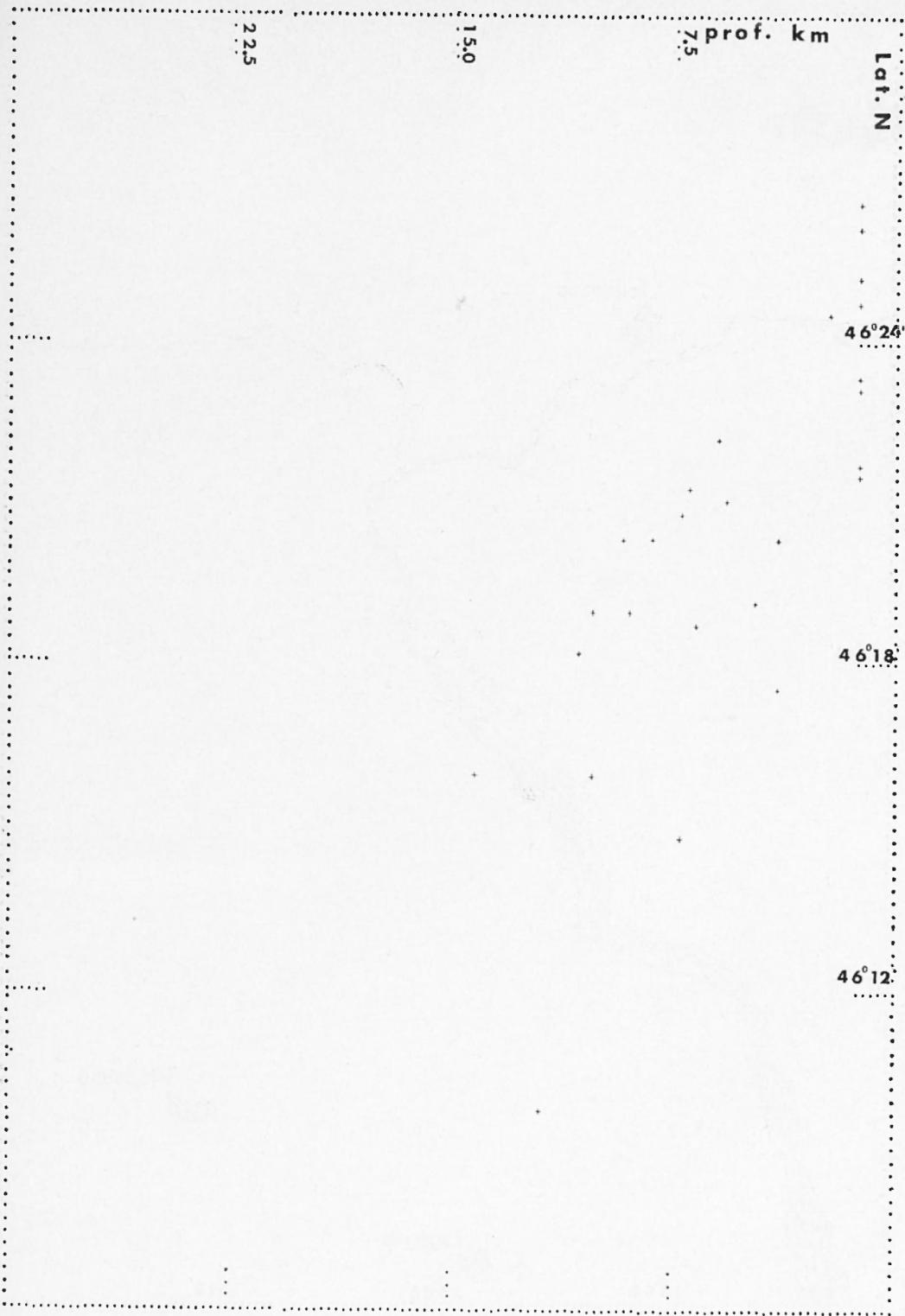


Fig. 24a-24b — From 9.25 20<sup>h</sup> 06<sup>m</sup> to 10.1 09<sup>h</sup> 08<sup>m</sup>.  
 The activity is scanty. The most important observation can be made for the alignment of the hypocenters still towards S.



10 1 9 8 57

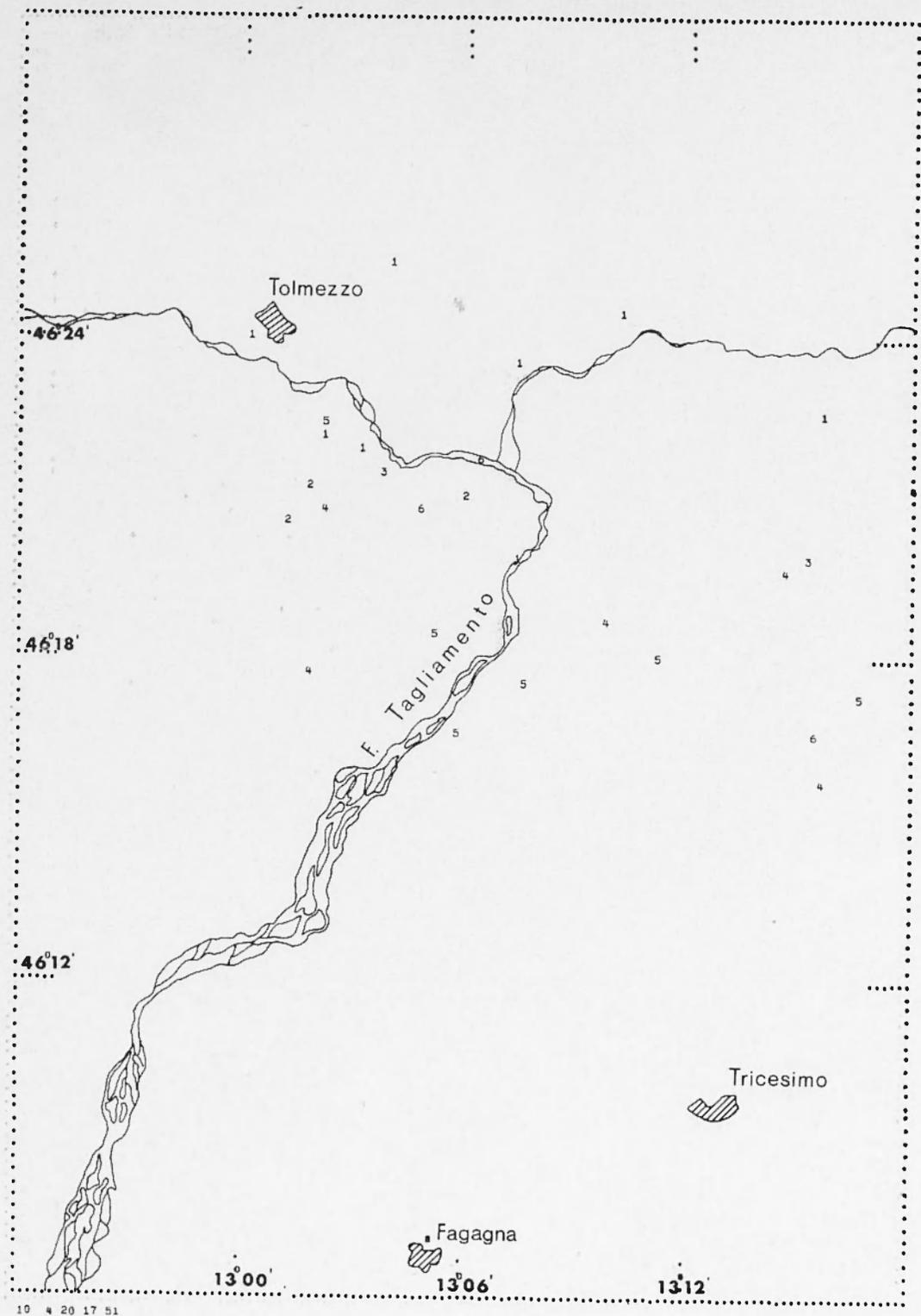
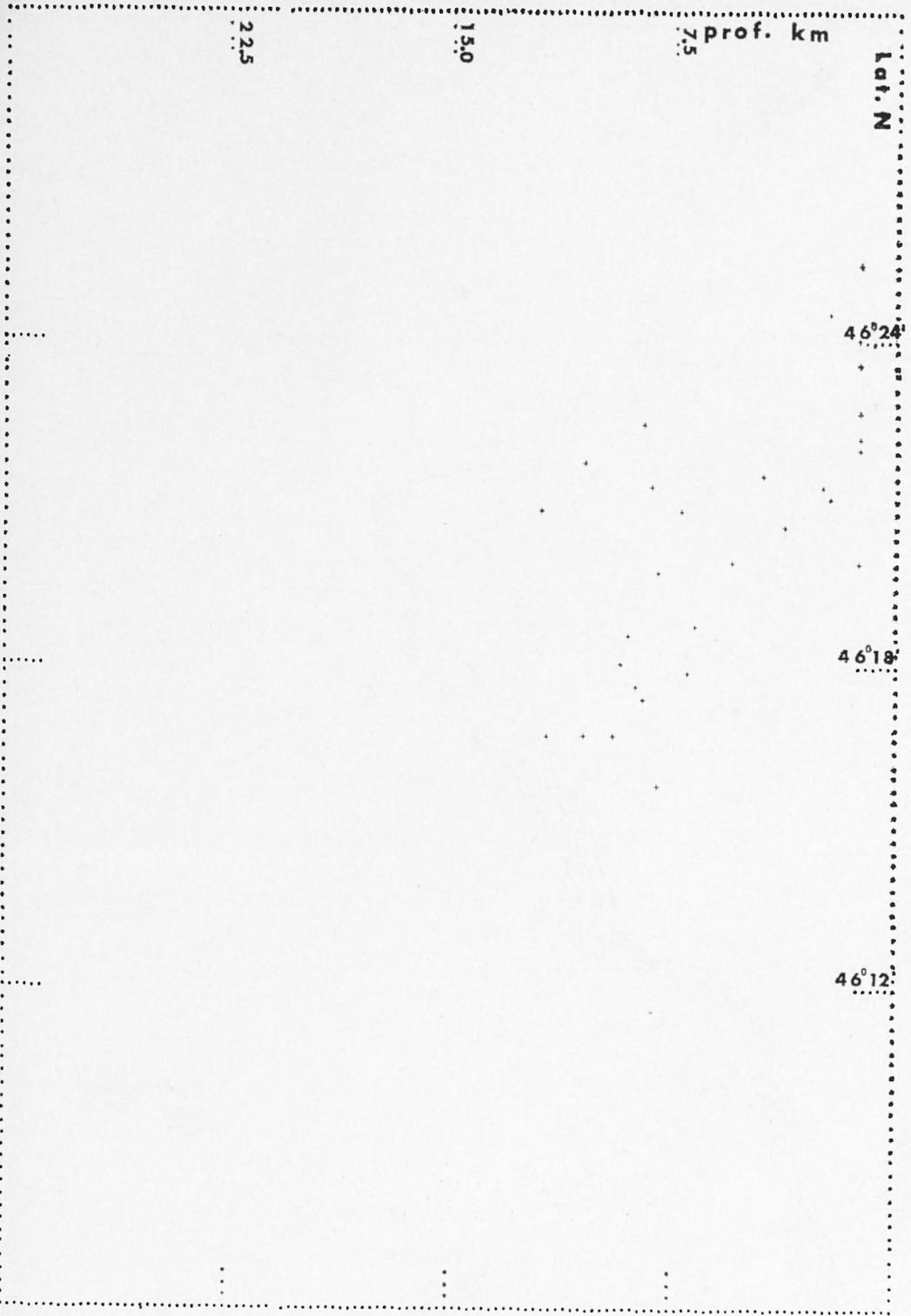


Fig. 25a-25b — From 10.1 09<sup>h</sup> 08<sup>m</sup> to 10.4 20<sup>h</sup> 17<sup>m</sup>.

In the northern sector the activity is continuing. Also the alignment of the hypocenters is still dipping towards S.



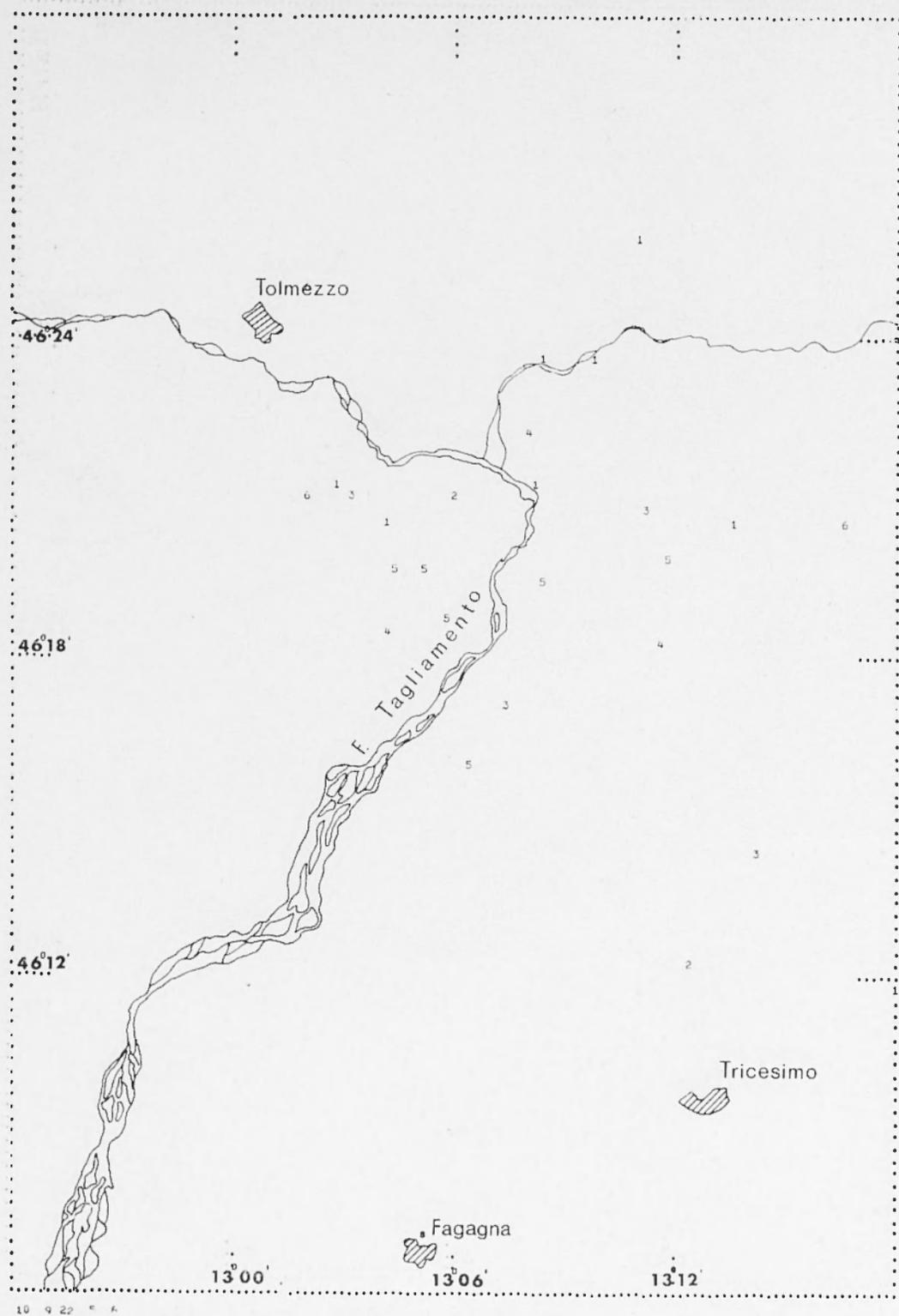
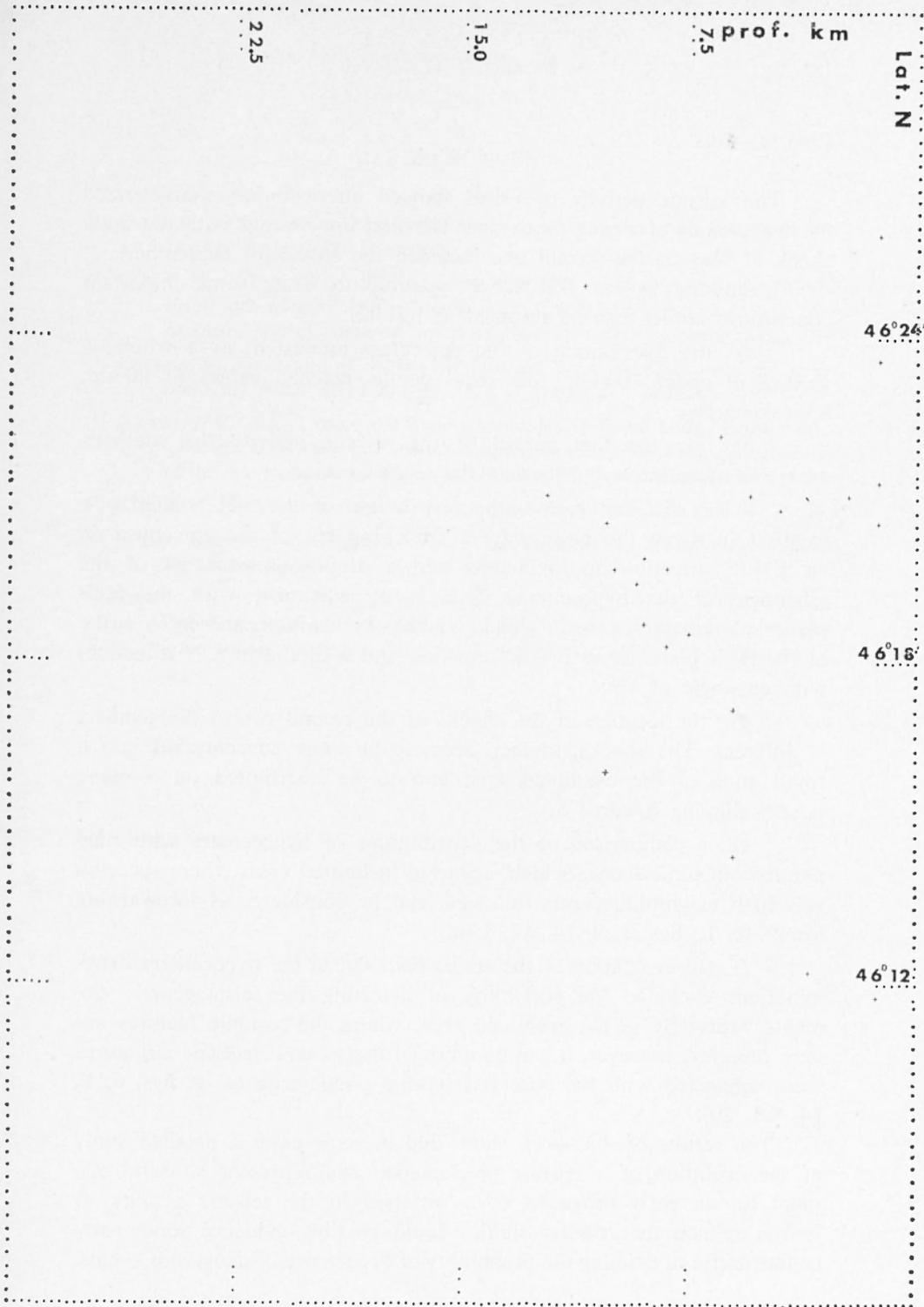


Fig. 26a-26b — From 10.4 20<sup>h</sup> 17<sup>m</sup> to 10.9 22<sup>h</sup> 05<sup>m</sup>.  
 The seismic activity has a low intensity, prevailing in the northern sector.



## CONCLUSIONS.

The seismic activity in Friuli showed an evolution characterized by two periods of strong paroxysm. The first one started with the main shock of May 6; the second one included the shocks of September.

Examining figures 1-26, it is possible to draw some important conclusions, which can be resumed as follows:

a) the distribution of the epicenters interested, as a whole, a surface of about 700 km<sup>2</sup>; the focal depths reached values of 20 km, approximately;

b) in the first period of the seismic activity the southern area was more interested than in the second period;

c) in the first period, the distribution of the foci pointed out, even if in a not too clear way, a stretching out of the epicenters in an *E-W* direction approximately and a dipping towards *N* of the alignment of the hypocenters. This is in agreement with the fault plane solution of the main shock which was characterized by a strike of the fault plane in an *E-W* direction and a dipping in a *N* direction, with an angle of 15°;

d) the location of the shocks of the second period (September) is different. The shocks, in fact, seem to be more concentrated into a small zone of the examined area and to be distributed on a plane steeply dipping towards *S*;

e) a comparison of the distributions of hypocenters with time pointed out some shocks which, occurred in limited areas where successively high magnitude events followed, can be considered as forewarning events (e. g., figs. 2, 9, 14, 17, 19);

f) an evaluation of the errors involved in the hypocenters determinations excluded the possibility of detecting the seismogenetic elements, especially in the examined area, where the tectonic features are very complex; however, it can be affirmed that several tectonic structures were connected with the observed seismic phenomena (e. g. figs. 6, 8, 14, 18, 20).

The results of this work show that in some cases a detailed study of the evolution of a seismic phenomenon can represent a useful element for an early prognosis of a variation in the seismic activity. It is our opinion that similar studies could possibly evidence some parameters useful in defining the probability of occurrence of dangerous events.

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